

*Circular 97*

STATE OF ILLINOIS  
DEPARTMENT OF REGISTRATION AND EDUCATION



*Groundwater Availability  
in Ford County*

by JAMES P. GIBB

ILLINOIS STATE WATER SURVEY  
URBANA  
1970

### *HOW TO USE THESE MATERIALS*

This circular provides a summary of all available information on water wells and groundwater conditions in Ford County. You can use these materials to find the possibilities of obtaining a dependable water supply at any location in the county.

First you will need the legal description (township, range, section, and portion of section) of your farm, home, or other location of interest. Then follow these steps.

- 1) Turn to appendix A and find your location (township, range, section) in the list of well numbers for existing wells.
- 2) Examine the records of all the wells in your section and in the adjoining sections. The depths of these wells, the water-bearing formations they tap, their nonpumping water levels, and other information give an immediate picture of existing water supplies, which is one indication of what is possible in your location of interest.
- 3) Continue to appendix B for the chemical quality of water in the existing wells in your location, which indicates the quality you may usually expect for the different depths and sources.
- 4) Now turn to the maps in the text which illustrate all of this information to show the possibilities for dependable wells throughout the county. Figures 6, 7, and 8 illustrate information for relatively shallow wells in the upper water-bearing deposits, and figures 9, 10, and 11 give information for deeper wells in the lower deposits.

An example of actual use of these materials for a specific location is presented on page 38, preceding appendix A.



## CONTENTS

	Page
Summary. . . . .	.1
Introduction. . . . .	.2
Geology. . . . .	.4
Groundwater. . . . .	.6
Occurrence and movement. . . . .	.6
Chemical character. . . . .	.10
Temperature. . . . .	.12
Aquifers.. . . .	.13
Upper                glacial                deposits.....	14
Development                and                availability.....	14
Chemical character. . . . .	.17
Lower glacial deposits. . . . .	.19
Development and availability. . . . .	.19
Chemical character. . . . .	.21
Bedrock formations. . . . .	.23
Development and availability. . . . .	.23
Chemical character. . . . .	.24
Wells. . . . .	.25
Types and drilling methods. . . . .	.25
Construction features. . . . .	.27
Casing. . . . .	.27
Screening. . . . .	.28
Gravel packing. . . . .	.28
Grouting. . . . .	.29
Disinfection . . . . .	.29
Methods of pumping water. . . . .	.30
Summary of major water supplies. . . . .	.30
Example use of materials. . . . .	.38
Appendix A. Records of wells. . . . .	.39
Appendix B. Chemical quality of groundwater. . . . .	.60
Selected references. . . . .	.65

## ILLUSTRATIONS

Figure	Page
1 Location of Ford County . . . . .	.3
2 Geologic maps of Ford County. . . . .	.5
3 Generalized graphic logs of glacial deposits and bedrock formations. . . . .	.7
4 Cycle of water movement (a) and generalized movement of water in Ford County (b). . . . .	.8
5 Water table contour map of Ford County. . . . .	.9
6 Probable maximum depth of wells in the upper glacial deposits. . . . .	.15
7 Nonpumping water levels (a) and water level elevations (b) in drilled wells tapping the upper glacial deposits. . . . .	.16
8 Selected chemical analysis (a), probable total dissolved minerals (b), and hardness (c) of water from the upper glacial deposits. . . . .	.18
9 Probable maximum depth of wells in the lower glacial deposits. . . . .	.20
10 Nonpumping water levels (a) and water level elevations (b) in wells tapping the lower glacial deposits. . . . .	.21
11 Selected chemical analysis (a), probable total dissolved minerals (b), and hardness (c) of water from the lower glacial deposits. . . . .	.22
12 Construction features used in Ford County wells . . . .	.27
13 Recommended construction features for large-diameter dug wells. . . . .	.28

## TABLES

Table	Page
1 Major dissolved elements and substances in groundwater in Ford County. . . . .	.11
2 Recommended chlorine dosages for well disinfection . . .	.30

## GROUNDWATER AVAILABILITY IN FORD COUNTY

by James P. Gibb

### SUMMARY

Groundwater in Ford County normally can be obtained from one of two primary water-bearing units within the glacial drift or from the underlying bedrock formations. The drift deposits consist of the Wisconsin, Illinoian, and Kansan age glacial materials underlain by Pennsylvanian, Mississippian, Devonian, and Silurian age bedrock formations.

Glacial deposits of Wisconsin age provide approximately 64 percent of the county's current water supply. The maximum recorded depth of wells tapping this upper aquifer system is about 240 feet. Individual yields of domestic wells finished in these deposits generally are limited by pump capacities of 5 or 10 gallons per minute (gpm). Municipal and industrial supply wells capable of producing in excess of 400 gpm have also been developed in this water-bearing unit. Adequate water for farm and domestic use is usually obtainable from the Wisconsin deposits throughout the county.

Deeper lying Illinoian and Kansan age glacial deposits also are tapped for domestic and municipal water supplies. Wells tapping this lower aquifer system may be as deep as 340 feet and yield in excess of 1000 gpm. The larger groundwater supplies have been developed from morainal deposits or glacial fill materials contained in the three buried bedrock valleys crossing the county. Groundwater from the glacial deposits is hard and normally contains objectionable concentrations of iron.

The upper bedrock formations have been tapped by a few wells in the northern portion of the county. These wells approach 300 feet in depth and yield from 5 to 150 gpm. Throughout the remainder of the county the bedrock has not been extensively tapped because of the productivity of the overlying drift. The quality of water from the bedrock is generally poor and at depths below 300 or 400 feet may become too "salty" for most uses.

An estimated 2.2 million gallons of water is pumped from the aquifers of Ford County each day to satisfy industrial, municipal, domestic, and rural needs. A much larger quantity of water, perhaps as much as 51 million gallons a day, could probably be withdrawn without overdevelopment. Maps and tables indicating the probable maximum depths of wells, water levels, chemical quality, and general groundwater conditions for each water-bearing unit at specific locations are presented to serve as a guide in the development and utilization of the groundwater resources of Ford County.

## INTRODUCTION

More than 500 requests for information concerning groundwater conditions in specific locations of Illinois are answered yearly by the Illinois State Water Survey. Approximately 40 percent of these requests are from individuals seeking advice on locating, developing, or treating home or farm groundwater supplies.

Many of these requests are answered with letter-type reports prepared jointly by the State Water Survey and State Geological Survey from available geohydrologic data in their basic record files. These reports, containing pertinent information on groundwater and geologic conditions at a specific site, permit meaningful appraisals for well construction which have saved considerable time, effort, and money in many cases. However, several thousand wells are constructed each year without the use of such information. If comprehensive summaries of groundwater conditions were available for all possible sites, great savings could result. This report presents such a summary for Ford County, where fairly complex groundwater conditions exist.

Ford County is located in the east-central part of the state (figure 1). It encompasses an area of 488 square miles and is mainly cultivated land. According to the 1960 census, the county has a population of 16,606 with 10,967 of the residents living in incorporated cities and villages. The county seat and largest city, Paxton, has a population of 4370.

The economy of the county is largely dependent on the production of farm crops such as corn, soybeans, and wheat and agriculturally oriented industry. Among the larger industries in Ford County are Central Soya Company, Inc., Stokely Van Camp Canneries, and M S W Gear Company all located in Gibson City.

Drainage from the county is predominantly to the south and west, but a small area southeast of Roberts drains to the northeast. Headwaters for the Iroquois, Mackinaw, North Fork Vermilion, and Vermilion Rivers are located in the upland areas of the county. The Sangamon River originating in neighboring McLean County flows through the extreme southwestern corner of the county.

Information on the streams and rivers in Ford County is not readily available in published form. Agencies in Illinois that may have data on file for these streams include: Illinois State Water Survey, Urbana; Illinois Division of Waterways, Springfield; and the U. S. Geological Survey, Champaign.

The rivers, streams, and creeks in Ford County are not sufficiently large to be considered as a source of water for most uses. All known water supplies in the county are from groundwater sources.

This report summarizes groundwater conditions in Ford County including pertinent geologic factors, occurrence and movement, temperature and chemical quality, and well development. Appendix A contains available records of all known wells, and appendix B lists the results of chemical analyses of water from all wells sampled.

This study is part of a continuing program of water-resource investigations being conducted by the Illinois State Water Survey under the general direction of Dr. William C. Ackermann, Chief, and H. F. Smith, Head of the Hydrology Section. The report was prepared under the direct guidance of William H. Walker. Grateful acknowledgment is made to the many well drillers, engineers, and public officials

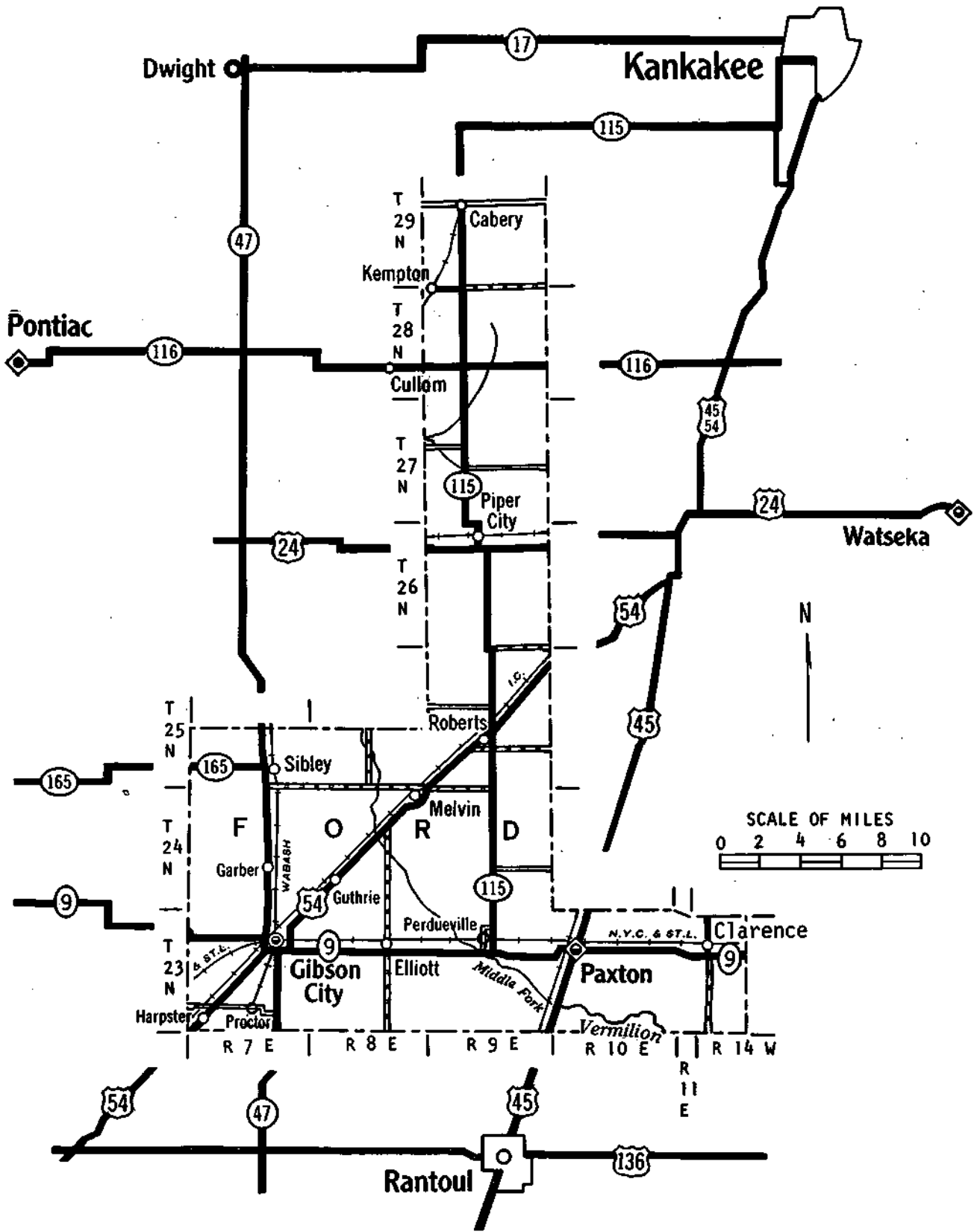


Figure 1. Location of Ford County



who provided invaluable information for use in this report. Special thanks is given to Verne Bear, science teacher from Roberts, who conducted the house to house inventory of wells in Ford County during the summer of 1967 which provided much of the basic data used in this report. Mrs. Dorothy Woller tabulated the well data and typed the manuscript, and John W. Brother, Jr., prepared the illustrations.

## GEOLOGY

The geology of Ford County is summarized in general terms in State Geological Survey Circular 248, "Groundwater Geology in East-Central Illinois." The following brief discussion of geologic conditions in the county is taken largely from that publication. For a more detailed definition of the geology in this portion of the state, the reader is referred to the State Geological Survey which is located on the University of Illinois campus, Urbana.

Glacial deposits blanket all of Ford County resulting in a relatively level plain broken only by isolated knobs, stream valleys, and long ridges formed at the front of the glaciers (end moraines). The glacial deposits include those of Wisconsinan (upper deposits), Illinoian, and Kansan (lower deposits) age. Information from wells and rock exposures indicates that the topography of the county has been shaped principally by ice and running water. Features produced by ice were developed long ago when the glaciers, nourished by snow accumulation in Canada, several times advanced across Ford County and melted away leaving vast quantities of rock debris. In front of the ice, sediment-laden meltwaters escaped down valleys, partially filling them with outwash deposits of sorted sand, gravel, and finer materials. Thick extensive till sheets of unsorted clay, silt, sand, and pebbles also were laid down under the advancing ice or dumped into place during melting. The thickness of the glacial deposits varies from about 50 to more than 400 feet, the thicker sections being associated with the bedrock valleys and the morainic ridge just northwest of Gibson City (figures 2a, b, and c).

The Wisconsinan deposits form the present-day land surface of Ford County. Running water continues to modify this surface by cutting into the land, carrying away soil and rock particles, and depositing the debris in river bottoms. This modification is a small-scale version of the changes made on the bedrock surface by the glacial meltwaters. The Wisconsinan deposits in the study area consist primarily of till materials interspersed with somewhat discontinuous pockets and lenses of sand and gravel.

Underlying the Wisconsinan materials from the Piper City area southward are the Illinoian age deposits. The materials in these deposits are more uniform and occur as relatively impermeable till units interbedded with continuous layers of sand and gravel. The Kansan age deposits occupy the basal position in the drift section and consist primarily of permeable sands and gravels, particularly in the bedrock channels.

The bedrock formations in Ford County are layers of consolidated rocks of Pennsylvanian, Mississippian, Silurian, and Devonian geologic age (*see figure 2d for areal distribution*). These rocks consist of beds of shale, sandstone, limestone, and dolomite arranged one upon the other; the top surface of these rocks is called the bedrock surface (*see figure 2c for surface configuration*). Origin-

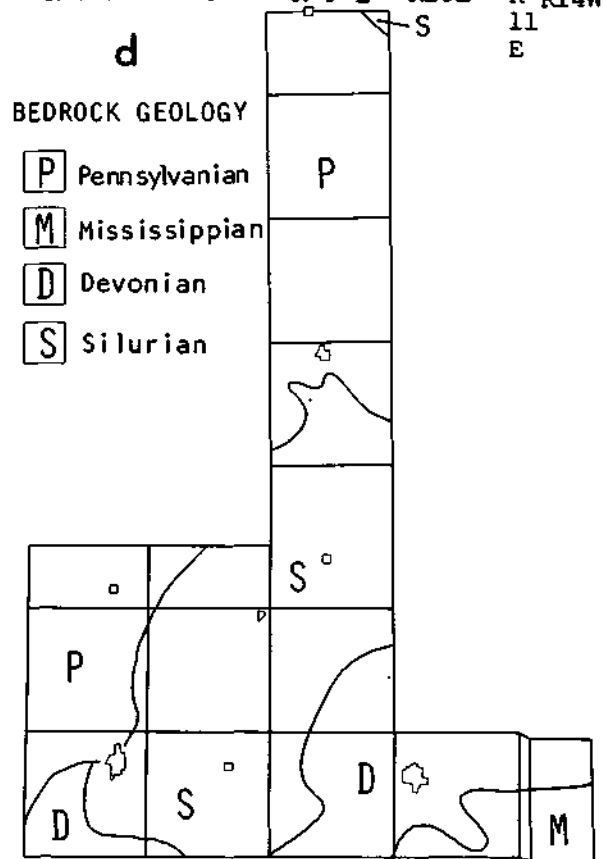
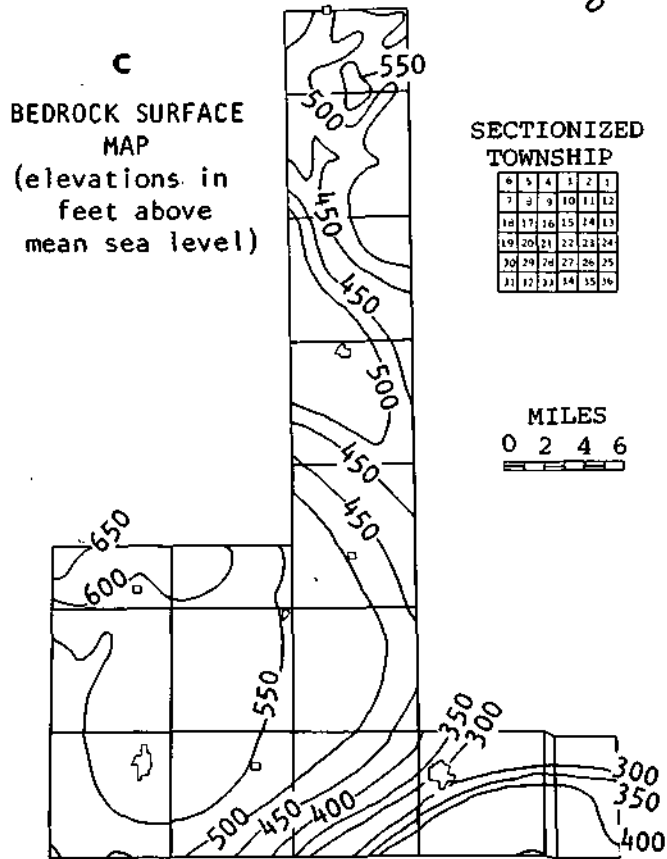
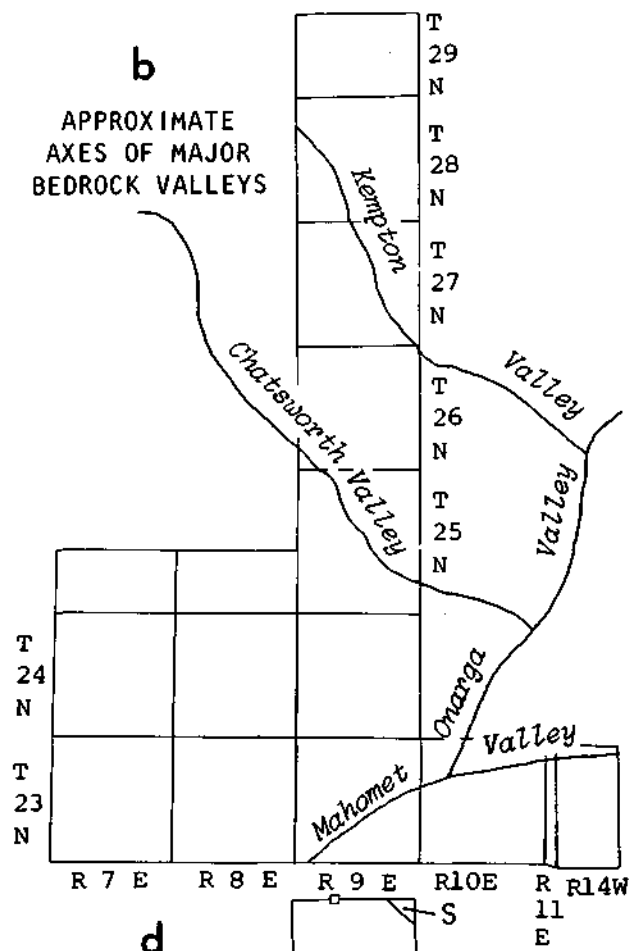
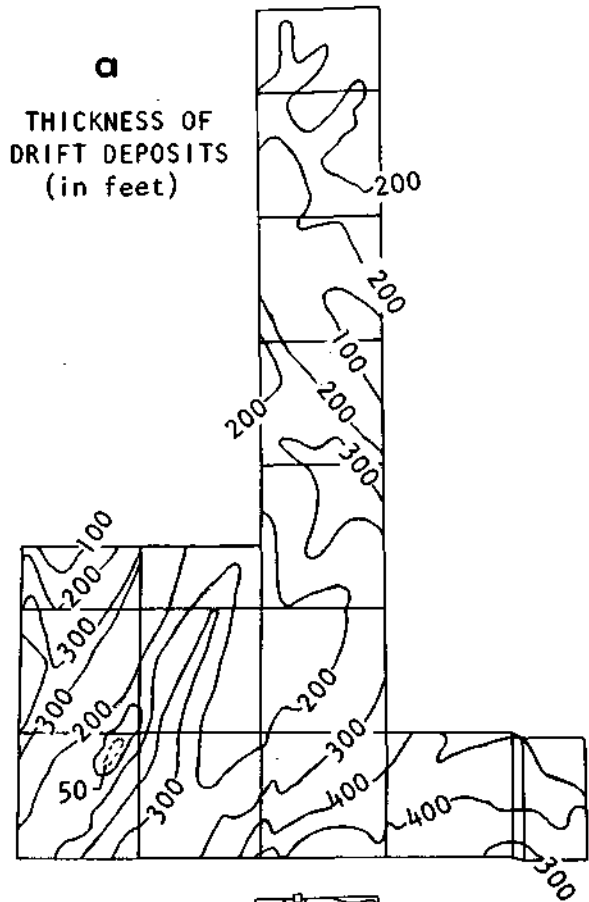


Figure 2. Geologic maps of Ford County

nally the bedrock formations were unconsolidated materials, deposited over many years as sediments in shallow seas or bordering marshes. They were then buried and hardened into solid rock during the several million years after the seas retreated from the area.

Erosion of the bedrock was not uniform through the county. In areas where soft shales and sandstone formations were exposed to weathering, valleys were formed by water and ice action. Hard sandstone and limestone formations in other areas resisted erosion and remained to form ridges and hills on the bedrock surface. Some of the old bedrock valleys coincide with present-day stream valleys, but some are partially or even completely buried by the glacial deposits so that there is little or no surface evidence of their presence. The principal buried valley system in Ford County occurs as part of the ancient Teays River, a master preglacial stream which headed in the Blue Ridge Mountains in North Carolina and discharged into the ancient Mississippi River west of Delavan in Tazewell County. The Teays River Valley, or Mahomet Valley as it is known in Illinois, enters the state in Iroquois County, northeast of Hoopston, and continues westward into Ford County just east of Paxton. The valley proceeds westward across the county south of Paxton turning southwest into Champaign County about 8 miles northwest of Rantoul. The major tributaries to the Mahomet Valley include the Onarga Valley, with its major tributaries, the Kempton and Chatsworth Valleys (figures 2b and c).

Generalized graphic logs of the glacial deposits and bedrock formations of Ford County are given in figure 3.

#### GROUNDWATER

Groundwater in Ford County begins as precipitation which seeps downward into the ground through the soils. Figure 4a shows the generalized cycle of water movement from the atmosphere as precipitation to the surface and into the ground, and then away from the area either through the ground and in flowing streams or again into the atmosphere through transpiration of plants and evaporation.

#### Occurrence and Movement

Water enters and filters slowly down through the ground until it reaches a level where all available voids are completely water-filled. Water thus contained in this zone of saturation is groundwater, and its upper surface is the water table. Figure 5 illustrates the general configuration of the water table surface in Ford County. The water table normally lies some 5 to 10 feet below ground level in the lowlands along the streams (points of discharge) and from 15 to 25 feet below ground level in the upland areas. Seasonal fluctuations in the water table levels should be expected to range from 5 to 15 feet.

In glacial drift deposits, water fills the voids between the soil particles that make up the formations. In bedrock, water occurs primarily in two ways - it is contained in the spaces between partially cemented grains of sandstone or in the fractures, bedding planes, and solution cracks of limestone formations. A saturated formation of sand, gravel, sandstone, or limestone that is capable of yielding water to wells in usable quantities is called an aquifer.

GLACIAL DRIFT SECTION

	STAGE	FORMATION THICKNESS (FT) <i>(not to scale)</i>	SECTION	MATERIALS DRILLERS TERMS	WATER-YIELDING CHARACTERISTICS
UPPER GLACIAL DEPOSITS	WISCONSINAN	50 - 200		TILL, GRAVEL, SAND, SILT, LOESS	WATER-YIELDING FROM SAND AND GRAVEL DEPOSITS THROUGHOUT MOST OF THE COUNTY. WELL YIELDS FROM 5 - 400 gpm.
				LOESS SILT, LOESS, PEAT	
LOWER GLACIAL DEPOSITS	ILLINOIAN	0 - 250		TILL, GRAVEL, SAND	WATER-YIELDING FROM SAND AND GRAVEL LAYERS IN SOUTH-CENTRAL PORTION OF THE COUNTY. WELL YIELDS FROM 10 - 600 gpm.
	KANSAN	0 - 275		TILL, GRAVEL, SAND, SILT	
BEDROCK					

After Horberg (1953)

UPPER BEDROCK SECTION

SYSTEM	SERIES OR GROUP	FORMATION THICKNESS (FT) <i>(not to scale)</i>	GRAPHIC LOG	ROCK TYPE (DRILLERS TERMS)	WATER-YIELDING CHARACTERISTICS; DRILLING AND WELL CONSTRUCTION DETAILS	
	PLEISTOCENE	0-500		UNCONSOLIDATED GLACIAL DEPOSITS, ALLUVIUM AND WIND-BLOWN SILT (DRIFT, SURFACE, OVERBURDEN)	WATER-YIELDING CHARACTER VARIABLE. LARGE YIELDS FROM THICKER SAND AND GRAVEL DEPOSITS IN BEDROCK VALLEYS. WELLS USUALLY REQUIRE SCREENS AND CAREFUL DEVELOPMENT. CHIEF AQUIFER IN AREA.	
PENNSYLVANIAN	MCLEANSBORO	0-1000		MAINLY SHALE WITH THIN LIMESTONE, SANDSTONE AND COAL BEDS (COAL MEASURES)	WATER-YIELDING CHARACTER VARIABLE. LOCALLY SHALLOW SANDSTONE AND CREVICED LIMESTONE YIELD SMALL SUPPLIES. WATER QUALITY USUALLY BECOMES POORER WITH INCREASING DEPTH. MAY REQUIRE CASING.	
	CARBONDALE	0-150				
	TRADEWATER CASEYVILLE	0-600				
MISSISSIPPIAN	CHESTER	0-500		LIMESTONE, SANDSTONE AND SHALE	TOO DEEP TO BE CONSIDERED AS A SOURCE OF GROUNDWATER IN THIS AREA.	
	VALHEYER	STE. GENEVIEVE		0-120	LIMESTONE	MAY BE WATER-YIELDING IN MASON COUNTY WHERE THESE FORMATIONS ARE PRESENT AT A SHALLOW DEPTH. IN THE REST OF THE AREA, TOO DEEP TO BE CONSIDERED AS A SOURCE OF GROUNDWATER.
		ST. LOUIS- SALEM		0-270	LIMESTONE	
		WARSAW		0-130	SHALE	
	KEOKUK- BURLINGTON	0-300	CHERTY LIMESTONE			
	KINDERHOOK	0-200		SHALE	NOT WATER-YIELDING	
DEVO- NIAN		0-70		LIMESTONE	WATER-YIELDING FROM CREVICES WHERE ENCOUNTERED AT A SHALLOW DEPTH. IN MOST OF THE AREA, TOO DEEP TO BE CONSIDERED AS A SOURCE OF GROUNDWATER.	
SILU- RIAN	NIAGARAN	0-350		DOLOMITE AND LIMESTONE		
	ALEXANDRIAN	0-100				
ORDOVICIAN	CINCINNATIAN	MAQUOKETA 0-200		SHALE WITH LIMESTONE AND DOLOMITE BEDS	NOT WATER-YIELDING AT MOST PLACES; CASING REQUIRED.	
	MOHAWKIAN	GALENA-PLATTEVILLE		300-430	LIMESTONE AND DOLOMITE	NOT IMPORTANT AS AQUIFERS. CREVICED DOLOMITE PROBABLY YIELDS SOME WATER TO WELLS DRILLED INTO UNDERLYING SANDSTONE.
		CHAZY		GLENWOOD-ST. PETER	150-300	SANDSTONE, CLEAN, WHITE, THIN DOLOMITE AND SHALE AT TOP (ST. PETER)
	PRAIRIE DU CHIEN	SHAKOPEE		200-410	CHERTY DOLOMITE THIN BEDS OF SANDSTONE	NOT IMPORTANT AS AQUIFER. LINERS IN LOWER ST. PETER SANDSTONE ARE COMMONLY SEATED IN UPPER PART OF SHAKOPEE.
				NEW RICHMOND	0-175	SANDSTONE AND DOLOMITE
	ONEOTA	300-500	DOLOMITE WITH SOME SANDSTONE BEDS (LOWER MAGNESIAN)			

After Selkregg & Kempton (1958)

Figure 3. Generalized graphic logs of glacial deposits and bedrock formations

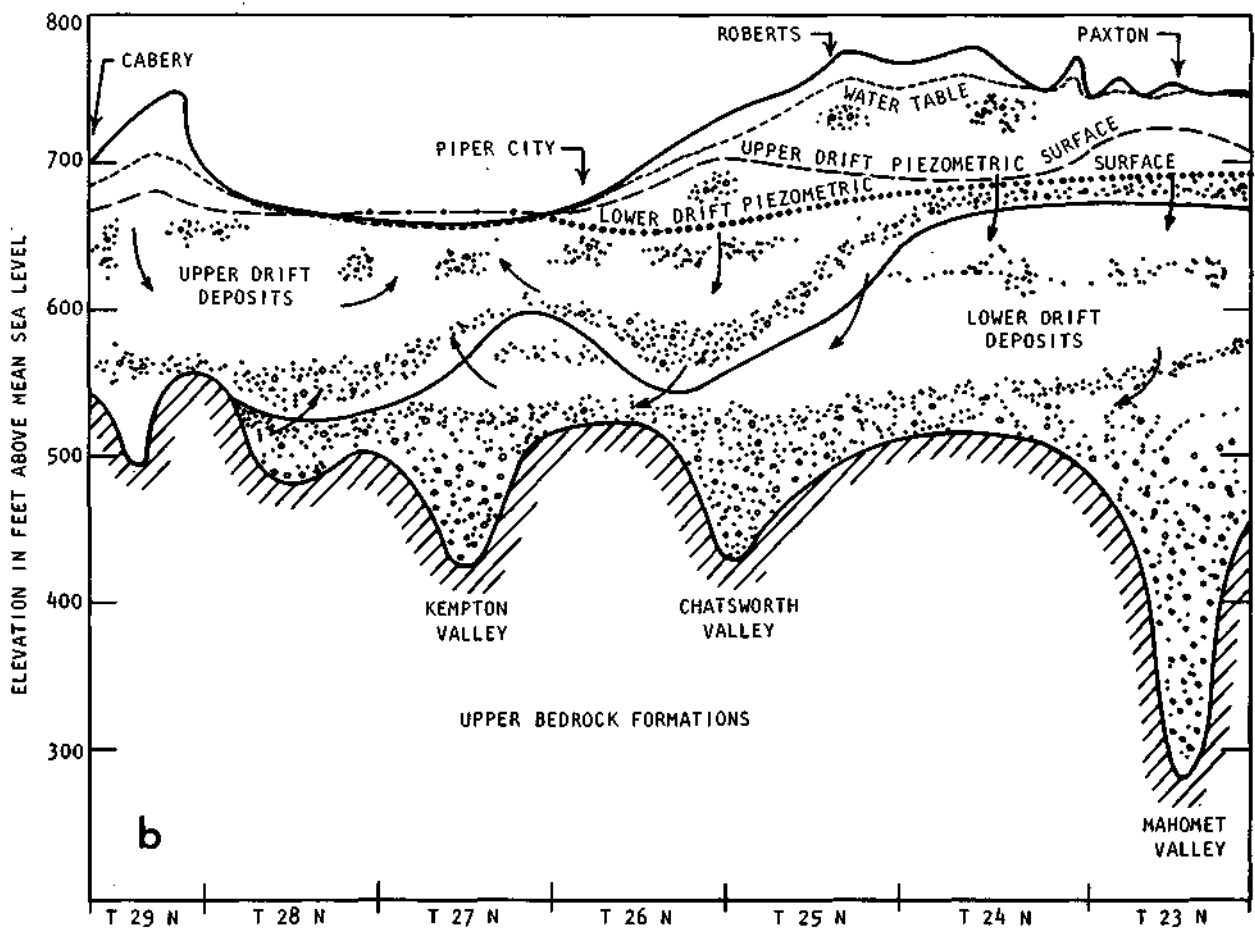
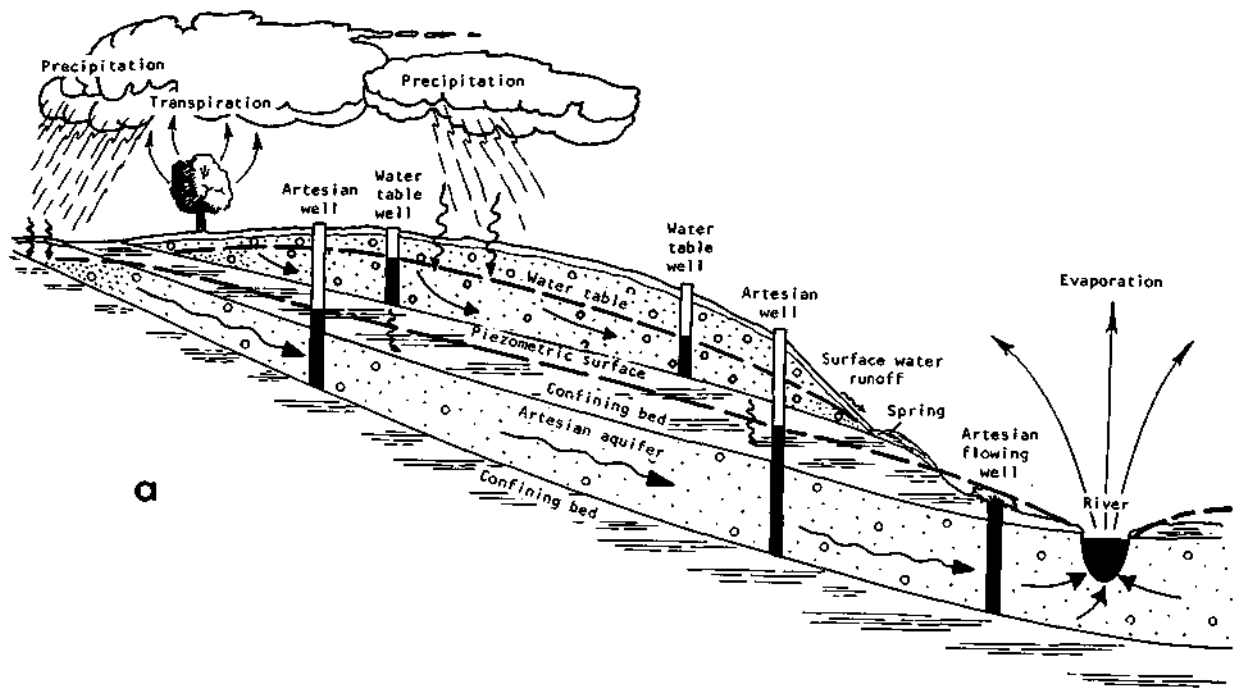


Figure 4. Cycle of water movement (a) and generalized movement of water in Ford County (b)

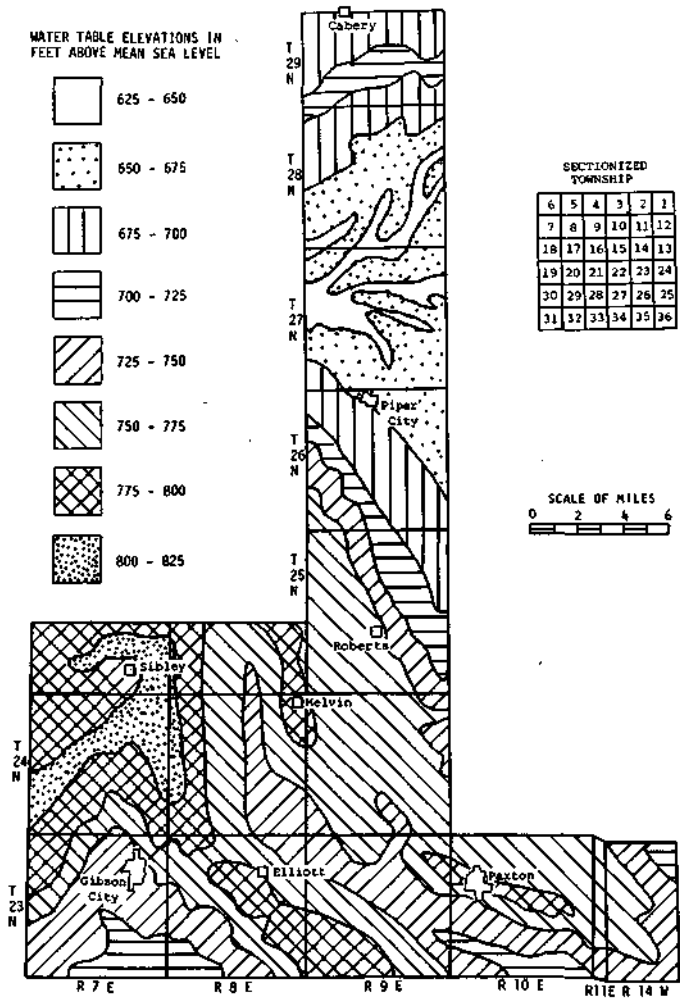


Figure 5. Water table contour map of Ford County

Under normal conditions, the upper glacial drift deposits are regularly recharged (refilled) by precipitation occurring in the immediate vicinity of the aquifer. Water continues to move freely downward under the influence of pressure and gravity to recharge the lower drift deposits and in some cases the underlying bedrock formations. However, layers of very dense (almost impermeable) materials separating water-bearing units may impede the downward movement of water. These layers, or confining beds, are usually clays or shales so compact that they cannot yield enough water to be classified as an aquifer. When such confining beds are present, water reaching the aquifer may come from a somewhat distant recharge area where the confining beds are missing or where the aquifer crops out at the land surface.

Water entering permeable formations in an outcrop or recharge area may become confined downslope beneath impermeable beds. Pressure is exerted on the groundwater in the confined aquifer by the weight of water at higher levels in the aquifer system. When a well penetrates such an aquifer downslope from the recharge area, the pressure forces the water to rise in the well above the

top of the aquifer. The water in this instance is confined (or artesian) water, the well is an artesian well, and the upper surface of the water in the well is the piezometric surface of the aquifer. When the piezometric surface of the aquifer is above land surface, wells tapping the aquifer are flowing artesian wells (see figure 4a).

Groundwater movement from recharge areas to discharge points is influenced by gravity and pressure differences. Major points of discharge include springs, lakes, streams, swamps, drainage tiles, and pumping wells. The rate of movement toward points of discharge may amount to a few hundred feet per year in unconsolidated materials and to only a few feet per year in sandstone formations. Water may be held in bedrock aquifers for many years.

The general direction of movement of groundwater in Ford County is illustrated in figure 4b. Precipitation falling in the upland areas south of Cabery and Piper City infiltrates into the upper drift deposits where a portion of it is diverted to discharge into local streams and drainage ditches (note the intersection of the water table surface and stream valleys just north of Paxton). The water not discharged locally continues to move downward to recharge the lower drift deposits. Along with the general downward migration of water in both

formations, there is movement downslope.(on the piezometric surfaces) toward the lowland drainage area just north of Piper City. Here the piezometric surfaces of both aquifers occur above land surface resulting in an upward movement of water from both systems to discharge into the streams and ditches draining the area.

### Chemical Character

The results of chemical analyses made by the State Water Survey usually are expressed in parts per million (ppm). A part per million is a unit weight of a constituent in a million unit weights of water; thus, a water sample containing 1 ppm of iron (Fe) would indicate 1 pound of iron in a million pounds of water. In order to express chemical dissociations and show water analyses graphically, chemically equivalent weights or equivalents per million are used. Equivalents per million are calculated by dividing the parts per million by the combining weight of the respective cation or anion. Analyses made by private chemical laboratories sometimes are reported in terms of grains per gallon (gpg). In the grain weight system of measure, one grain per gallon is considered a 1/7000 of a pound of a mineral dissolved in a gallon of water. One grain per gallon is equal to 17.1 parts per million.

The sources and significance of the major dissolved elements and substances in the groundwater of Ford County are shown in table 1. The U. S. Public Health Service drinking water standards for these major constituents also are included in the table. These standards have been accepted by the American Water Works Association as desirable limits for public water supplies, and should serve as a guide to owners of farm and home water supplies in evaluating the quality of their water.

As generally may be inferred from the information in table 1, the dissolved minerals in groundwater are derived chiefly from the earth materials through which the water flows. The soils and glacial materials above bedrock are particularly rich in calcium, magnesium, iron, and other minerals which are readily absorbed by the groundwater as it passes over and through these deposits. Calcium and magnesium are responsible for hardness of water, and iron causes reddish-brown staining. The natural chemical quality of groundwater is sometimes altered by highly mineralized surface water that seeps down along the casings in poorly constructed wells. This type of seepage also can result in bacterial pollution of the well and contamination of the aquifer.

Groundwater from glacial deposits throughout the county varies from moderately hard to extremely hard (150 to 2000 ppm), but normally hardness can be successfully removed by home water-softening units that are now readily obtainable. The iron content of water from the drift deposits ranges between 0.1 and 15.0 ppm, and is most often well above the recommended limit of 0.3 ppm. Iron can be removed by units similar to home water softeners; however, for domestic users, tolerance rather than removal is the usual practice. The chemical quality of water from the various glacial formations is discussed in more detail later in this report.

The chemical quality of water from the bedrock aquifers in Ford County varies considerably depending on the formation tapped and depth of penetration into the aquifer. Generally speaking, water from the bedrock becomes more highly mineralized with depth. Chloride, sulfate, and sodium are normally present in larger concentrations than in drift water. A more detailed discussion of the

Table 1. Major Dissolved Elements and Substances in Groundwater in Ford County

Constituent and recommended upper limit <sup>1</sup>	Source	Remarks
Iron (Fe) 0.3 ppm	Dissolved from common iron-bearing minerals present in practically all rocks, clays, and soils; may also be derived from iron pipes, pumps, and other equipment.	On exposure to air, iron oxidizes to a reddish-brown sediment. More than about 0.3 ppm stains laundry and porcelainware reddish brown; objectionable for food and beverage processing. Small concentrations may be removed by water softeners, larger concentrations by chlorination or aeration and filtration.
Manganese (Mn) 0.05 ppm	From soils and sediments. Less abundant in rocks than is iron.	Resembles iron in chemical behavior and has same objectionable features, except stains are brown to black. The same types of treatment used for iron are generally effective.
Nitrate (NO <sub>3</sub> ) 45 ppm	Results from decayed organic matter such as that from barnyards, feedlots, manure piles, septic tank fields, as well as from silage juices and animal tissue. Usually occurs in waters from shallow wells of less than 50-foot depth, often as the result of poor well construction permitting drainage into the well at or near the surface.	Values higher than a few ppm may suggest pollution. More than about 45 ppm nitrate may cause methemoglobinemia (blue babies) when such water is used in preparation of infant feeding formulas. Nitrate poisoning of livestock has also been reported. Removal by demineralization is possible but usually prohibitive in cost.
Chloride (Cl) 250 ppm	Dissolved from rocks and found in large amounts in ancient brines and sea water. Industrial and domestic waste also may contribute appreciable quantities to shallow aquifers.	In concentrations over about 250 ppm chloride gives a salty taste to water and increases its corrosiveness. Concentrations over 1000 ppm are very objectionable for domestic use; livestock tolerance is considerably higher.
Sulfate (SO <sub>4</sub> ) 250 ppm	Dissolved from rocks and soils containing gypsum, iron sulfides, and other sulfur compounds. Present in waters from coal mine drainage and some industrial wastes.	Sulfate in water containing calcium forms a hard scale. In concentrations over about 750 ppm sulfate in combination with sodium or magnesium has a laxative effect, most noted by infrequent users of the water.
Alkalinity [bicarbonate (HCO <sub>3</sub> ) and carbonate (CO <sub>3</sub> )]	Results from action of carbon dioxide or acid in water on carbonate rocks such as limestone and dolomite.	Alkalinity is present almost entirely in the form of bicarbonates. In the presence of calcium, carbonates formed may produce a carbonate scale; they decompose on heating with release of carbon dioxide gas and attendant formation of calcium carbonate scale.
Hardness (as CaCO <sub>3</sub> )	Caused by calcium and magnesium which occur in almost all rocks but especially in limestone, dolomite, and gypsum.	Before a lather will form, hard water precipitates soap, forming a sludge which causes deposits on bathtubs and is responsible for gray laundry and dingy glassware. Hard water also forms scale in boilers, hot water heaters, and pipes. Hardness normally can be removed by standard home water softening units.
Total dissolved minerals 500 ppm	Includes all mineral constituents dissolved from rocks and soil.	Mineralization of more than 500 ppm is normally detectable by taste; over 1000 ppm is undesirable for most domestic purposes; livestock may tolerate concentrations up to 7000 ppm. <sup>2</sup>

<sup>1</sup>U. S. Public Health Service. 1962. *Drinking water standards*. Publication No. 956.

<sup>2</sup>*Salinity and livestock water quality*. 1959. South Dakota State College Agricultural Experiment Station. Bulletin 481.



chemical character of bedrock water is included in the bedrock formations section of this report.

Water from wells throughout Ford County contains varying quantities of carbon dioxide. A relatively small area in and around Sibley also has water known to contain methane gas. These gases are colorless, odorless, and tasteless. Methane gas is lighter than air whereas carbon dioxide is heavier. When methane gas is mixed with air in concentrations of 5 to 15 percent, it is highly explosive if ignited. Most dangerous points of concentration are in the well house, within the air cushion of pressure tanks, and in hot water heaters. All such points should be vented to the outside air if methane gas is detected in a water supply. All new wells constructed should be checked for methane gas by the driller before the installation is placed in service.

Further, no one should ever enter a large-diameter dug well without previously checking for the presence of methane gas or carbon dioxide, both of which can cause asphyxiation. These gases can be readily removed from water by standard aeration procedures.

Nitrates, or simple nitrogen compounds that occur in water as mineral constituents, are considered harmful to people, particularly children, if concentrated in drinking water supplies in excess of 45 ppm. Nitrate poisoning of livestock has also been reported. Excessive concentrations of nitrate have been detected in only a few groundwater samples from Ford County. Among wells sampled during the 1967 well inventory, three bored wells (61, 84, and 122 feet deep) and three dug wells (16, 20, and 35 feet deep) contained more than 45 ppm nitrate. Approximately 2 percent of the sampled wells of less than 50 feet deep had nitrate concentrations greater than *h*5 ppm, and only 3 percent had concentrations greater than 20 ppm.

Primary sources of the nitrate contamination in these wells were nearby septic tanks, old privies, or drainage from feedlots and pastures. Leachates (seepage) from these sources percolating through the soil or flowing overland to the wells have been determined to be the source of nitrate pollution in practically all such cases. Nitrogen fertilizers have not yet been demonstrated to be of importance in Illinois groundwater pollution. However, these may become a significant future source as larger quantities of nitrogen-rich fertilizers are applied to the soils of the state.

The treatment of water supplies containing nitrate poses a difficult problem. Boiling the water does not help, but rather results in concentrating the nitrate. Nitrates and similar mineral constituents such as chlorides and sulfates can be reduced or removed by demineralization, a process not economically desirable for private water supplies. It is usually easier to abandon the source of high nitrate water and develop a supply either at a location not susceptible to the nitrates or in another aquifer horizon.

Mineral analyses of groundwater from throughout the county are included in appendix B of this report.

#### Temperature

The temperature of groundwater varies with the location and depth of the aquifer, the origin and time of occurrence of recharge, and the proximity of the

aquifer to bodies of surface water. In Ford County the primary source of recharge is precipitation which enters the groundwater reservoirs mostly in the early spring and late fall when the precipitation is fairly cool. After infiltrating into the ground, little temperature change occurs because of the insulating effect of the surrounding earth materials. The lower groundwater temperatures generally are associated with the shallower aquifers, and temperatures slowly increase at a rate of about 3/4 to 1 degree Fahrenheit for each 100 feet in depth. Therefore, water obtained from a 1000-foot deep bedrock well should be expected to be approximately 10 degrees warmer than that from a shallow glacial drift well. Groundwater temperatures range from about 54 degrees Fahrenheit in the upper glacial deposits to 65 degrees in the deeper bedrock formations tapped in Ford County (see appendix B).

#### AQUIFERS

Aquifer selection for farm and domestic well construction in the past was often influenced by the quantity of water required, the type of drilling equipment available, and perhaps the amount of money the farmer or homeowner was willing to pay. In most cases, the shallowest water-bearing sand and gravel deposit encountered was capable of satisfying the relatively small water demands, could be easily developed, and provided the cheapest solution to the water supply problem. However, with the increased use of water on the farm and in the home, higher yielding wells than those previously constructed are often required.

Throughout most of Ford County there are two glacial aquifers, each containing one or more layers or zones of water-bearing sand and gravel. In many areas, the deeper deposits are more productive than the shallower sands and gravels currently being tapped. For these reasons, drilling for a farm and domestic well should continue until a satisfactory supply is obtained or until the underlying bedrock is encountered.

For larger capacity municipal or industrial wells, it is advisable to construct a test hole penetrating the entire drift section to determine the presence and relative potentials of the different glacial aquifers prior to completion of the final production well.

For either type of development, the services of a competent well driller experienced in constructing sand and gravel wells should be obtained to maximize the use of available resources. Drilling into the underlying bedrock aquifers is recommended only if a satisfactory supply cannot be developed from the glacial materials.

For the purposes of this study, the various water-bearing units are grouped into three general aquifer systems. The Wisconsin age glacial deposits outlined by Horberg (1953) are hereafter designated as the upper glacial deposits or upper aquifer system.

The underlying Illinoian and Kansan age glacial deposits defined by Horberg (1953) appear to be hydraulically interconnected throughout most of the county and generally respond to pumpage as a unit. For this reason these deposits are considered as one geohydrologic unit, termed the lower glacial deposits or lower glacial aquifer.

The third aquifer system, the bedrock formations, is composed of all fresh-water bedrock units underlying the drift deposits throughout the county. Detailed discussions on the occurrence, water-yielding characteristics, and quality of water of each aquifer system are presented in the following sections.

### Upper Glacial Deposits

Water-bearing sand and gravel deposits contained within the Wisconsin drift section serve as a source of water for approximately 70 percent of the individual farm and domestic water supplies in the county. These deposits occur as scattered pockets in the upper portions of the aquifer, as fill materials in the Kempton Bedrock Valley, and as extensive sheet or striplike deposits present at or near the base of the upper glacial materials throughout most of the county.

**Development and Availability.** Wells tapping the upper glacial aquifer are of four general types: dug, bored, driven, and drilled. Records of approximately 150 large-diameter dug and bored wells were collected from Water Survey files and from the direct inventory conducted during the summer of 1967. These wells range in depth from 20 to 60 feet below ground level and are from 12 to 48 inches in diameter. The yields of these wells are never large and are often barely adequate for domestic use. Most of these installations were constructed many years ago when water demands were small and mechanized drilling equipment was not always available. Today's larger water requirements usually cannot be obtained from such wells, and the availability of aquifers suitable for drilled well development in Ford County make dug wells undesirable for most uses.

Records are available for 27 driven wells tapping the upper glacial aquifer to depths of about 20 feet. These wells are from 1-1/2 to 2 inches in diameter and usually yield less than 5 gpm, but generally are adequate for domestic use. This type of well construction in Ford County is limited to a small area south of Gibson City where relatively shallow sands and gravels are present.

The Water Survey files include records of 945 drilled wells finished in the upper aquifer system. These installations range in depth from 50 to 75 feet in the lowland areas of the major drainage courses of the county to about 200 feet on the higher grounds just south of Cabery, northwest of Roberts, north of Gibson City, and in the general vicinity of Paxton.

Figure 6 illustrates the probable maximum depth of a well finished in this aquifer anywhere in Ford County. In many locations several different water-bearing zones are present and conceivably could be penetrated before the maximum depths indicated would be reached. To insure that a suitable aquifer is not overlooked, homeowners are again urged to obtain the services of a competent well driller experienced in constructing screened wells in sand and gravel.

Farm and domestic drilled wells in this aquifer system range from 2 to 6 inches in diameter. However, well diameters between 4 and 6 inches are more desirable than the smaller wells because of the greater selection in types of pumps which can be used and the increased accessibility to the pump for inspection and repairs. The diameters of municipal and industrial wells in the upper deposits range from 6 to 24 inches, usually determined by the desired pump size requirements.

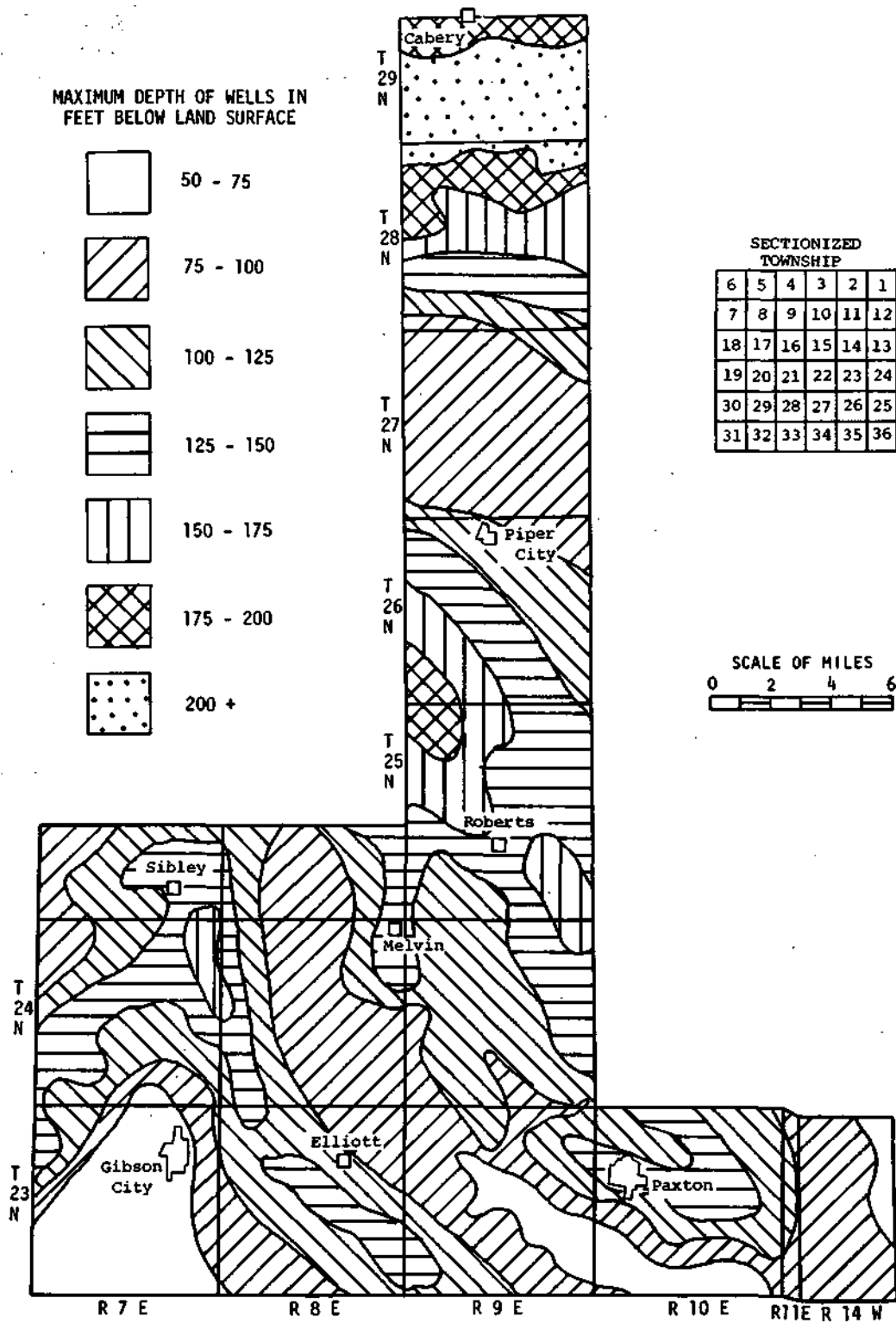


Figure 6. Probable maximum depth of wells in the upper glacial deposits

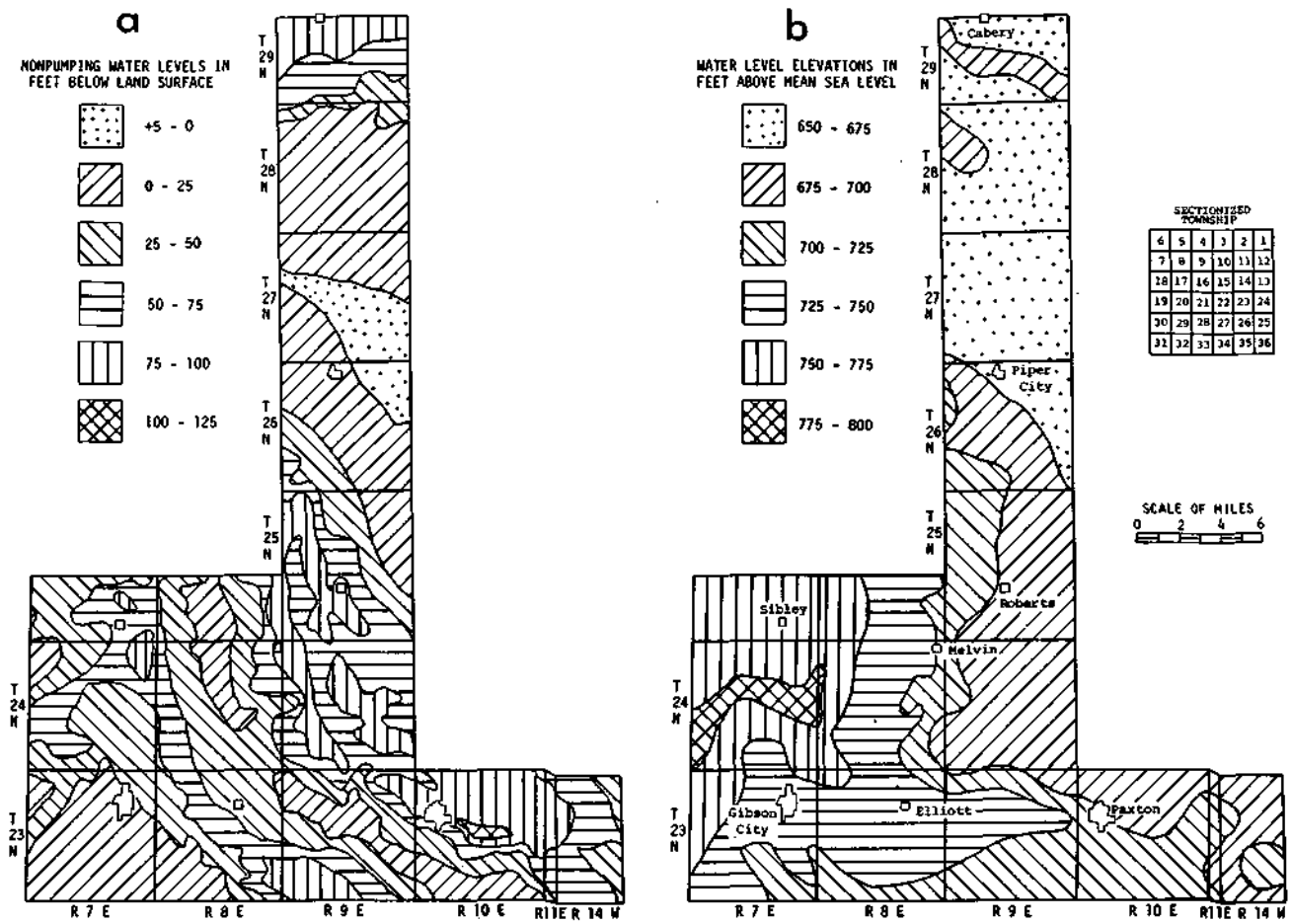


Figure 7. Nonpumping water levels (a) and water-level elevations (b) in drilled wells tapping the upper glacial deposits

The nonpumping water levels of the drilled wells finished in the upper aquifer are shown in figure 7a. The depths to water are to a large extent influenced by the land surface topography. Seasonal fluctuations as great as 5 or 10 feet may be experienced in the shallower wells and normally will be less in the deeper wells. Figure 7b illustrates the mean water-level elevations that occur in the drilled wells in the upper glacial deposits. The general direction of movement of water in this aquifer system can also be determined from this illustration since water always flows downward (from high to lower elevations) and perpendicular to the individual contours.

The yields of wells tapping the upper glacial deposits range from 5 gpm for farm and domestic wells (generally fixed by pump capacities) to over 400 gpm from the municipal wells finished in the outwash deposits at Gibson City. Municipal and industrial wells finished in this aquifer at Elliott, Gibson City, Kempton, Piper City, and Sibley produce approximately 60 percent of the municipal pumpage in the county. Installed pump capacities in these wells are 75 gpm, 250 to 400 gpm, 100 gpm, 140 to 170 gpm, and 58 gpm, respectively.

Total groundwater pumpage from the upper glacial aquifer is an estimated 1.4 mgd. Much larger quantities of water could probably be withdrawn without overdevelopment. Pumping tests and aquifer evaluations would be required to accurately determine the groundwater potential in localized areas.

*Chemical Character.* The chemical quality of water from the upper glacial deposits varies considerably within Ford County. Results of analyses of water from approximately 110 drilled wells finished in this aquifer system are included in appendix B. Typical analyses are shown graphically in figure 8a.

Available chemical and hydrologic data suggest that the variation in the chemical constituents in this aquifer system generally can be explained by the period of time the water has been underground and in contact with soil particles from which mineral constituents are dissolved. Relatively low mineralized water, indicating a short period of residence, is usually found in areas of recharge. More highly mineralized waters, suggesting longer periods of contact, are associated with regions of groundwater discharge. The validity of this approach is indicated by the similarity in general configuration of the variation in total dissolved minerals in figure 8b and the movement of water through these deposits shown in figure 7b.

Localized disturbances in the overall general pattern of chemical quality variability are evident in an area east of Roberts and in the extreme northern portion of the county (figures 8b and c). The more highly mineralized waters encountered in the area east of Roberts are probably the result of two separate phenomena. The wells in this area are somewhat deeper (*see figure 6*) and water recharged to their water-yielding deposits must percolate downward through a greater thickness of overlying materials. Therefore, water finally reaching these deposits would have experienced a longer period of residence during which minerals from the surrounding materials could be dissolved than water derived from shallower depths. Also, available logs of wells in this area suggest the presence of several scattered peat zones which normally would increase the acidity of water coming into contact with these beds, and in turn would result in an increase in total dissolved minerals.

The higher mineral content in the northern portion of the county also can be partially attributed to the increased depth of the wells in that area. In addition, the upper glacial deposits in this area lie directly on the Pennsylvanian age bedrock formations which characteristically contain highly mineralized water. Piezometric surface data show a gradient from the bedrock formations toward the upper aquifer indicating that the more highly mineralized water from the bedrock is being discharged upward into these deposits.

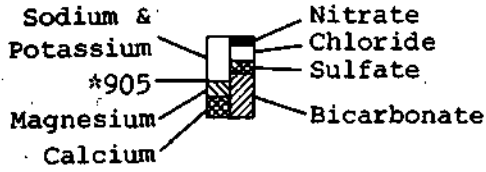
Total mineral concentrations of water from the upper glacial deposits are illustrated in figure 8b. Water from the upper aquifer system is below the recommended 500 ppm upper limit set by the U. S. Public Health Service in all areas except for the two locations just discussed.

The hardness of water obtained from these deposits throughout the county ranges from moderately hard (200 ppm) to extremely hard (over 400 ppm). Figure 8c suggests the probable hardness content which may be expected from this aquifer for any given location. In all areas, the general quality of water from these deposits could be improved by the use of standard home water-softening units.

Iron concentrations above the recommended 0.3 ppm upper limit set by the U. S. Public Health Service were detected in over 90 percent of the sampled wells. Water containing iron in such quantities usually causes staining of laundry and porcelain fixtures unless some type of iron removal equipment is used. Manufacturers of home water softeners advertise that these units will effectively remove as much as 5.0 ppm "dissolved" iron. It should be noted,

EXPLANATION

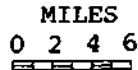
\* NUMBER AT TOP OF MAGNESIUM IS HARDNESS (as CaCO<sub>3</sub>) IN PARTS PER MILLION (ppm). NITRATE IS SHOWN WHEN MORE THAN 10 ppm.



WELL LOCATION

SECTIONIZED TOWNSHIP

6	5	4	3	2	1
7	8	9	10	11	12
13	14	15	16	17	18
19	20	21	22	23	24
25	26	27	28	29	30
31	32	33	34	35	36



EQUIVALENTS PER MILLION

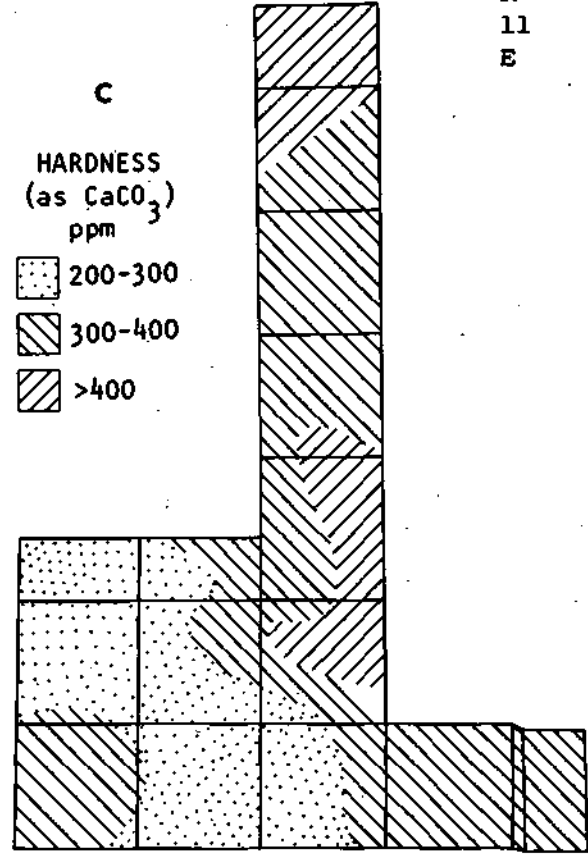
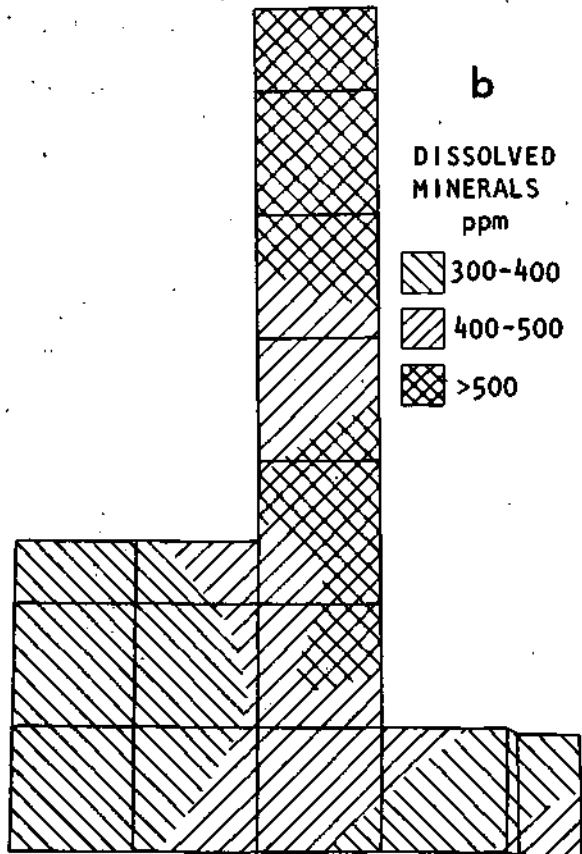
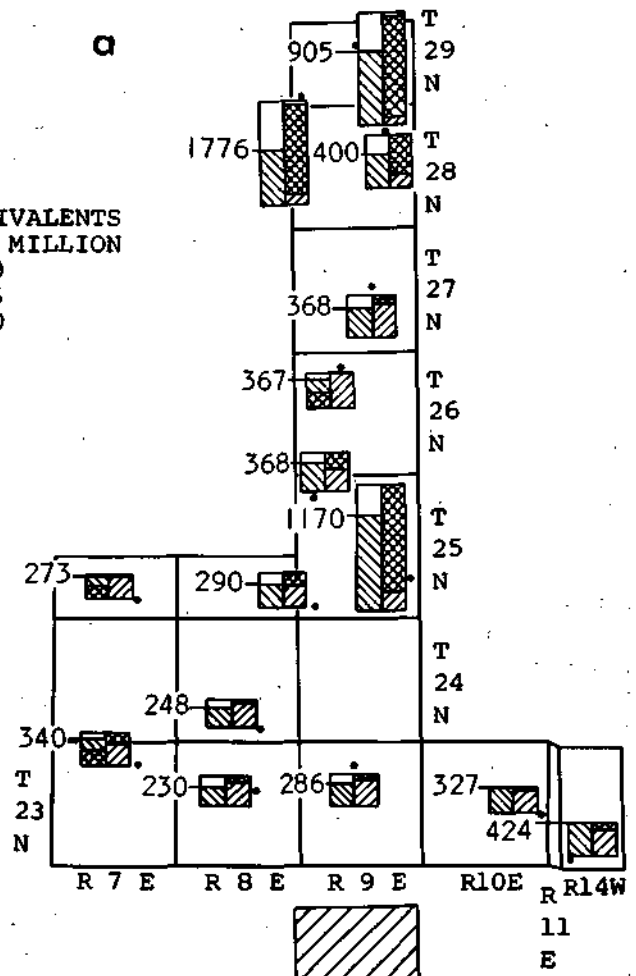
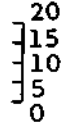


Figure 8. Selected chemical analysis (a), probable total dissolved minerals (b), and hardness (c) of water from the upper glacial deposits

however, that the iron found in the water in Ford County is normally in an insoluble state and can create serious plugging problems in water-softening units. The use of an oxidizing iron removal unit in conjunction with a water softener is usually required to provide continuous and effective treatment over a long period of time.

### Lower Glacial Deposits

Sand and gravel deposits of Illinoian and Kansan age serve as a source of water for approximately 20 percent of the farm and domestic water supplies in the county. Continuous layers of water-bearing sands and gravels occur near the base of the Illinoian deposits throughout all but the northern one-fourth of the county (see figure 9 for northern boundary). Thicker more permeable sections of sand and gravel (Kansan age) also occur in the fill materials of the ancient Chatsworth and Mahomet Bedrock Valleys.

*Development and Availability.* Private farm and domestic drilled wells tapping the lower glacial deposits range in depth from about 125 feet in the lowland area just north of Piper City to over 250 feet in the upland areas of the southern portion of the county. Figure 9 illustrates the probable maximum depth that wells of this type may be expected to reach in this aquifer system. Larger capacity wells tapping the deeper deposits in the bedrock valleys may range in depth from 250 to 340 feet below land surface.

The diameters of farm and domestic wells in this aquifer range from 2 to 6 inches, similar to the wells in the upper aquifer. Because these wells are deeper than those in the upper deposits, the larger selection in pump types and the accessibility for pump repair and maintenance afforded by the 4- and 6-inch wells become even more important. Municipal and industrial wells in the lower deposits are from 8 to 16 inches in diameter depending on the pump size requirements.

The nonpumping water levels of wells finished in the lower glacial aquifer are shown in figure 10a. Seasonal water-level fluctuations in these deposits are insignificant since water stored in the overlying materials normally is available to the lower aquifer system during prolonged periods of drought. Figure 10b illustrates the mean water-level elevations that occur in wells finished in the lower glacial deposits. The general direction of movement of water in this aquifer system also can be determined from this illustration since water always flows downward and perpendicular to the individual contours.

The yields of wells tapping the lower glacial aquifer range from 10 gpm for the farm and domestic wells (generally limited by the installed pump capacity) to about 1000 gpm for the larger capacity municipal and industrial wells finished in the permeable fill materials of the Mahomet Bedrock Valley. Municipal wells finished in the lower drift deposits at Melvin, Paxton, and Roberts produce approximately 38 percent of the current municipal pumpage in the county. Installed pump capacities in these municipal wells are 60 to 200 gpm, 100 to 1000 gpm, and 105 to 130 gpm, respectively. Maximum well yields obtained from the lower aquifer throughout the county generally are limited to about 200 gpm except for those finished in the productive Mahomet Valley deposits.

Total groundwater pumpage from the lower glacial aquifer is estimated to be 0.6 mgd. Much larger quantities of water can be withdrawn without overdevel-



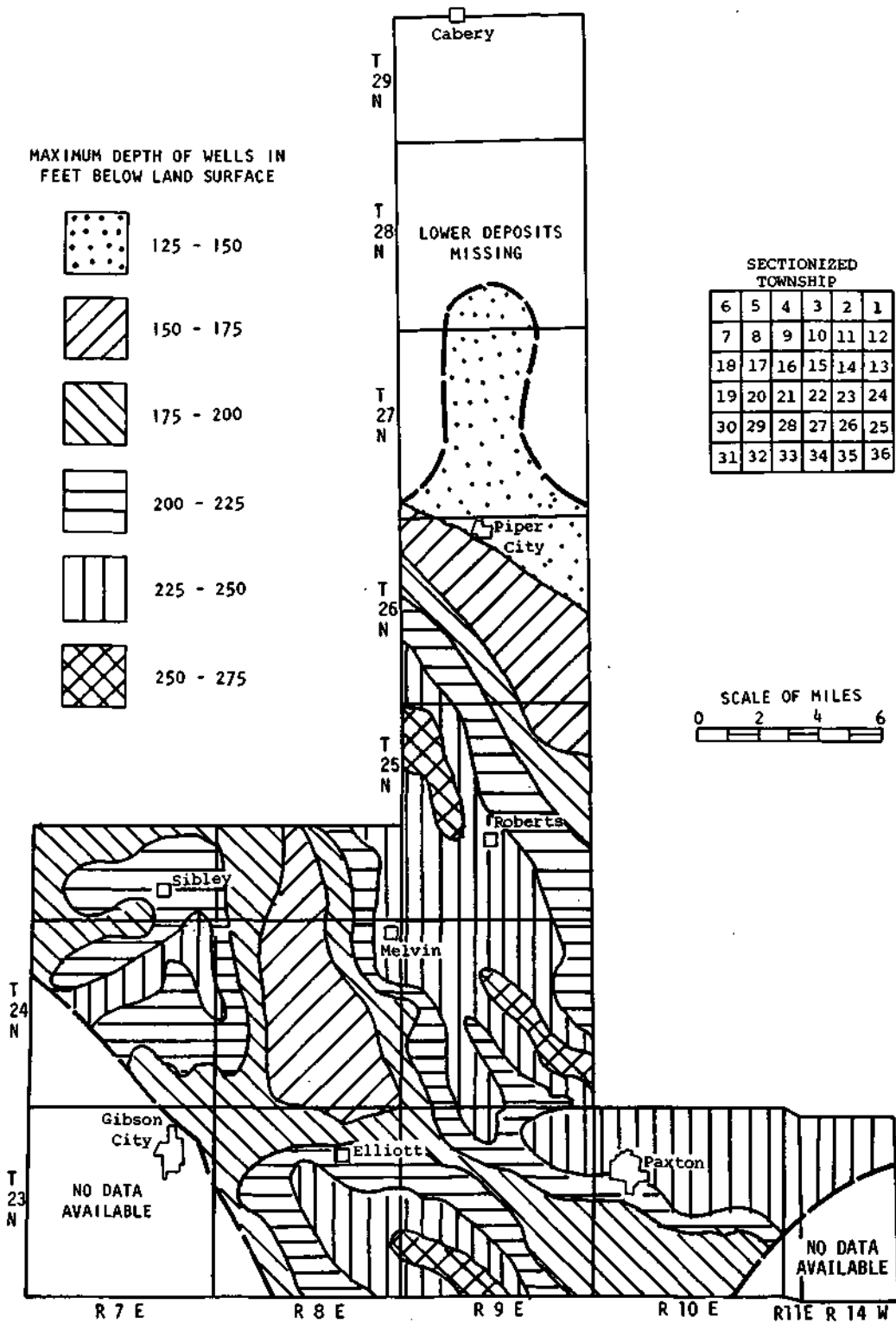


Figure 9. Probable maximum depth of wells in the lower glacial deposits

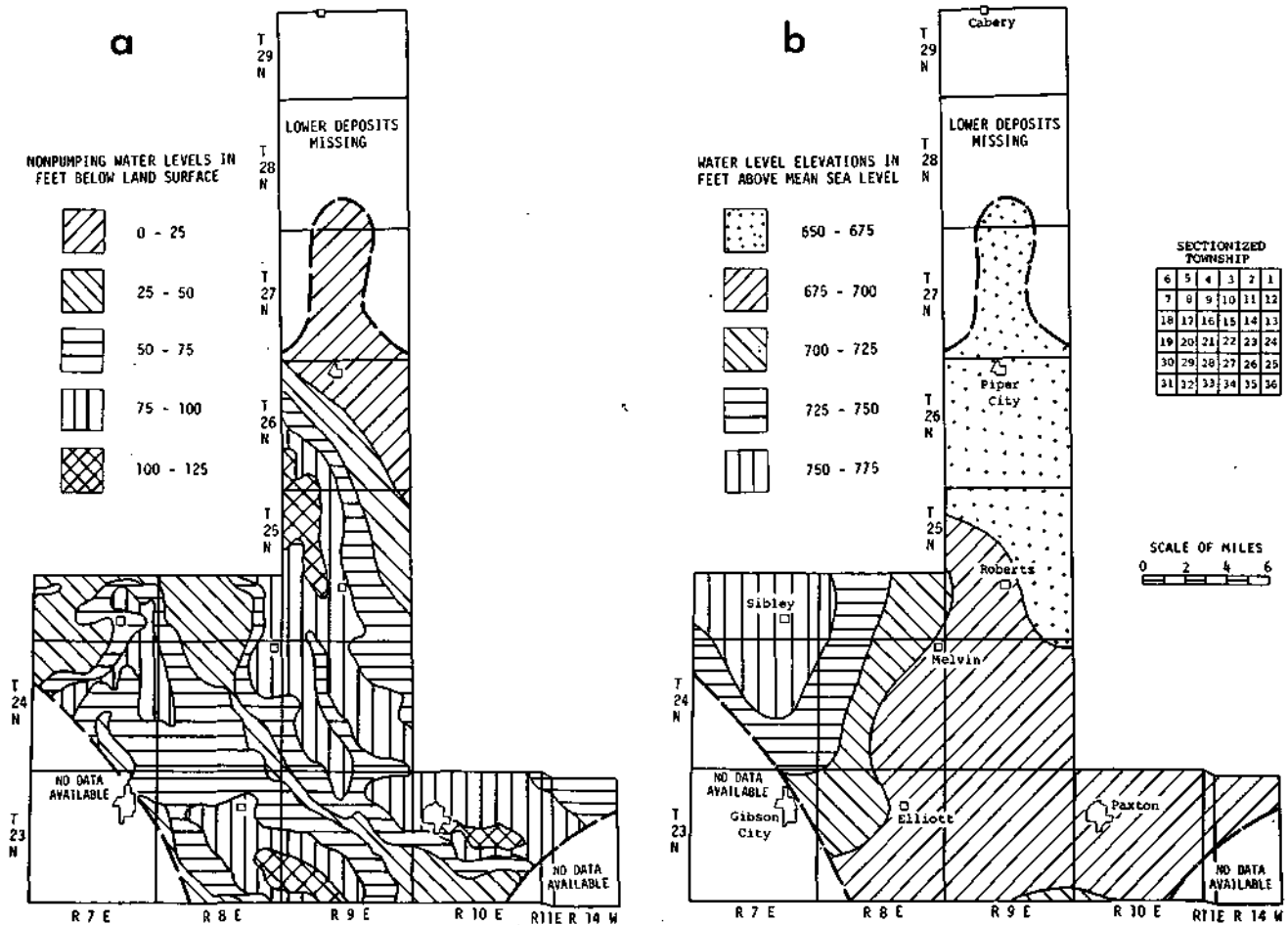


Figure 10. Nonpumping water levels (a) and water-level elevations (b) in wells tapping the lower glacial deposits

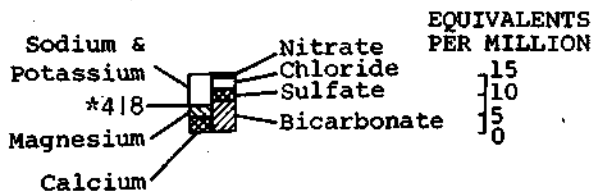
opment particularly in the areas overlying the Mahomet Bedrock Valley. The limited development of this aquifer system to date is largely due to the availability of water from the upper glacial deposits. Test drilling, pumping tests, and aquifer evaluations would be required to accurately determine the ground-water potential in localized areas.

*Chemical Character.* The chemical quality of water from the lower glacial deposits is less variable than that from the upper aquifer but generally is higher in total mineral content. Results of analyses of water from approximately 40 wells finished in this aquifer system are included in appendix B. Typical analyses are shown graphically in figure 11a.

Available chemical and hydrologic data generally imply that the length of time the water has been in the ground, and in contact with soil particles from which mineral constituents are dissolved, may account for most of the variation in the chemical constituents of water in this aquifer. Since water in the lower glacial deposits must first pass through the overlying upper aquifer, it follows that it should be slightly more mineralized than water from the upper glacial deposits. Relatively low mineralized water indicating the shortest period of residence is noted in the southern portion of the county (figure 11b) where recharge to the lower aquifer system occurs rapidly and readily because of the

**EXPLANATION**

\* NUMBER AT TOP OF MAGNESIUM IS HARDNESS (as CaCO<sub>3</sub>) IN PARTS PER MILLION (ppm). NITRATE IS SHOWN WHEN MORE THAN 10 ppm.



• WELL LOCATION

SECTIONIZED TOWNSHIP

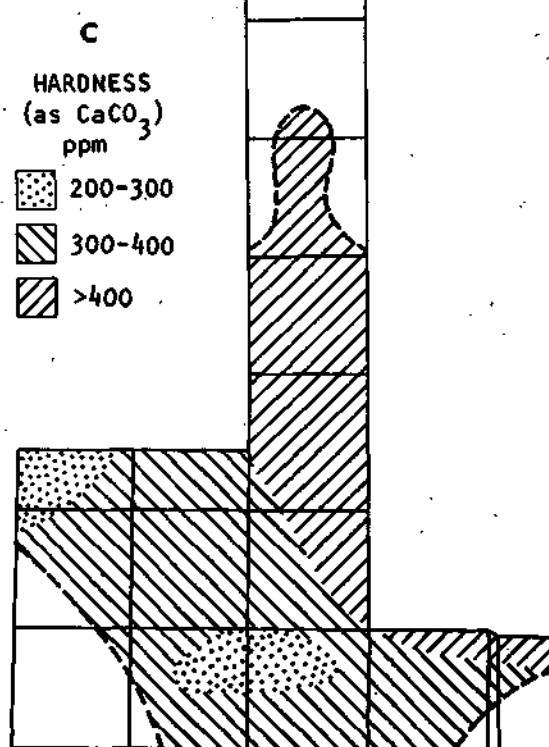
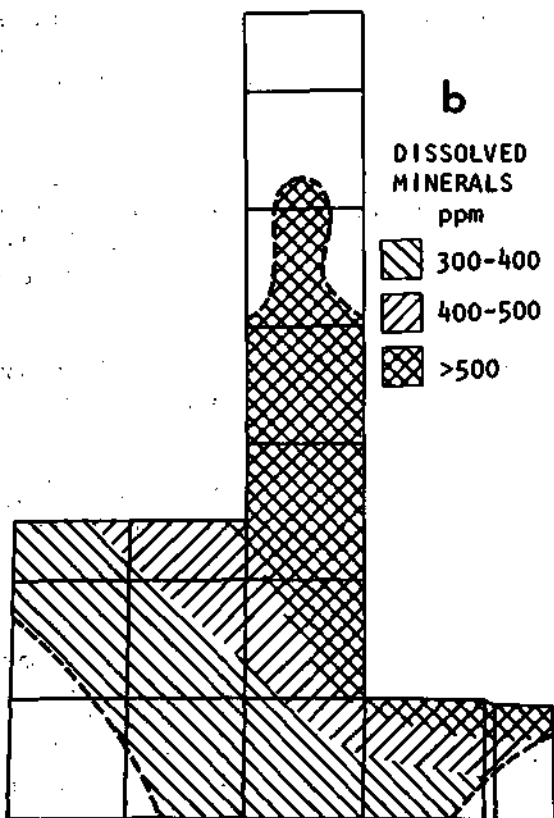
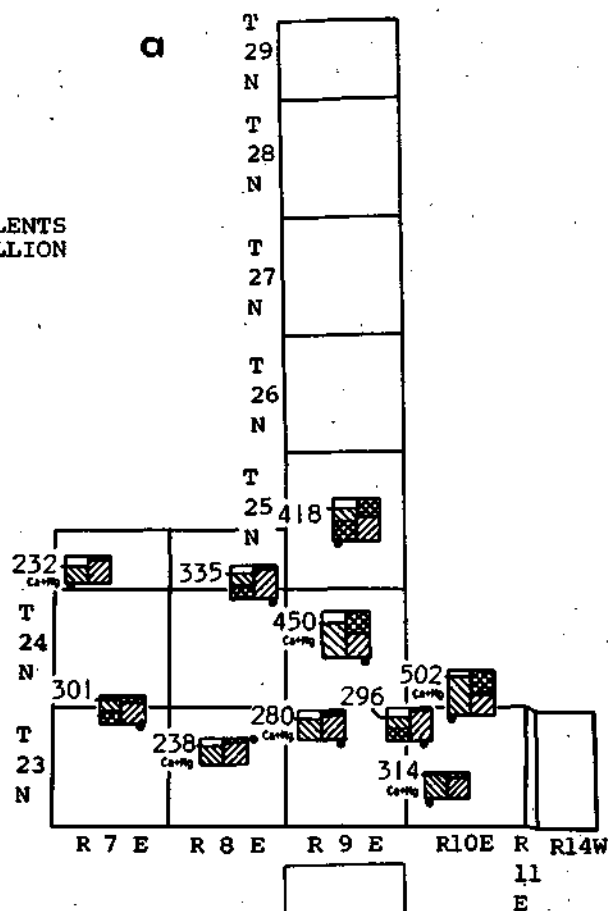
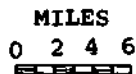
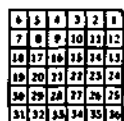


Figure 11. Selected chemical analysis (a), probable total dissolved minerals (b), and hardness (a) of water from the lower glacial deposits

absence of an overlying confining clay layer. Then as the water slowly moves northward (see figures 4b and 10b ) to discharge through the upper aquifer system, it gradually increases in total mineral content.

Total mineral concentrations of water from the lower glacial deposits are illustrated in figure 11b. Water from this aquifer system is below the recommended 500 ppm upper limit set by the U. S. Public- Health Service in the southwestern portion of the county.

The hardness of water obtained from the lower deposits ranges from moderately hard (200 ppm) to extremely hard (400 ppm). Figure 11c suggests the probable hardness content which may be expected from this aquifer for any given location. In all areas, standard home water-softening units could improve the general quality of water from these deposits.

Iron concentrations above the recommended 0.3 ppm upper limit set by the U. S. Public Health Service were detected in over 95 percent of the sampled wells. The use of an iron removal unit in conjunction with a water softener is usually the most efficient and long-lasting method for treating water of this quality.

### Bedrock Formations

The upper bedrock formations in Ford County consist of Pennsylvanian, Mississippian, Devonian, and Silurian age rocks (see figures 2a and d for surface configuration and areal distribution ). Water-bearing layers of sandstone, limestone, and dolomite contained within these geologic systems serve as a source of water for approximately 9 percent of the farm and domestic water supplies in the county.

*Development and Availability.* Private farm and domestic wells finished in the Pennsylvanian rocks range in depth, from 150 to 250 feet in the northern two townships of the county and from 350 to 450 feet in the area northwest of Sibley. Throughout the remaining portions, of the study area, the Pennsylvanian rocks are either not present or have not as yet been tapped by wells.

Farm and domestic wells finished in the Pennsylvanian rocks range in diameter from 4 to 8 inches. Two municipal wells owned by the village of Cabery tapping this aquifer are 6 and 8 inches in diameter.

The nonpumping water levels of wells tapping the Pennsylvanian formations range from about 25 to over 50 feet below land surface. Seasonal water-level fluctuations are negligible due to the availability of water from the thick overlying glacial materials.

The yields of wells finished in this formation range from 2 or 3 gpm from the thin layers of sandstone and limestone normally encountered in the upper portions of these rocks to over 100 gpm from thicker, more permeable water-bearing units occasionally found at greater depths. Municipal Well No. 3 (drilled 357 feet deep) owned by the village of Cabery is reportedly capable of producing 125 gpm for 4 hours with a drawdown of 34 feet from a nonpumping water level of 48 feet below ground level (specific capacity =3.4 gpm/ft of drawdown). However, individual well yields from the Pennsylvanian rocks throughout most of Ford

County should not be expected to exceed 3 gpm, and yields in excess of 25 gpm as at Cabery should be considered anomalies.

Other more promising bedrock aquifers in the county are as yet relatively unexplored. The Silurian and Devonian age formations have been tapped by only one or two wells in the vicinity of Gibson City. However, yield data from these wells and previous studies by Csallany and Walton (1963) indicate that individual well yields ranging from about 50 to 250 gpm may be obtainable from these formations.

The deeper lying St. Peter sandstone of Ordovician age may also be a promising aquifer in the northernmost part of the county. This deeply buried formation (1000 to 1200 feet below land surface) is the deepest known fresh-water aquifer in this part of Illinois. It has been tapped by the nearby towns of Cullorn, Chatsworth, and Fairbury, and the State Reformatory for Women in neighboring Livingston County. Individual production rates for these wells range from 75 to 200 gpm.

Total groundwater pumpage from all bedrock formations is estimated to be 0.2 mgd. Greater quantities of water could certainly be withdrawn without overdevelopment. However, because large quantities of groundwater are generally available from the overlying glacial aquifers throughout the county, the bedrock formations will probably always be explored and developed only as a last resort.

*Chemical Character.* The chemical quality of water from the bedrock formations varies considerably both areally and with depth. Results of analyses of water from the bedrock are included in appendix B.

Water from the Pennsylvanian age rocks in Ford County is generally highly mineralized. Restricted circulation of water through the relatively tight sandstone and limestone layers of these rocks probably accounts for the high degree of mineralization and the normal increase in chemical constituents with depth. Water obtained from depths below 300 or 400 feet is likely to be too highly mineralized for domestic use. Analysis number 116366 from the Cabery Village Well No. 2 generally illustrates the quality of water available from the Pennsylvanian rocks.

Analysis No. H6366

<u>Chemical constituent</u>	<u>Symbol</u>	<u>Concentration (ppm)</u>	<u>Chemical constituent</u>	<u>Symbol</u>	<u>Concentration (ppm)</u>
Iron	Fe	1.4	Silica	SiO <sub>2</sub>	14.3
Manganese	Mn	Tr	Fluoride	F	1.0
Calcium	Ca	171.5	Chloride	Cl	21.0
Magnesium	Mg	67.5	Nitrate	NO <sub>3</sub>	5.4
Ammonium	NH <sub>4</sub>	1.4	Sulfate	SO <sub>4</sub>	1059.8
Sodium	Na	278.8	Alkalinity (as CaCO <sub>3</sub> )		180
			Hardness (as CaCO <sub>3</sub> )		707
			Total dissolved minerals		1744

Although few data are currently available concerning the chemical quality of water from the Silurian and Devonian age formations in Ford County, regional

data suggest that chemical constituents similar to those noted in water from the overlying glacial materials should be expected. Water contained in these formations throughout the county normally originates from recharge areas where the Pennsylvanian rocks are missing and a free exchange of water from the overlying drift materials occurs. Recharge from the Pennsylvanian rocks in other areas is usually prohibited by the presence of tight shale layers commonly found near the base of that formation. Analysis number 109960 of water from Central Soya Well No. 3 generally illustrates the quality of water which may be expected from the Silurian and Devonian formations.

Analysis No. 109960

<u>Chemical constituent</u>	<u>Symbol</u>	<u>Concentration (ppm)</u>	<u>Chemical constituent</u>	<u>Symbol</u>	<u>Concentration (ppm)</u>
Iron	Fe	0.9	Silica	SiO <sub>2</sub>	21.3
Manganese	Mn	0.0	Fluoride	F	0.4
Calcium	Ca	58.8	Chloride	Cl	1.0
Magnesium	Mg	33.5	Nitrate	NO <sub>3</sub>	0.6
Ammonium	NH <sub>4</sub>	5.6	Sulfate	SO <sub>4</sub>	2.5
Sodium	Na	49.7	Alkalinity (as CaCO <sub>3</sub> )		404
			Hardness (as CaCO <sub>3</sub> )		285
			Total dissolved minerals		406

The chemical quality of water from the deeper lying St. Peter sandstone in the northern portions of the county can only be inferred. Data from wells tapping this formation in nearby Livingston County suggest that water from the St. Peter sandstone in this part of Illinois might be similar to that obtained at Cullom (Fe = 0.5 ppm, Cl = 330 ppm, SO<sub>4</sub> = 72.0 ppm, hardness = 122 ppm, and total dissolved minerals = 975 ppm).

In many attempted developments in the past, the St. Peter sandstone has been erroneously accused of containing very highly mineralized water and in some cases a very strong hydrogen sulfide (H<sub>2</sub>S) or rotten egg odor. Several such cases have been the result of poor well construction allowing entrance of highly mineralized water and hydrogen sulfide from overlying formations to contaminate the St. Peter water. For this reason, careful quality monitoring and proper well construction are very important in attempting to finish a well in this formation.

WELLS

Types and Drilling Methods

Wells may be classified into types according to the method used in sinking the hole into the ground. The four types commonly found in Ford County are dug or augered, bored, driven, and drilled. The type of well constructed for a given location depends on the aquifer to be tapped and the needs and economic limitations of the user.

Dug or augered wells 2 to 5 feet in diameter are commonly used where water-bearing materials are not highly permeable (cannot transmit much water) and where they are less than 40 or 50 feet below land surface. Many of the large-diameter wells in use today are very old wells which were excavated with hand tools and lined with uncemented brick or stone. These wells are often subject to contamination by surface seepage and may be unsuitable for domestic use.

Current methods for constructing large-diameter wells involve the use of a rotary bucket drilling rig for the-excavating process. A large cylindrical bucket with auger type cutting blades on the bottom is rotated until the bucket is loaded with the materials being excavated. When full, the bucket is raised and swung aside to be dumped. Sections of precast large-diameter porous concrete tile are then placed to case the hole. This type of operation has proven most successful in areas where clay formations are present and caving of overlying materials into the bore hole is at a minimum. '

Bored wells 6 to 18 inches in diameter were commonly sunk prior to 1930 when home and farm demands were relatively small. Water enters this type of well only through the bottom opening which limits its yield capabilities. Because the bored well has a small capacity for receiving and storing water from the aquifer, it usually is inadequate for present-day water demands.

Continuous-flight spiral augers are normally used to construct bored wells. This method of drilling is also limited to regions where sufficient clay is present so that the bore hole will not cave in.

Driven wells, constructed by driving a small-diameter pointed screen and attached pipe directly into the aquifer, are feasible only where the aquifer is shallow (less than about 30 feet below the surface) and the overlying materials are easily penetrated. These wells usually are 1 to 2 inches in diameter, and are pumped by shallow well suction lift type pumps.

Drilled wells, which are most common in Ford County, may be constructed by the cable-tool or hydraulic-rotary methods. In the cable-tool method, the earth materials are broken into small fragments by the alternate raising and dropping of a heavy chisel-edged bit, and these fragments are removed from the hole at intervals by a bailer. In unconsolidated formations, an open hole is maintained by driving a stringer of casing as drilling progresses. After the aquifer has been penetrated, a well screen usually is placed opposite the water bearing formation, the casing pulled upward to expose the screen, and the screen sealed to the casing.

In the conventional hydraulic-rotary method, the drill pipe with a bit attached to the lower end is rotated to break the material into small particles. A thin mud is pumped through the drill pipe, out through openings in the bit, and up to the surface through the space between the drill pipe and the walls of the hole. The circulating mud thus removes the drill cuttings and prevents caving by plastering and supporting the formations penetrated with a thin mud cake on the bore wall until the final well casing and screen are placed in the hole.

In reverse hydraulic-rotary drilling, the flow of the drilling fluid is reversed from that in the conventional rotary method. The drilling fluid, usually a relatively clear water rather than mud, moves slowly down through the opening between the drill pipe and the bore hole, picks up formation cuttings loosened by the drill bit, enters the drill pipe through holes in the bit, and by suction

pumping moves to a surface pit where the cuttings settle. From there the clear water is again circulated down the hole. The fluid level in the hole must be kept at ground level at all times, since the difference in pressure between the water column in the hole and that in the aquifer prevents caving of the hole until the final well casing with attached screen is installed.

Construction Features

Construction features vary with the type of well and the characteristics of the aquifer to be utilized. Some of the features commonly employed in various types of wells in the county are casing, screening, gravel packing, and grouting (figure 12). These features, along with those for a dug well shown in detail in figure 13, are recommended by the Bureau of Public Water Supply of the Illinois Environmental Protection Agency. They are designed to minimize contamination from the surface. Detailed requirements are given in the rules and regulations of the Illinois Water Well Construction Code.

**Casing.** Wells are cased to maintain an open hole and to assist in protecting the quality of the water supply. Wells penetrating bedrock aquifers are cased opposite the overlying unconsolidated materials (figure 12a) and opposite any bedrock formations subject to caving. Drilled wells tapping water-bearing sand and gravel deposits are cased to the top of the well screen (figures 12b and c). Steel casing is used in drilled wells and some large-diameter dug wells;

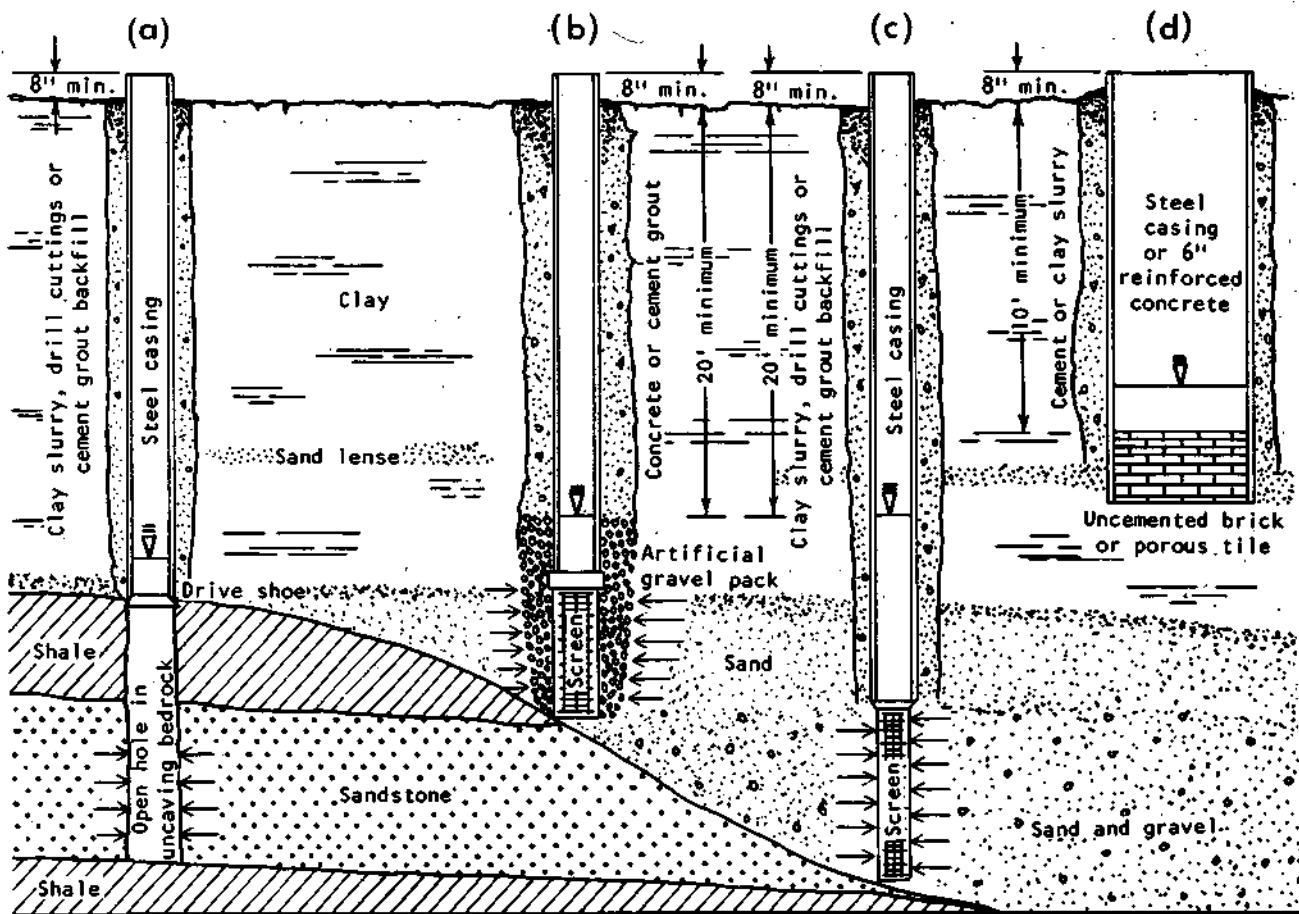


Figure 12. Construction features used in Ford County wells



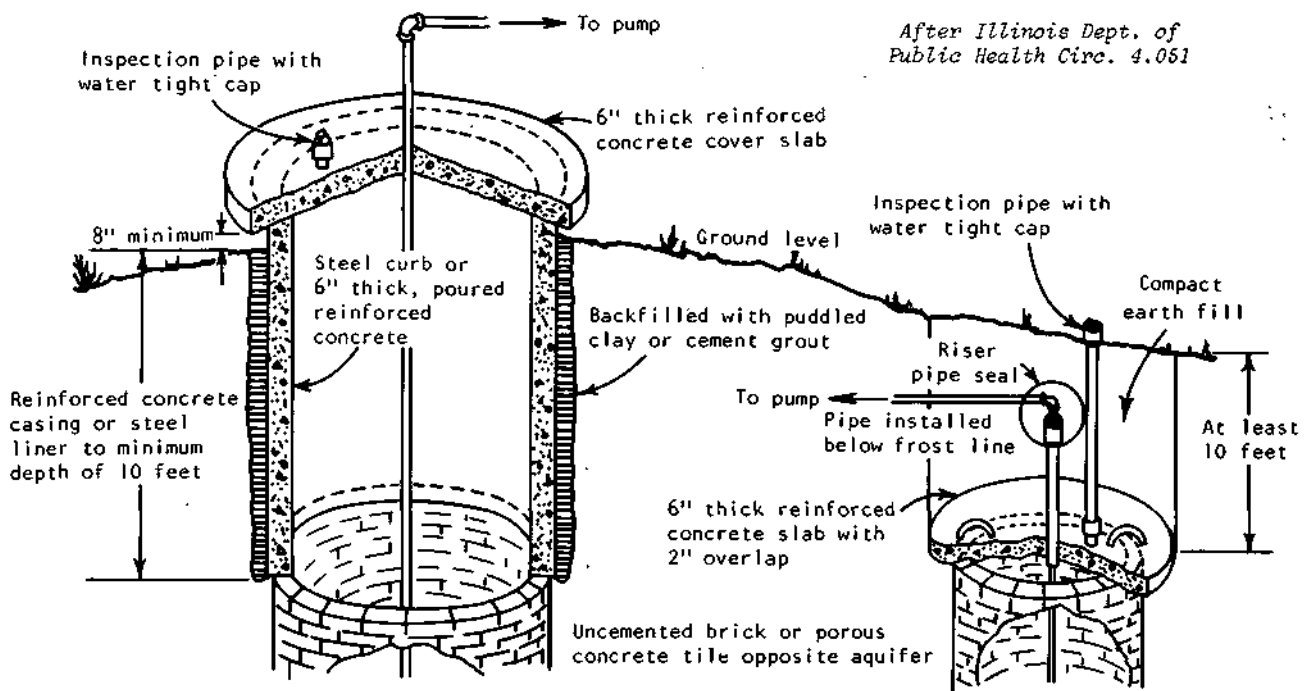


Figure 13. Recommended construction features for large-diameter dug wells

bored and dug wells may be cased with 6-inch thick reinforced concrete from at least 8 inches above ground level to a minimum depth of 10 feet below land surface with the lower portion usually lined with clay or porous concrete tile or uncemented brick (figures 12 and 13). Pitless adaptor units or surface well seals are normally used to provide sanitary protection to small-diameter wells equipped with steel casings. Sanitary protection for bored and dug wells also may be provided by placing a concrete slab at a point at least 10 feet below ground level and by filling in above the slab with compacted earth (figure 13).

*Screening.* Most successful drilled wells tapping sand and gravel are equipped with a length of commercially made well screen placed opposite the water-bearing formation (figures 12b and c). A properly selected and installed screen is designed to retain the aquifer material yet permit water to freely enter the well. Torch-cut and hand-sawed slotted casing sometimes is substituted for commercially made well screens because of the cheaper initial cost; this practice is not recommended because the openings in such a casing are usually too large to retain the aquifer material and too few to allow free water flow into the well. Most wells so equipped have a history of silt or sand pumping, low yield, and short production life, thus often proving to be more costly on a long-term basis than the well equipped with a commercially made screen.

Wells finished in bedrock aquifers not subject to caving do not require well screens.

*Gravel Packing.* Drilled wells finished in sand and gravel (figure 12c) usually are equipped with a screen that will retain the coarser 30 to 60 percent of the aquifer material immediately adjacent to the well screen; the remaining fine grains are removed from this area by surging, pumping, and bailing. This development practice causes a layer of coarse materials (a natural gravel pack) to accumulate around the screen. If the aquifer is uniformly fine-grained

(figure 12b) and the natural development methods are not possible, an artificial gravel pack envelope at least 6 inches thick may be required around the outside of the screen to prevent migration of fine material into the well. The grain size of this gravel pack should be about five times as large as the average grain size of the water-bearing material.

Some drillers partially fill the well casing with large-diameter gravel in an attempt to hold back the aquifer and term this gravel packing. This procedure is a very poor substitute for a true gravel pack, as it greatly reduces the yield-capability of the well and usually permits fine-grained materials to move into the well to plug it up or "chew up" the pump.

*Grouting.* The annular space between the casing and the bore hole must be sealed to minimize the chance of contamination from the surface. In drilled, dug, or bored wells (figures 12a, c, d) a clay slurry or cement grout must be used to seal the opening between the casing and the upper part of the bore hole. For artificial gravel-pack wells (figure 12b), a cement or concrete grout is required to insure an adequate seal. Minimum depths for grouting depend on the materials penetrated at the well site and vary from 10 feet for dug wells to 20 or more feet for the drilled wells.

### Disinfection

New wells, or old installations after rehabilitation, usually are bacterially contaminated and should be disinfected before being placed in service. After the disinfection is completed, the well should be sealed to safeguard against future contamination. The Illinois Environmental Protection Agency, Bureau of Public Water Supply, recommends disinfection procedures using a **strong chlorine laundry bleach** (5.25 percent chlorine). The correct amount to use can be determined from table 2, as explained in the instructions which follow.

- 1) Measure the depth of water in the well if possible. (Considering the well full of water will be satisfactory in most cases since a slight overdose does no harm.)
- 2) Determine the amount of laundry bleach (from table 2) and mix it in about 10 gallons of water. For example, a 6-inch diameter well with 75 feet of water would require 3 cups of laundry bleach.
- 3) Pour this solution into the well between the casing and the drop pipe. (This may involve raising the pump about 4 inches to allow sufficient space for the addition of the solution and for the placement of a sanitary well seal.)
- 4) Connect one or more hoses from faucets on the discharge side of the pressure tank to the top of the well, and while pumping the well, let water from these flow back into the well for at least 15 minutes. Then open each faucet in the system and let the water run until a chlorine odor or taste is detected. Close all faucets. Seal the top of the well casing.
- 5) Let stand for several hours, preferably overnight.

- 6) Operate the pump, discharging water from all outlets until all chlorine odor and taste disappears. Faucets on fixtures discharging to septic tank systems should be throttled to a low flow to avoid overloading the disposal system.

Table 2. Recommended Chlorine Dosages for Well Disinfection

Diameter of well (inches)	Amount of chlorine ( <i>cups</i> ) for given depth of water in well ( <i>feet</i> )					
	5	10	25	50	75	100
2	0.5	0.5	0.5	0.5	0.5	0.5
3	0.5	0.5	0.5	0.5	1.0	1.0
4	0.5	0.5	0.5	1.0	1.5	2.0
6	0.5	0.5	1.0	2.0	3.0	4.0
8	0.5	1.0	2.0	3.5	5.5	7.0
10	1.0	1.5	3.0	6.0	9.0	12.0
12	1.0	2.0	4.0	8.0	12.0	16.0
18	2.0	3.5	9.0	18.0	27.0	36.5
24	3.0	6.5	16.0	32.5	48.5	64.5
30	5.0	10.0	25.0	50.5	76.0	----
36	7.0	14.5	36.0	72.5	----	----
48	13.0	26.0	64.5	----	----	----
60	20.0	40.3	----	----	----	----

Chlorine always should be used outside or in well-ventilated places because breathing the fumes is dangerous. In heavy concentrations, chlorine also is harmful to the skin and clothing.

Additional instructions on safe water supplies from wells can be obtained from the Bureau of Public Water Supply of the Illinois Environmental Protection Agency (formerly Illinois Department of Public Health), Springfield.

#### Methods of Pumping Water

Most wells in Ford County are equipped with electrically driven pumps of the suction, jet, cylinder, or turbine types. Suction pumps can be used only where the pumping level is less than about 18 feet. For greater lifts, deep-well jet, cylinder, or vertical and submersible turbine pumps are required. Farm and domestic pumps generally are of the deep-well jet, cylinder, or submersible turbine types. Sizes of commercially available submersible pumps limit their use to wells with minimum inside diameters of 4 inches. Large-capacity municipal and industrial wells in the county utilize both the submersible and vertical turbine type pumps.

#### SUMMARY OF MAJOR WATER SUPPLIES

Municipalities and industries in Ford County pump about 1.4 mgd from wells. Groundwater pumped for farm and domestic purposes is estimated to be 0.8 mgd.

The major portion of municipal pumpage occurs at Gibson City and Paxton which pump about 500,000 and 365,000 gpd, respectively. In addition, industrial pumpage at Gibson City is an estimated 200,000 gpd. All municipal water supplies are obtained from wells tapping sand and gravel deposits of the drift materials with the exception of Cabery which obtains its water supply from the Pennsylvanian bedrock formations. Wells tapping the upper glacial deposits yield from 100 to 400 gpm; two wells tapping the lower deposits in the buried Mahomet Valley at Paxton yield about 1000 gpm. The bedrock wells at Cabery reportedly yield 125 . gpm.

The municipal and industrial groundwater supplies described in this section should provide a general indication as to what could be obtained from other such installations in areas where similar aquifer conditions are present. Probable maximum well yields can be estimated using specific capacity (yield per foot of drawdown) data in conjunction with available drawdown information at the area of interest. For example: Paxton City Well No. 6 has a reported specific capacity of 10 gpm/ft and nonpumping water level 102 feet below land surface. The top of the screen in this well is 133 feet below ground level providing an available drawdown of about 30 feet (133 feet minus 102 feet = 31 feet available drawdown). The estimated maximum short-term yield of this well should therefore be about 300 gpm (10 gpm/ft times 31 feet = 310 gpm).

In the descriptions below, population figures are taken from the 1960 census; pumping figures are the most recent available and in most cases are for 1969.

### Cabery

The municipal water supply for the village of Cabery (population 293) is obtained from two wells (Nos. 2 and 3), located on the south edge of town, that tap the Pennsylvanian age bedrock formations. Well No. 1, the old Park well, was abandoned in 1930.

The older well (No. 2) was drilled in 1920 by Lars Jensen, Clifton, to a depth of 233 feet below ground level. The well is cased with 200 feet of 6-inch pipe and is used only for standby purposes. The other well (No. 3) was drilled in 1956 by J. Bolliger and Sons, Fairbury, to a depth of 357 feet. The well is cased with 8-inch pipe from 3 feet above the surface down to 214 feet 10 inches, below which the hole was finished 8 inches in diameter. Upon completion, the well reportedly produced 125 gpm for 4 hours with a drawdown of 34 feet from a nonpumping water level of 48 feet (specific capacity = 3.6 gpm/ft of drawdown).

Analyses of water samples from the two wells show the water to have the following mineral constituents in parts per million (ppm) :

Well No.	Hardness	Iron	Total dissolved minerals
2	684	1.6	1406
3	640	0.2	1619

The water is not treated.

## Central Soya Company, Inc. (Gibson City)

Five wells have been drilled at the Central Soya Company processing plant. Three wells have been abandoned, and the plant now uses Well No. 4, tapping the bedrock, and Well No. 5, finished in the upper glacial materials, as a supplemental supply to water purchased from the city. Well No. 4 was drilled by L. F. Swanson and Sons, Gibson City, in 1947 to a depth of 552 feet below ground level. It is cased with 7-inch diameter pipe from 1 foot above land surface to 395 feet, below which is open hole. The well is equipped with a vertical turbine pump rated at about 150 gpm against 300 feet of head. Upon completion, the well produced 120 gpm for 30 hours with 59 feet of drawdown from a nonpumping water level 75 feet below ground level (specific capacity = 2.0 gpm/ft of drawdown).

Well No. 5 was constructed in 1947 by plant employees to a depth of 34 feet below ground level. It is cased with 60-inch porous concrete pipe from 2 feet above land surface to a depth of 34 feet. The well is equipped with a vertical turbine pump rated at 100 gpm against 75 feet of head.

Average daily pumpage is estimated to be about 200,000 gallons.

Analysis of a sample from Well No. 4 (appendix B, Lab. No. 110829) showed the water to have a hardness of 236 ppm, total dissolved minerals of 346 ppm, and an iron content of 0.6 ppm.

The water is not treated.

## Elliott

The village of Elliott (population 343) uses one well (No. 2) finished in the upper glacial formation as a source of municipal supply. Well No. 1 drilled in 1911 is no longer in use. Well No. 2 was constructed in 1950 by J. Bolliger and Sons, Fairbury, to a depth of 126 feet below land surface. The well is cased with 8-inch pipe to a depth of 120 feet and is equipped with 6 feet of No. 80 slot (0.080 inch) Johnson Everdur well screen. Upon completion, the well produced 120 gpm for 5 hours with a drawdown of 53.4 feet from a nonpumping water level of 69.8 feet below land surface (specific capacity = 2.2 gpm/ft of drawdown). This well is equipped with a 75-gpm vertical turbine pump.

Average daily pumpage is reported to be 20,000 gallons.

Analysis of a sample (appendix B, Lab. No. 144933) showed the water to have a hardness of **244** ppm, total dissolved minerals of 363 ppm, and an iron content of 2.1 ppm.

The water is not treated.

## Gibson City

The municipal water supply for Gibson City (population 3453) is obtained from three wells, located in the north end of town, tapping the upper glacial formation. Well No. 1 was drilled by American Water Company, Aurora, in 1927 to a depth of 58 feet below ground level. It has 10 feet of 24-inch screen with a gravel-pack annulus 7 inches thick. Wells No. 2 and 3 were drilled by Layne-

Western Company, Aurora, in 1939 and 1949 to depths of 56 and 58 feet below land surface, respectively. Both are equipped with 20 feet of 18-inch screen with gravel-pack annuli 9 inches thick. Pumping tests conducted on the wells indicate they have specific capacities from 22 to 24 gpm/ft of drawdown and should be capable of yielding 300 to 400 gpm each on a short-term basis. Pump capacities in the three wells range from 250 gpm to 400 gpm.

Average daily pumpage is reported to be 575,000 gallons, of which approximately 100,000 gallons per day is used by Central Soya Company.

Analyses of water samples from the three wells show the water to have the following mineral content in parts per million (ppm):

<u>Well No.</u>	<u>Hardness</u>	<u>Iron</u>	<u>Total dissolved minerals</u>
1	304	2.0	355
2	360	1.5	388
3	340	1.3	382

The water is fluoridated.

### Kempton

Four wells have been drilled at the village of Kempton (population 252). Wells No. 1 and 3 have been abandoned and filled. Wells No. 2 and 4, finished in the upper glacial formation, serve as the source for municipal water supply.

The older well (No. 2) was drilled in 1931 by E. W. Johnson, Bloomington, to a depth of 238 feet below ground level. It is 8 inches in diameter and is equipped with 10 feet of Johnson welded screen, the top 7 feet having No. 20 (0.020 inch) slot openings and the lower 3 feet having No. 30 (0.030 inch) slots. When completed, the well produced 110 gpm for 4 days with a drawdown of 100 feet from a nonpumping water level of 80 feet (specific capacity = 1.1 gpm/ft of drawdown).

Well No. 4 was drilled in 1962 by J. Bolliger and Sons, Fairbury, to a depth of 238 feet. It is an 8-inch well equipped with 5 feet 8 inches of No. 20 (0.020 inch) slot Johnson Everdur screen. Upon completion, the well produced 100 gpm for 4 hours with a drawdown of 61.5 feet from a nonpumping water level of 87.5 feet (specific capacity = 1.6 gpm/ft of drawdown). The long-term safe yields of the two wells are estimated as 75 and 50 gpm (108,000 and 72,000 gpd), respectively. Both wells are equipped with 100-gpm submersible turbine pumps.

Analyses of water from the two wells show the water to have the following mineral contents in parts per million (ppm):

<u>Well No.</u>	<u>Hardness</u>	<u>Iron</u>	<u>Total dissolved minerals</u>
2	734	2.1	1688
4	730	2.2	1776

The supply is continuously chlorinated for bacterial protection.

### Melvin

The village of Melvin (population 559) utilizes one well (No. 4), finished in the lower glacial formation, as a source of water supply. A second well (No. 3) is maintained for emergency use. Two earlier wells (Nos. 1 and 2) are no longer in service. Well No. 4 was constructed in 1954 by J. Bolliger and Sons, Fairbury, to a depth of 260 feet. It is an 8-inch well equipped with 20 feet 8 inches of No. 12 (0.012 inch) slot Johnson Everdur screen. Upon completion, the well produced 153 gpm for 4.5 hours with a drawdown of 13 feet from a nonpumping water level of 118 feet (specific capacity =11.8 gpm/ft of drawdown). Well No. 3 was drilled in 1923 by E. W. Johnson, Bloomington, to a depth of 265 feet. It is a 6-inch well, equipped with 25 feet of No. 10 (0.010 inch) slot screen. Wells No. 4 and 3 are equipped with 200-gpm and 60-gpm pumps, respectively.

Average daily pumpage is reported to be approximately 27,000 gallons.

Analyses of water from the two wells show the water to have the following mineral contents in parts per million (ppm):

<u>Well No.</u>	<u>Hardness</u>	<u>Iron</u>	<u>Total dissolved minerals</u>
3	335	0.6	427
4	316	0.7	426

The supply is continuously chlorinated for bacterial protection.

### Paxton

Eight wells have been drilled at the city of Paxton (population 4370). Four wells (Nos. 1, 2, 3, and 4) have been abandoned and sealed. The city now uses two wells (Nos. 7 and 8), located approximately 2 miles west of town, as the primary source for municipal water supply. These wells tap the lower glacial (Kansan) deposits. Wells No. 7 and 8 were drilled in 1956 and 1959 by J. P. Miller Artesian Well Company, Brookfield, to depths of 340 feet. Both are gravel-packed wells with 16-inch diameter casings, 10-inch thick annuli, and 100 feet of 16-inch No. 80 (0.080 inch) slot screen.

Two older wells (Nos. 5 and 6) located in town, which are also in the lower glacial (Illinoian) deposits, are maintained for emergency use. Well No. 5 was drilled in 1945 by Woollen Brothers, Wapella, to a depth of 149 feet. It is an 8-inch well and is equipped with 23 feet of Johnson Armco-iron screen. Well No. 6 was drilled in 1950 by Hayes and Sims, Champaign, to a depth of 153 feet. It is a 10-inch well and is equipped with 20 feet of Johnson screen. The top 6 feet of screen has No. 30 (0.030 inch) slot openings, the middle 5 feet has No. 25 (0.025 inch) slots, and the bottom 9 feet has No. 18 (0.018 inch) slots.

Upon completion the individual wells were tested to determine their potential yields. A summary of the test data is as follows:

Well No.	Pumping rate (gpm)	Length of test (hr)	Drawdown (ft)	Nonpumping water level (ft below land surface)	Specific capacity (gpm/ft)
5	155	1	13	101	12
6	200	4	20	102	10
7	800-1900	2k	1k	66	135 (avg)
8	1200	2k	8	68	150

Installed pump capacities for the four wells are: No. 5, 100 gpm; No. 6, 150 gpm; No. 7, 1000 gpm; and No. 8, 1000 gpm.

Average daily pumpage is reported to be 450,000 gallons.

Analyses of water samples from the four wells show the water to have the following mineral contents in parts per million (ppm):

Well No.	Hardness	Iron	Total dissolved minerals
5	194	1.4	311
6	295	1.3	451
7	332	1.6	368
8	350	1.7	418

The water is treated with phosphate to sequester the iron (help hold it in solution) and is continuously chlorinated for bacterial control. The water is also fluoridated.

### Piper City

Seven wells have been drilled at the village of Piper City (population 807). The first five wells constructed have been abandoned and the village now uses two wells, finished in the upper glacial formation, as a source of water supply. The two wells (Nos. 6 and 7) were drilled in 1944 and 1953 by Hayes and Sims, Champaign, to depths of 90 feet and 127 feet, respectively. Both wells are 16-inch gravel-packed wells with 8-inch diameter inner casings and screens. Well No. 6 is equipped with 19 feet of No. 60 (0.060 inch) slot Johnson Everdur screen.

Upon completion, Well No. 7 yielded 100 gpm for 1 hour with 8.5 feet of drawdown from a nonpumping water level 9 feet below land surface (specific capacity = 12 gpm/ft of drawdown). Wells 6 and 7 are equipped with vertical turbine pumps rated at 170 and 140 gpm, respectively.

Average daily pumpage is reported to be 55,000 gallons.

Analyses of water samples from the wells show the water to have the following mineral constituents in parts per million (ppm):



<u>Well No.</u>	<u>Hardness</u>	<u>Iron</u>	<u>Total dissolved minerals</u>
6	304	1.2	380
7	319	2.1	405

The water is not treated.

### Roberts

The village of Roberts (population 504) uses two wells (Nos. 5 and 6), finished in the lower glacial deposits, as a source of municipal water supply. Four earlier wells (Nos. 1, 2, 3, and 4) have been abandoned and filled.

The older well (No. 5) was drilled in 1950 by J. Bolliger and Sons, Fairbury, to a depth of 226 feet below land surface. It is an 8-inch well with 9 feet of continuous slot screen exposed to the water-bearing formation. The upper 3 feet has No. 10 (0.010 inch) slot openings and the lower 6 feet has No. 20 (0.020 inch) slot openings.

Upon completion the well was pumped at rates varying from 75 to 105 gpm for about 4 hours with 34.6 feet of drawdown from a nonpumping water level 81 feet below land surface (final specific capacity = 3 gpm/ft of drawdown). The well is equipped with a 105-gpm vertical turbine pump.

The other well (No. 6) was drilled in 1960 by J. Bolliger and Sons, Fairbury, to a depth of 228 feet below land surface. It is an 8-inch well with 13 feet of No. 16 (0.016 inch) slot Johnson Everdur screen. Upon completion, this well produced 128 gpm for 5 hours with a drawdown of 19.8 feet from a nonpumping water level of 86.6 feet from land surface (specific capacity = 6.5 gpm/ft of drawdown). The well is equipped with a 130-gpm submersible turbine pump.

Average daily pumpage is estimated to be 20,000 gallons.

Analyses of water samples from the wells show the water to have the following mineral constituents in parts per million (ppm):

<u>Well No.</u>	<u>Hardness</u>	<u>Iron</u>	<u>Total dissolved minerals</u>
5	446	1.6	685
6	426	1.2	681

The water is treated with a chlorinated solution of polyphosphate to sequester the iron (hold it in solution) and is continuously chlorinated for bacterial control.

### Sibley

The village of Sibley (population 386) uses one well (No. 1), finished in the upper glacial formation, as a source of municipal water supply. It was

constructed in 1907 by Otto Stiegman, Roberts, to a depth of 117 feet. It is an 8-inch well with 8.5 feet of enlarged No. 10 (0.010 inch) slot Cook screen. The well is equipped with a 58-gpm vertical turbine pump.

Analysis of a sample (appendix B, Lab. No. 116216) showed the water to have a hardness of 273 ppm, total dissolved minerals of 310 ppm, and an iron content of 0.8 ppm.

The water is not treated.

Stokely Van Camp Canning Co. (Gibson City)

The Stokely Van Camp Canning Co. uses three wells (Nos. 2, 3, and 4), all finished in the upper glacial deposits, as a source of water. One earlier well (No. 1) has been abandoned and sealed. Wells No. 2, 3, and 4 were drilled by L. F. Swanson and Sons, Gibson City, in 1932, 1951, and 1967 to depths of 47, 58, and 59 feet, respectively. Wells No. 2 and 3 are 6 and 8 inches in diameter, have 15-foot long screens, and are equipped with vertical turbine pumps capable of pumping about 400 gpm. Well No. 4 is an 8-inch well, is equipped with a total of 20 feet of No. 30 and 50 (0.030 and 0.050 inch) slot screen, and has a submersible turbine pump rated at about 600 gpm. Upon completion, the well reportedly produced 569 gpm for 4 hours with 2 feet of drawdown from a nonpumping water level 10.5 feet below ground level (specific capacity = 284.5 gpm/ft of drawdown).

No average pumpage data are available. However, peak usage during the canning season may approach as much as 750,000 gpd.

Water used in the processing plant is continuously chlorinated for bacterial control.

## EXAMPLE USE OF MATERIALS

The following brief discussion illustrates how the tables and maps in this circular may be used to evaluate the groundwater conditions at any given location in the county. Assume that a well is desired for a dependable farm water supply (5 to 10 gpm) in the Southeast corner of Section 31, Township 25 North, Range 9 East, Wall Township, Ford County (FRD 25N9E-31.1a).

A quick search of the data tabulated in appendix A shows three wells located in the section of interest. Two of these wells tap the upper glacial deposits at depths of 101 and 120 feet and one has a reported nonpumping water level about 90 feet below ground level. The third well is finished in the lower glacial deposits at a depth of 222 feet below ground level and also has a non-pumping water level 90 feet below ground level. Records of 28 additional wells located in the adjoining sections (24N8E-1, 24N9E-5 & 6, 25N8E-25 & 36, and 25N9E-29, 30, S 32) are also tabulated in appendix A. Twenty of these wells tap the upper deposits between depths of 60 and 143 feet and have reported nonpumping water levels from 50 to 96 feet below ground level. The remaining eight wells are finished in the lower deposits at depths from 210 to 260 feet and have reported nonpumping water levels from 75 to 156 feet below ground level. No records of wells tapping the underlying bedrock formations are recorded in the general area of interest.

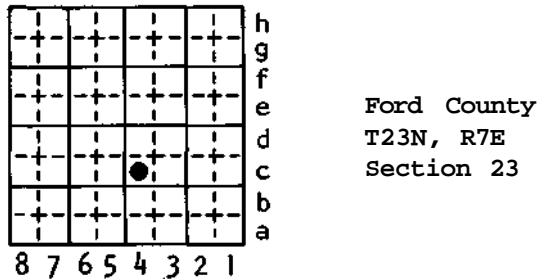
Most of the wells near this location are 2 or 4 inches in diameter and are equipped with lengths of commercially made screen designed to hold back the aquifer materials yet permit free entry of water into the well. Available information suggests that 3 or more feet of water-bearing sand (and screen) are normally required to insure an adequate farm supply. Although many of the wells now in use in Ford County are 2 inches in diameter, increased water usage and ease of pump maintenance make 4- or 6-inch wells more desirable.

The chemical quality of water from each aquifer is illustrated in appendix B. Water from the upper glacial deposits in the area of interest contains 3.5 ppm iron, 290 ppm hardness, and 488 ppm total dissolved minerals (Lab. No. 172369). Analysis of a sample (Lab. No. 116241) shows water from the lower deposits to have 0.6 ppm iron, 335 ppm hardness, and 427 ppm total dissolved minerals.

Maps in the text indicate that a satisfactory farm well can probably be developed from either the upper or lower glacial formations. A drilled well less than about 100 feet deep (figure 6) with a nonpumping water level near 90 feet below ground level (figure 7a) should be obtainable from the upper deposits. Total dissolved minerals would be expected to range between 400 and 500 ppm (figure 8b) with a hardness content between 300 and 400 ppm (figure 8c). Providing no deposits worthy of development are encountered at these depths, drilling into the lower deposits is recommended. A well tapping these materials should be less than about 240 feet deep (figure 9) with a nonpumping water level of about 90 feet below ground level (figure 10a). Total dissolved minerals and hardness content would be expected to range between 400 and 500 ppm and 300 and 400 ppm, respectively (figures 11b and c).

APPENDIX A - RECORDS OF WELLS

The well-numbering system used in this report is based on the location of the well, and uses the township, range, and section for identification. The well number consists of five parts: county abbreviation (FRD), township (T), range (R), section, and coordinate within the section. Sections are divided into rows of 1/8-mile squares. Each 1/8-mile square contains 10 acres and corresponds to a quarter of a quarter of a quarter section. A normal section of 1 square mile contains 8 rows of 1/8-mile squares; an odd-sized section contains more or fewer rows. Rows are numbered from east to west and lettered from south to north as shown in the diagram.



The number of the well shown is FRD 23N7E-23.4c. Where there is more than one well in a 10-acre square they are identified by arabic numbers after the lower case letter in the well number. Any number assigned to the well by the owner is shown in parentheses after the location well number.

In the listing of wells owned by municipalities, the place-name is followed by V, T, or C in parentheses to indicate whether it is a village, town, or city, except where the word City is part of the place-name.

Owners are listed according to the most current information available -- the 1969 plat book and recent well records for Ford County.

Symbols and abbreviations shown indicate the following:

- = constructed before year given, exact date unknown
- \* = test hole not developed as well
- \*\* = abandoned well
- drv = driven well
- drl = drilled well
- bor = bored well
- drl-GP = drilled well, gravel packed

The types of wells and methods of construction used in Ford County, their susceptibility to surface contamination, and methods of disinfection are discussed in the text of this report.















### Appendix A (Continued)

Well number	Owner	Well			Screen			Land surface elevation (ft above msl)	Non-pumping water level (ft)	Draw-down (ft)	Pumping rate (gpm)	Observed specific capacity (gpm/ft)	Length of test (hr)	Water-bearing formation and depth (ft)	Driller
		Year constructed	Type	Depth (ft)	Diameter (in)	Length (ft)	Diameter (in)								
T23W, R14W (Continued)															
6.5h	M. Sanders**	1924	dr-1	106	2	--	--	750	38	--	--	--	--	Sand at 106	Herrington
6.8b	M. Gavlick Est.	1910	dr-1	112	2	--	--	750	95	--	4	--	--	Sand at 112	Stiegman
6.8h1	Wildred Kennow	--	dr-1	80	2	--	--	749	--	--	--	--	--	--	--
6.8h2	Wildred Kennow	1874	dug	32.4	36	--	--	749	9.8	--	--	--	--	Sand at 32.4	--
7.1b	W. T. Morrison Est.	1900	dr-1	102	2	--	--	770	2	--	--	--	--	Rock	Swartz
7.2d	Walter Congram	--	dr-1	106	2	--	--	765	60	--	--	--	--	Sand & gravel at 106	Rardin
7.2e	L. Sheltenburger	1883	dr-1	87.5	2	--	--	770	55	--	--	--	--	Sand at 87.5	Sheltonburger
7.5h	Fred Menzke	1958	dr-1	--	2	--	--	770	--	--	--	--	--	--	--
7.6a	J. A. Johnson Est.	--	dr-1	106	2	--	--	773	--	--	--	--	--	--	Rardin
8.2a	Doris W. Johnson	1921	dr-1	115	2	--	--	750	50	--	--	--	--	Sand & gravel at 115	Weburg
8.3h	Lourence Ulrich	--	dr-1	98	2	--	--	759	--	--	--	--	--	--	--
8.8h	R. R. Hutchison	1917	dr-1	90	3	--	--	750	75	--	--	--	--	Sand & gravel at 90	Stiegman
9.3a1	Mary J. Peterson**	1894	bor	21	10	--	--	743	14	4	--	--	--	Sand & gravel at 21	McCornick
9.3a2	Mary J. Peterson**	1894	bor	80	10	--	--	743	40	5	--	--	2	Sand & gravel at 80	McCornick
9.6a1	Ed Brocksmith	--	dr-1	124	2	--	--	740	40	--	--	--	--	--	--
9.6a2	Ed Brocksmith	--	dr-1	150	2	--	--	740	--	--	--	--	--	--	Beck
9.6b	Elmer Rief	1944	dr-1	135	2	--	--	742	70-80	--	--	--	--	Sand at 135	Sharp & Stiegman
T24N, 47E															
1.1b	Miran Sibley	1957	dr-1	379	10-6	--	--	842	80	--	--	--	--	Sand & gravel, 110-120; sand, 225-240, 250-260; sand & gravel, 260-279	Zink
1.4h	Louise Smith Corp.	1895	dr-1	110	2	--	2	850	65-50	--	--	--	--	Sand at 110	Swanson
1.7a1	Urling S. Iselin	1899	dr-1	120	2	--	2	850	85	--	--	--	--	Sand at 120	Swanson
1.7a2	Urling S. Iselin	1963	dr-1	283	4	4	4	.008-.012	850	149	42	15	4	Dirty sand, 149-275; sand, 275-283	Swanson
1.7b1	Louise Smith Corp.	1904	dr-1	100	2	--	2	845	45	--	--	--	--	Sand at 100	Swanson
1.7b2	Louise Smith Corp.	1927	dr-1	--	6	--	6	845	--	--	--	--	--	Sand at 100	Swanson
2.5b	1st National Bank, Champaign, Tr.**	1921	dr-1	100	2	--	--	830	45	--	--	--	--	Sand at 100	Swanson
2.6g	1st National Bank, Champaign, Tr.**	--	dr-1	100	2	--	2	810	50	--	--	--	--	Sand at 100	Swanson
3.1d	Anne Sibley	1915	dr-1	100	2	--	2	820	70	--	--	--	--	Sand at 100	Swanson
3.8a	Anne Sibley	1890	dr-1	104	2	--	2	820	55	--	--	--	--	Sand at 104	Swanson
4.1g	Oscar A. Fentus	--	dr-1	18	4	--	4	785	5	--	--	--	--	Sand at 18	--
4.3a1	Suzanne Stroth Est.	1894	dug	35	40	--	--	802	29	--	--	--	--	Sand at 35	--
4.3a2	Suzanne Stroth Est.	--	dr-1	35-40	2	4	2	802	--	--	--	--	--	Sand at 40	--
4.7a1	F. G. Bielfeldt**	1893	dr-1	35	2	--	2	792	12	--	--	--	--	Sand at 35	Swanson
4.7a2	F. G. Bielfeldt**	1948	dr-1	31.5	4	--	--	792	13	5	14	2.8	3	Dirty sand & gravel, 20-29; sand & gravel, 29-31.5	Swanson
4.7a3	F. G. Bielfeldt	1964	dr-1	52	6	--	6	792	--	--	--	--	--	Sand at 52	Taylor
5.5a	Louise Smith Corp.	1905	dr-1	18	2	--	2	775	5	--	--	--	--	Sand at 18	Swanson
6.1b	C. Morris	1910	dr-1	195	4	--	4	780	15	--	--	--	--	Sand at 195	Swanson
6.1f1	James S. Watson, Jr.	1910	dr-1	16	2	3	2	772	5	--	--	--	--	Sand, 15-16	Swanson
6.1f2	James S. Watson, Jr.	1962	dr-1	140	4	--	4	772	--	--	--	--	--	Sand at 140	Taylor
T24W, 47E															
7.2a	Don & Cary Busling	--	dr-1	90	2	--	2	788	15	--	--	--	--	Sand at 90	--
7.4h	Wilton & Bates**	1930	dr-1	152	2	--	--	772	10	--	--	--	--	Sand at 152	Swanson
7.6a	Louise Stroth	1916	dr-1	29	2	--	2	780	2	--	--	--	--	Sand at 29	Anderson
8.1c	Esther Vehrs	1951	dr-1	80	4	--	--	.014	810	40	15	15	1.0	Sand, 40-45, 76-80	Swanson
8.1f	Berthe Davidson	1894	dr-1	108	2	--	2	792	25	--	--	--	--	Sand at 108	--
8.7h1	Charibel Fracty	--	dr-1	19	4	--	4	780	17.4	--	--	--	--	Sand at 19	--
8.7h2	Charibel Fracty**	1931	dr-1	42	4	4	4	780	10	--	--	--	--	Sand at 42	Kaffer
9.5h	Harold Sheppelman	1895	dr-1	45	2	--	2	791	15	--	--	--	--	Sand at 45	--
9.8c	Norman B. Ashley	1895	dr-1	90	2	--	2	802	20	--	--	--	--	Sand at 90	--
10.1a	H. & E. Mussell	1900	dr-1	125	2	--	2	813	65	--	--	--	--	Sand at 125	Swanson
10.1h	1st National Bank, Champaign, Tr.	--	dr-1	150	2	--	2	830	65	--	--	--	--	Sand at 150	Swanson
10.5h	Louise Smith Corp.	1950	dr-1	96.5	4	--	--	.018	820	52	21	18	.9	Sand & gravel, 74-94; sand, 94-96.5	Swanson
T25W, 47E															
10.8b	Louise Smith Corp.	1933	dr-1	103	2	--	--	831	55	--	--	--	--	Sand at 103	Swanson
11.1b	Urling S. Iselin	1948	dr-1	130	4	--	--	831	84	10	8	.8	6	Sand, 127-130	Swanson
11.4h	Louise Smith Corp.	--	dr-1	138	2	--	--	832	60	--	--	--	--	Sand at 138	Swanson
11.8f	Louise Smith Corp.	--	dr-1	100	2	--	--	833	25	--	--	--	--	Sand at 100	Swanson



























Appendix A (Concluded)

Well number	Owner	Well			Screen			Land surface elevation (ft above mol)	Non-pumping water level (ft)	Draw-down (ft)	Pumping rate (gpm)	Observed specific capacity (gpm/ft)	Length of test (hr)	Water-bearing formation and depth (ft)	Driller	
		Year constructed	Type	Depth (ft)	Diameter (in)	Length (ft)	Diameter (in)									Slot size (in)
T29M, R9E (Continued)																
23.6h3	Dorothy Moroff	1959	drl	376	5	--	--	722	70	120	2	.02	--	Dirty sand, 180-191; limestone with shale beds, 191-376	Griffy	
24.1a1	Helen L. Hughes	1889	drl	976	6-3	--	--	737	84	--	5	--	--	--	Sykes	
24.1a2	Helen L. Hughes	1960	drl	1250	6-5	--	--	740	--	--	--	--	--	Limestone, 210-440, 515-1025; sandstone, 1025-1245	Mehling	
24.1b	Helen L. Hughes**	1925	drl	523	--	--	--	740	--	--	--	--	--	Limestone, 184-523	Jensen	
24.3h	R. N. Hughes	1917	dug	84.4	30	--	--	731	9.2	--	--	--	--	--	--	
24.6g	Leslie Hummel	--	drl	143.5	4	--	--	720	44	--	--	--	--	--	--	
24.6e	Orville Rouk	1963	drl	326	6	--	--	740	75	--	15	--	--	Sand, 103-107; limestone, 216-223, 228-236, 240-252, 278-326	Bolliger	
25.3c	Mrs. J. Rinehart	1956	drl	--	6	--	--	750	--	--	--	--	--	--	--	
25.3h	Mrs. J. Rinehart**	--	drl	178	4	--	--	740	40	--	--	--	--	--	--	
25.6a1	Minnie Schaforth	--	dug	--	48	--	--	730	--	--	--	--	--	--	--	
25.6a2	Minnie Schaforth	1950	drl	234	6	--	--	.020	735	90	9	--	--	Sandstone, 218-234	Bolliger	
26.8a	Vern H. Dorn	1955	drl	220	6	--	--	740	35-40	--	--	--	--	--	Bolliger	
27.4a1	Wilson Bros.	--	dug	70	60	--	--	740	--	--	--	--	--	--	--	
27.4a2	Wilson Bros.	1954	drl	165	6	--	--	740	--	--	--	--	--	--	--	
28.3a	Cecil Cochran	1957	drl	380	4	--	--	745	--	--	--	--	--	Limestone, 180-280	Jensen	
29.1d	Earl Taylor Est.	1931	drl	230	4	--	--	740	--	--	--	--	--	--	Jensen	
29.1h1	Edward Ohrt	--	dug	52.5	54	--	--	735	26	--	--	--	--	--	--	
29.1h2	Edward Ohrt	1967	drl	470	6	--	--	737	--	--	--	--	--	--	Bolliger	
29.8a	J. & M. Donaghue	--	dug & bor	75.2	48-12	--	--	731	21.4	--	--	--	--	--	--	
30	Eldon Sargeant	1953	drl	318	--	--	--	--	--	--	2	--	--	Sand at 152; rock, 152-170	Bolliger	
30.1b	Dean Benson	1924	drl	206	4	--	--	725	38	--	--	--	--	Sand at 206	Jensen	
31.1b1	Earl Taylor Est.	--	drl	240	4	--	--	740	40	--	4	--	--	Rock	--	
31.1b2	Earl Taylor Est.	1942	drl	81	6	--	--	730	--	--	--	--	--	Sand & gravel, 80-81	Bolliger	
31.1h1	James Malone	1899	drl	735	6-3-2	--	--	731	100	--	4.5	--	--	Lime rock, 220-260; rock, 260-735	Munger & Burns	
31.1h2	James Malone	1954	drl	236	--	--	--	.020	730	90	8	--	--	Sand, 212-215, 230-236	Bolliger	
31.1h3	James Malone	1960	drl	213	6	--	--	.020	730	80	5	--	--	Sand, 153-155, 210-215	Bolliger	
31.5a(4)	Sampton (v)	1942	drl	238	8	6.7	8	.020	740	87.5	61.5	100	1.6	4	Dirty sand, 210-231; sand, 231-238	Bolliger
32.1h1	John Leonard, Jr.	--	drl	--	4	--	--	740	--	--	--	--	--	--	--	
32.1h2	John Leonard, Jr.	--	bor	--	10	--	--	740	--	--	--	--	--	--	--	
32.3a	Robert Wagner	--	drl	180	4	--	--	710	--	--	--	--	--	--	Matfield	
32.7a	Berens Bros.	--	drl	--	4	--	--	745	--	--	--	--	--	--	--	
33.1d1	Paul Weaver	--	drl	175	4	--	--	735	38	--	3	--	--	--	--	
33.1d2	Paul Weaver	1960	drl	400	4	--	--	735	84	16	9	.6	12	Limestone at 400	Jensen	
33.4g	Eula Farley	--	drl	209	4	--	--	740	50	--	--	--	--	--	--	
34.2h	Richard Morn	--	drl	150	3	--	--	732	35-40	--	--	--	--	--	--	
34.4a	Thomas Fenton	--	drl	200	4	--	--	707	45	--	4	--	--	Rock	--	
34.5h1	O. H. Basham	--	bor	109	18	--	--	750	30	--	--	--	--	--	--	
34.5b2	O. H. Basham**	--	dug	--	42	--	--	750	--	--	--	--	--	--	--	
34.8d1	Fred Crane	~1932	drl	300	6	--	--	715	--	--	1	--	--	Sand & gravel, 135-136	Bolliger	
34.8d2	Fred Crane	1932	drl	851	4	--	--	715	--	--	--	--	--	*Limestone, 226-660, 663-850	Burns	
34.8d3	Fred Crane	1955	drl	340	6	--	--	715	50	--	30	--	--	Sand, 108-112, 142-165, 187-201; lime rock, 201-240	Bolliger	
35.2h	Leroy Dalchman	--	drl	--	4	--	--	714	--	--	--	--	--	--	--	
35.7h	Luther Berratt	--	drl	170	4	--	--	702	30	--	--	--	--	Rock	--	
35.8c	Carrol Anderson	--	drl	80+	2.5	--	--	697	--	--	--	--	--	--	--	
36.1e	J. N. Bergen**	1925	drl	187	--	--	--	705	--	--	--	--	--	Sand & gravel at 187	Jensen	
36.1f	J. N. Bergen	1926	drl	221	4-3	--	--	705	30	--	--	--	--	Lime rock, 183-221	Jensen	
36.4h	J. N. Bergen**	--	drl	170	4	--	--	727	45	--	--	--	--	--	--	
36.6h	Irwin Taylor	1957	drl	--	6	--	--	740	--	--	--	--	--	--	--	



APPENDIX B - CHEMICAL QUALITY OF GROUNDWATER

Tabulated data of mineral content for groundwater supplies in Ford County follow.

Symbols used in the tabulations are:

- D = glacial drift
- L = limestone
- S = sandstone
- \* = State Bureau of Public Water Supply  
chemical analyses

The sources and significance of the major dissolved elements and substances in groundwater in waters of Ford County and U. S. Public Health Service drinking water limits (1962) are included in table 1.







Appendix B (Concluded)

Well number	Owner	Depth ft	Source	Laboratory number	Iron Fe	Manganese Mn	Ammonium NH <sub>4</sub>	Sodium Na	Calcium Ca	Magnesium Mg	Silica SiO <sub>2</sub>	Fluoride F	Nitrate NO <sub>3</sub>	Chloride Cl	Sulfate SO <sub>4</sub>	Alkalinity (as CaCO <sub>3</sub> )	Hardness CaCO <sub>3</sub>	Total dissolved minerals	Temperature °F
<b>T28N, R9E</b>																			
2.1d	Florence Schaub	100.6	D	171665	1.8	--	--	--	--	--	--	--	33.2	32	--	172	850	1586	63
6.5h1(1)	Kempton (V)	404	L	35950	.3	--	.5	351	5.7	1.5	15	--	.8	43	--	--	20	821	--
6.5h1(1)	Kempton (V)	404	L	50541	.2	.0	.4	309	6.6	2.3	14	--	.0	43	1	678	26	825	--
6.5h1(1)	Kempton (V)	404	L	72733	.0	--	.3	321	2.0	.8	7	3.7	.5	47	0	640	8	801	--
6.5h2(2)	Kempton (V)	238	D	83928	2.8	.0	1.5	257	165.9	90.2	11	--	.9	25	1140	128	785	1809	--
6.5h2(2)	Kempton (V)	238	D	116251	2.1	.1	1.6	241	162.4	79.6	12	.9	.5	21	1049	140	734	1688	55
6.5h2(2)	Kempton (V)	238	D	931539	1.5	--	--	--	--	--	--	.6	.0	54	1100	132	700	1260	--
6.5h3(3)	Kempton (V)	386	L	117100	--	--	--	--	--	--	--	1.0	--	30	--	288	579	1324	54
6.5h3(3)	Kempton (V)	386	L	117231	2.4	--	--	--	--	--	--	3.5	--	45	166	652	141	998	--
6.5h3(3)	Kempton (V)	386	L	931540	4.0	--	--	--	--	--	--	1.1	.0	36	1025	280	460	1220	--
7.1f	Dora Brown	200+	D	171851	2.6	--	--	--	--	--	--	--	--	20	--	128	562	1117	59
10.1b2	Harold Kelly	72.3	D	171844	.9	--	--	--	--	--	--	--	--	12	--	130	820	1334	61
11.3h	Louis Guatandi	97	D	171867	1.2	--	--	--	--	--	--	--	--	10	--	152	400	854	59
17.5a	Oswald Keller, Sr.	61.3	D	171854	.1	--	--	--	--	--	--	--	206.0	295	--	234	1660	2538	59
20.8a2	Howard Ferry	200	L	171853	4.4	--	--	--	--	--	--	--	--	20	--	156	708	1181	61
21.4a	Ford County Farm Service, Inc.	190	L	171966	8.6	--	--	--	--	--	--	--	--	167	--	248	96	560	65
22.8d2	John Riebe	206	S	171852	3.7	--	--	--	--	--	--	--	--	194	--	258	128	605	56
23.8d	W. & E. Malpin	40	D	171967	.3	--	--	--	--	--	--	--	45.8	19	--	254	374	514	59
24.4h	Donald Read	62	D	171959	.6	--	--	--	--	--	--	--	9.3	13	--	120	716	1506	59
29.8f	James McDermott	--	D	171856	Tr	--	--	--	--	--	--	--	95.6	65	--	260	650	1034	65
32.8e	M. & L. Flessner	56.7	D	171965	1.6	--	--	--	--	--	--	--	--	7	--	280	340	556	57
35.1h	J. Montellus Est.	180	S	171960	1.5	--	--	--	--	--	--	--	--	63	--	180	62	300	59
35.8b	Clara Switzer	40	D	171963	.8	--	--	--	--	--	--	--	--	64	--	238	364	752	58
<b>T29N, R9E</b>																			
13.1g	Emery DuBols	117.7	D	171632	.7	--	--	--	--	--	--	--	1.4	11	--	120	890	1571	--
16.1g2	Mrs. N. J. Selcher	275	L	171635	.2	--	--	--	--	--	--	1.8	--	22	--	358	43	675	59
16.6g	--	230	L	87831	1.5	--	--	--	56.4	26.2	--	--	--	21	535	248	249	1090	--
16.6h	Ernst Reising	230	L	87833	7.8	--	--	--	41.6	31.6	--	--	--	20	531	144	234	949	--
16.8g(3)	Cabery (V)	357	L	141967	.5	.0	--	--	--	--	1.0	.1	--	20	--	268	544	1260	54
16.8g(3)	Cabery (V)	357	L	152585	.2	.0	Tr	254	150.8	64.0	10	1.0	9.0	21	923	196	640	1619	54
16.8g(3)	Cabery (V)	357	L	91945	.3	.03	4.9	11	148	62	12	.9	.0	22	900	212	600	1636	--
16.8h(2)	Cabery (V)	233	L	92604	.8	--	--	--	--	--	--	--	--	21	--	172	820	1630	53
16.8h(2)	Cabery (V)	233	L	116366	1.4	Tr	1.4	279	171.5	67.5	14	1.0	5.4	21	1060	180	707	1744	55
16.8h(2)	Cabery (V)	233	L	140893	1.6	--	--	--	--	--	--	--	--	20	--	196	684	1406	--
18.1h1	F. & L. Paradies	40	D	171633	.4	--	--	--	--	--	--	--	397.0	315	--	224	2020	3147	52
19.4h	W. A. Hockett	122	D	171638	.1	--	--	--	--	--	--	--	118.0	51	--	234	890	1302	60
20.3h	Mrs. J. Jensen	100+	D	171634	2.6	--	--	--	--	--	--	--	--	13	--	92	1280	2435	--
22.8h	J. Ogilvie Est.	185	D	171637	.8	--	--	--	--	--	--	--	4.0	18	--	98	905	1919	60
24.3h	K. M. Hughes	84.4	D	171636	2.7	--	--	--	--	--	--	--	54.4	64	--	252	608	829	54
26.8d	Vern H. Down	220	D	171666	2.2	--	--	--	--	--	--	--	--	19	--	108	482	2519	58
29.1h1	Edward Ohre	52.5	D	171671	Tr	--	--	--	--	--	--	--	10.7	34	--	264	534	720	61
31.1b2	Earl Taylor Est.	81	D	171669	15.0	--	--	--	--	--	--	--	--	5	--	152	162	384	57
31.5a(A)	Kempton (V)	238	D	158802	2.2	.1	--	--	--	--	--	.5	.2	20	--	130	730	1776	54
34.5h1	O. H. Basham	109	D	171668	1.4	--	--	--	--	--	--	--	--	19	--	100	760	2027	61

## SELECTED REFERENCES

- Bennison, E. W. 1947. *Ground water, its development, uses and conservation*. Edward E. Johnson, Inc., St. Paul, Minnesota.
- Buswell, A. M., and T. E. Larson. 1937. *Methane in ground water*. Journal of the American Water Works Association v. 29(12):1978-1982.
- Csallany, Sandor, and W. C. Walton. 1963. *Yields of shallow dolomite wells in northern Illinois*. Illinois State Water Survey Report of Investigation 46.
- Ground water and wells*. 1966. Edward E. Johnson, Inc., St. Paul, Minnesota.
- Hanson, Ross. 1950. *Public ground-water supplies in Illinois*. Illinois State Water Survey Bulletin 40 (also Supplement 1, 1958, and Supplement 2, 1961).
- Horberg, Leland. 1950. *Bedrock topography of Illinois*. Illinois State Geological Survey Bulletin 73.
- Horberg, Leland. 1953. *Pleistocene deposits below the Wisconsin drift in northeastern Illinois*. Illinois State Geological Survey Report of Investigation 165.
- Illinois Department of Public Health. *Dug wells*. Division of Sanitary Engineering Circular 4.051.
- Illinois Department of Public Health. *Drilled wells*. Division of Sanitary Engineering Circular 4.052.
- Illinois Department of Public Health. 1967. *Water well construction code, rules and regulations*.
- Illinois State Water Survey. 1966. *Nitrate in water supplies*. Technical Letter 6.
- Larson, T. E. 1963. *Mineral content of public ground-water supplies in Illinois*. Illinois State Water Survey Circular 90.
- Meents, Wayne F. 1960. *Glacial-drift gas in Illinois*. Illinois State Geological Survey Circular 292.
- Piskin, Kemal, and R. E. Bergstrom. 1967. *Glacial drift in Illinois: thickness and character*. Illinois State Geological Survey Circular 416.
- Pree, H. L., Jr., W. H. Walker, and L. M. MacCary. 1957. *Geology and ground-water resources of the Paducah area, Kentucky*. U. S. Geological Survey Water Supply Paper 1417.
- Sanderson, E. W. 1967. *Groundwater availability in Shelby County*. Illinois State Water Survey Circular 92.
- Selkregg, Lidia, and John P. Kempton. 1958. *Ground-water geology in east-central Illinois*. Illinois State Geological Survey Circular 248.

- Smith, H. F. 1954. *Gravel packing water wells.* Illinois State Water Survey Circular 44.
- Salinity and livestock water quality.* 1959. South Dakota State College Agricultural Experiment Station, Brookings, Bulletin 481..
- Spafford, H. A., and C. W. Klassen. *Hazards of methane gas in water wells and suggested method for elimination.* Illinois Department of Public Health mimeograph report.
- Walton, W. C. 1960. *Leaky artesian aquifer conditions in Illinois.* Illinois State Water Survey Report of Investigation 39.
- Walton, W. C. 1965. *Ground-water recharge and runoff in Illinois.* Illinois State Water Survey Report of Investigation 48.
- Water for Illinois, a plan for action.* 1967. Illinois Department of Business and Economic Development, Springfield.
- Willman, H. B., and others. 1967. *Geologic map of Illinois.* Illinois State Geological Survey.