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DEPARTMENT OF REGISTRATION AND EDUCATION



Groundwater Availability in Shelby County

by E. W. SANDERSON

ILLINOIS STATE WATER SURVEY

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GROUNDWATER AVAILABILITY IN SHELBY COUNTY, ILLINOIS

by E. W. Sanderson

SUMMARY

Throughout Shelby County adequate water supplies for normal household and farm use generally can be obtained from wells tapping glacial deposits of sand and gravel. In the central part of the county, wells capable of producing more than 100 gallons per minute (gpm) have been developed from extensive deposits of permeable sand and gravel contained in the preglacial Kaskaskia valley. In 1966 more than 2 million gallons of water was pumped from wells each day to satisfy industrial, municipal, and domestic needs. A much larger quantity of groundwater can be withdrawn without overdevelopment. Maps and tables of data illustrating localized groundwater conditions and quality are included in this report to serve as a guide in the proper development and utilization of available groundwater resources.

INTRODUCTION

More than 500 requests for information concerning groundwater conditions in specific locations of Illinois are answered yearly by the Illinois State Water Survey. Approximately 40 percent of these requests are from individuals seeking advice on locating, developing, or treating home or farm groundwater supplies. Many of these requests are answered with letter-type reports prepared jointly by the State Water Survey and State Geological Survey from available geohydrologic data in the basic record files. These reports, containing pertinent information on groundwater and geologic conditions at a specific site, permit meaningful cost appraisals for well construction which have saved considerable time, effort, and money in many cases. However, several thousand wells are constructed each year without the use of such information. If comprehensive summaries of groundwater conditions were available for all possible sites, great savings could result. This report presents such a summary for Shelby County, where fairly complex groundwater conditions exist.

Shelby County is located in the south-central part of the state (figure 1 and plate 1). It encompasses an area of 772 square miles and is mainly cultivated land. According to 1960 figures, the county has a population of 23,404, and 11,697 of the residents live in incorporated cities and villages. The county seat and largest city, Shelbyville, has a population of 4821 (1960 census). All known water supplies in the county are from groundwater sources.

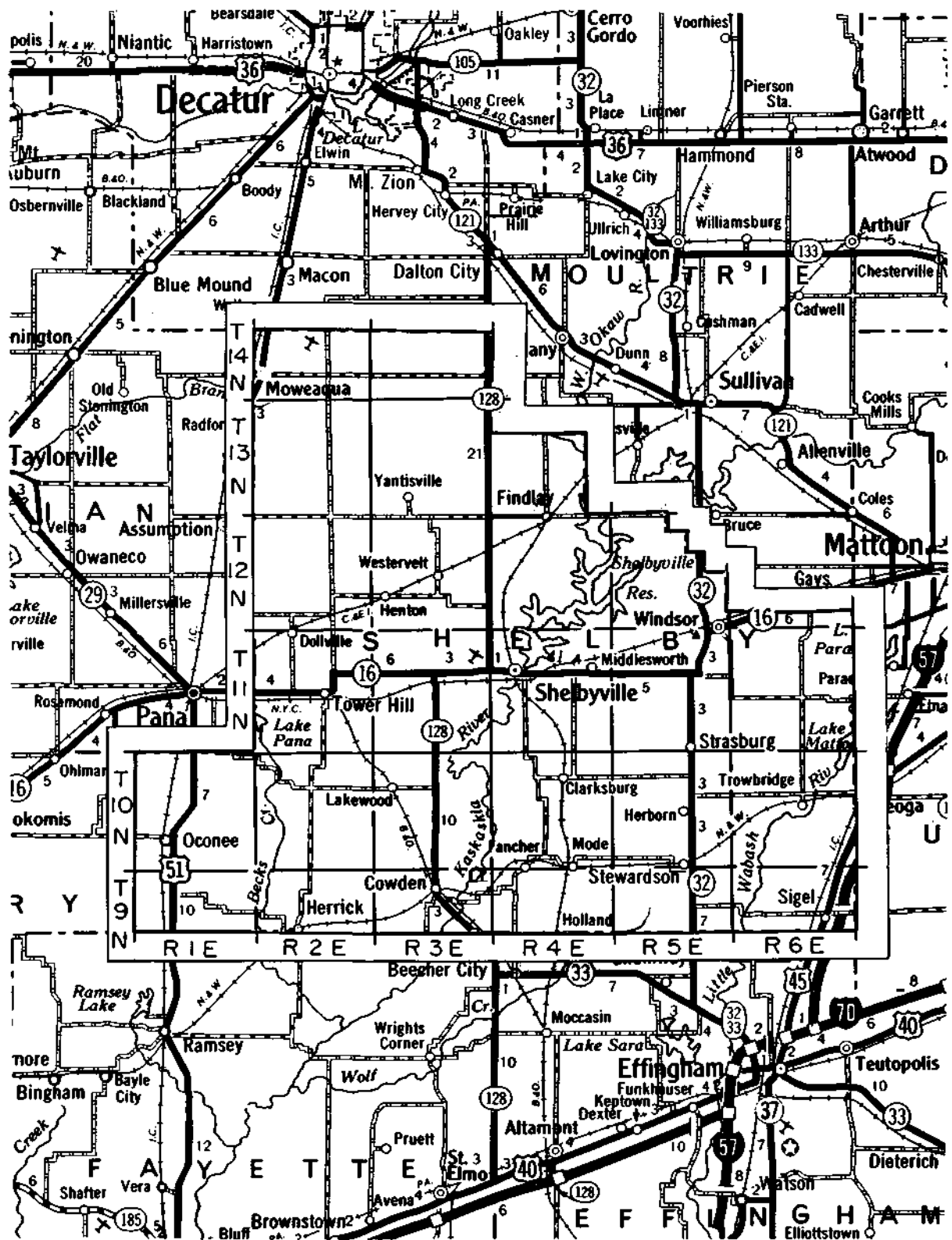


Figure 1. Location of Shelby County

Although the present economy of the county is largely dependent on the production of farm crops such as corn, soybeans, and wheat, a noticeable economic contribution from industrial, commercial, and recreational developments should be realized after the Shelbyville Reservoir on the Kaskaskia River is completed in 1968. These and other allied developments will create additional demands for dependable water supplies of good chemical quality. Surface water from the Shelbyville Reservoir undoubtedly will furnish part of that quantity. However, it is probable that groundwater supplies always will predominate in the county.

This report summarizes groundwater conditions in Shelby County including pertinent geologic factors, occurrence and movement, temperature and chemical quality, and well development. Appendix A lists the results of chemical analyses of all wells sampled, and appendix B contains available records of existing wells.

This report is part of a continuing program of water-resource investigations being conducted by the Illinois State Water Survey under the direction of William C. Ackermann, Chief, and H. F. Smith, Head of the Hydrology Section. Grateful acknowledgment is made to the many well drillers, engineers, and public officials who provided invaluable information for use in this report. Mrs. Dorothy Woller tabulated the well data and typed the manuscript, and John W. Brother, Jr. prepared the illustrations. The suggestions and constructive criticisms of William H. Walker were of great value in the preparation of this report.

GEOLOGY

The geology of Shelby County is summarized in general terms in State Geological Survey Circular 225, "Groundwater Geology in South-Central Illinois." The following brief discussion of geologic conditions in the county is taken largely from that publication. In addition to the summary in Circular 225, the files of the State Geological Survey are available for greater definition of the geology in this portion of the state.

Information from wells and exposures of rocks indicate that the land surface of Shelby County has been shaped principally by ice and running water. The features produced by ice were developed long ago when glaciers, nourished by snow accumulation in Canada, several times advanced across Shelby County and melted away leaving a vast quantity of rock debris. In front of the ice, sediment-laden meltwaters escaped down valleys, partially filling them with outwash deposits of sorted sand, gravel, and finer material. Thick extensive till sheets of unsorted clay, silt, sand, and pebbles also were laid down under the advancing ice or dumped in place during melting. Glacial deposits blanket practically all of Shelby County resulting in a relatively level plain broken only by isolated knobs, stream valleys, and long ridges formed at the front of the glacier (end moraines).

Running water continues to modify this surface today by cutting into the land, carrying away soil and rock particles, and depositing the debris in river

bottoms. This modification is a small-scale version of the changes made on the bedrock surface by glacial meltwaters.

Below the glacial deposits in Shelby County are layers of consolidated rocks representing several geologic ages. The uppermost consolidated rocks consist of beds of shale, sandstone, and limestone arranged one upon the other; the top surface of these rocks is called the bedrock surface. Originally the bedrock formations also were unconsolidated materials, deposited over many years as sediments in shallow seas or bordering marshes. They were then buried and hardened into solid rock during the several million years after the seas retreated from the area.

Erosion of the bedrock was not uniform through the county. In areas where soft shales and sandstone formations were exposed to weathering, valleys were formed by water and ice action, while hard sandstone and limestone formations in other areas resisted erosion and remained to form ridges and hills on the bedrock surface. Some of the old bedrock valleys coincide with present-day stream valleys, but some are partially or even completely buried by the glacial deposits so that there is little or no surface evidence of their presence. In parts of the county, the bedrock surface is exposed in dry washes and gullies in the higher lands, and in some of the creek and river valley lowlands.

GROUNDWATER

Groundwater in Shelby County begins as precipitation which seeps downward through the soils. Figure 2 shows the cycle of water movement from the atmosphere as precipitation to the surface and into the ground, and then away from the area either through the ground and in flowing streams or again into the atmosphere through transpiration of plants and evaporation.

Occurrence and Movement

Water enters and filters slowly down through the ground until it reaches a level where all available voids are completely water-filled. Water thus contained in this zone of saturation is groundwater, and its upper surface is the water table.

In glacial drift deposits, water fills the voids between the grains that make up the formations. In bedrock, water occurs primarily in two ways--it is contained in the spaces between partially cemented grains of sandstone strata or in the fractures, bedding planes, and solution cracks of limestone formations. A saturated formation of sand, gravel, sandstone, or limestone that is capable of yielding water to wells in usable quantities is called an aquifer.

Usually, glacial drift aquifers are regularly recharged (refilled) by rainfall occurring directly on the soil surface. If a glacial drift aquifer contacts

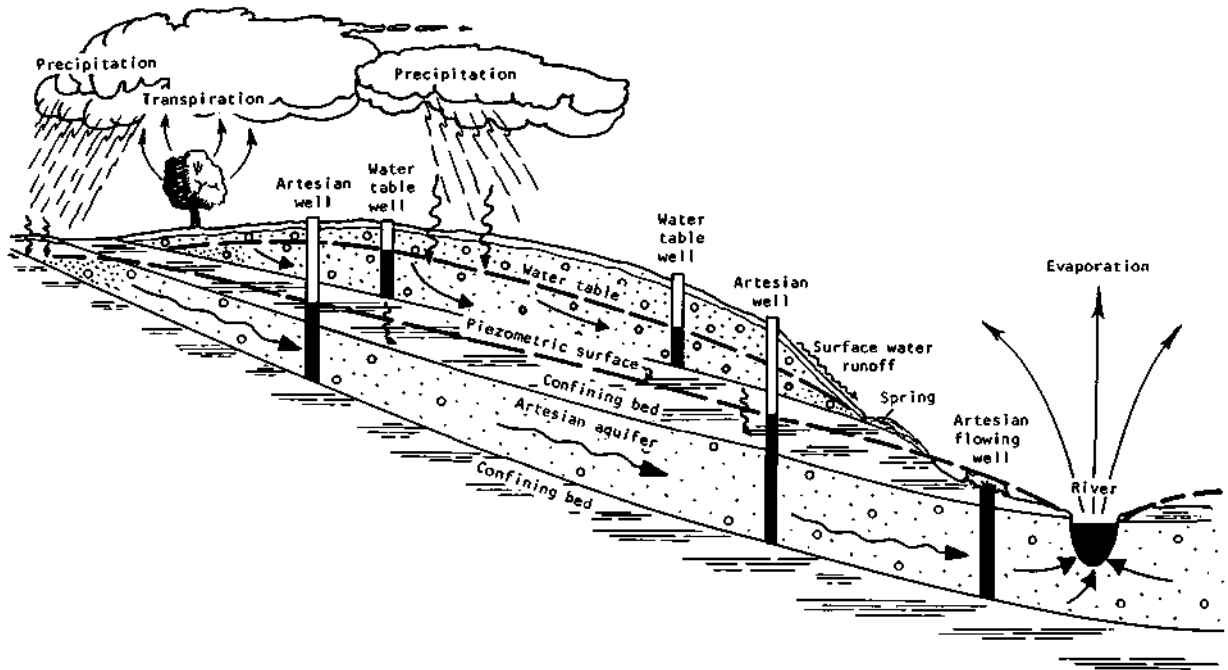


Figure 2. Cycle of water movement

a bedrock water-bearing unit below it, the water continues freely downward to recharge the bedrock aquifer. However, layers of very dense (almost impermeable) materials separating water-bearing units impede the downward movement of water. These layers, or confining beds, are usually clays or shales so dense that they cannot yield enough water to be classified as an aquifer. When such confining beds are present, most of the water reaching the aquifer may come from a distant recharge area where the confining beds no longer exist or where the aquifer crops out at the land surface.

Water entering permeable strata in an area of outcrop may become confined downslope beneath relatively impermeable beds. Pressure is exerted on the groundwater in a confined aquifer by the weight of water at higher levels in the aquifer. When a well penetrates such an aquifer downslope from the recharge intake area, the pressure forces the water to rise in the well above the top of the aquifer. The water in this instance is confined (or artesian) water, the well is an artesian well, and the upper surface of the water in the well is the piezometric surface. Thus, the piezometric surface is the level to which water from a given artesian aquifer will rise under its full head in tightly cased wells.

Groundwater moves under the influence of gravity or other pressure differences from recharge areas to discharge points of lower pressure. Major points of discharge are springs, lakes, streams, swamps, drainage tiles, or pumping wells.

The movement toward the discharge points may amount to only a few hundred feet a year in unconsolidated materials and to only a few feet a year in sandstone formations. Water may be held in bedrock aquifers for many years.

Availability

Glacial drift sand-and-gravel aquifers offer the best possibilities for developing groundwater supplies in Shelby County. Throughout much of the county, the upper bedrock is composed primarily of shale and other deposits which do not bear water. Only in the vicinity of Tower Hill and Clarksburg-Stewardson are there shallow fresh-water sandstone formations. In the remainder of the county, deeper lying bedrock aquifers yield water too highly mineralized for most purposes.

Plate 1 (see pages 8-9) depicts the availability of groundwater in glacial drift aquifers. Areas of excellent, good, fair, and poor groundwater possibilities are delineated. Also noted are the probable maximum depths for wells and the areas where the bedrock is known to contain fresh water. The information on plate 1 was based on groundwater data including: 1) the chemical analyses and well records in appendices A and B; 2) topographic and bedrock surface maps; 3) geologic reports; 4) electrical earth resistivity surveys; and 5) drillers logs from oil and coal test holes.

According to these data, groundwater availability in the county is largely controlled by geologic factors such as the thickness, nature, and origin of glacial drift deposits; their degree of interconnection with the upper bedrock aquifers; and the occurrence and permeability of water-bearing sandstone and limestone units in the bedrock.

Parts of the county covered by the Shelbyville moraine and areas having preglacial bedrock valleys filled with material from the glacial front are underlain by fairly extensive sheets or strips of water-bearing sand and gravel. Farm and home water supplies nearly always can be obtained from these aquifers, and the chances are also good for developing higher capacity wells for industrial and municipal use, although test drilling may be necessary to locate the more permeable and thicker sections of water-bearing material required for this type of installation. In the preglacial Kaskaskia valley, a major drainage way from the Shelbyville moraine of Wisconsinan age, wells capable of producing more than 100 gpm on a long-term basis have been constructed (see Shelbyville and Findlay municipal well data in appendix B).

In those parts of the county covered with about 40 or more feet of older glacial deposits (Illinoian), water-bearing sand and gravel formations are generally thinner, less permeable, and discontinuous. However, well records indicate that supplies adequate for normal domestic needs usually can be developed from small-diameter wells, although in some locations several test holes may have to be drilled to find a suitable aquifer. The chances of developing dependable supplies from drilled wells are poor only where the Illinoian drift mantle is

less than about 40 feet thick. In these areas, one or more large-diameter dug or augered wells may be required to furnish an adequate home or farm supply.

QUALITY

The sources and significance of the major dissolved elements and substances in groundwater and their ranges of concentration in waters of Shelby County are shown in table 1. The U. S. Public Health Service drinking water standards (1962) for these major constituents also are included in the table. These standards have been accepted by the American Water Works Association as minimal for public water supplies, and should serve as a guide to owners of farm and home water supplies in evaluating their water quality.


Chemical Character


As may be generally inferred from the information in table 1, the dissolved minerals in groundwater are derived chiefly from the earth materials through which the water flows. The soils and glacial materials above bedrock are particularly rich in calcium, magnesium, iron, and other minerals which are readily absorbed by the groundwater as it passes over and through these deposits. Calcium and magnesium are responsible for hardness of water, and iron causes reddish-brown staining.


Groundwater from glacial deposits throughout the county is generally considered very hard, from 250 to 450 parts per million (ppm), but normally hardness can be successfully removed by home water-softening units that are now readily obtainable. The iron content of these waters usually is between 1.0 and 5.0 ppm, well above the recommended limit of 0.3 ppm. Iron can be removed by units similar to home water softeners; however, for domestic users, tolerance rather than removal is the usual practice.


The deeper bedrock aquifers in Shelby County contain highly mineralized groundwater not suited for most farm and domestic uses. Chloride, sulfate, and sodium are present in especially high concentrations. In most cases, these bedrock aquifers are sandwiched between beds of shale, coal, and fire clay so that free exchange between these formations and the upper glacial drift containing fresh water has not occurred. For this reason, the highly mineralized water contained in these aquifers has not been flushed out by fresh water. Available geochemical data for the county indicate that groundwater from elevations lower than about 475 feet above sea level is unusable for most domestic purposes. Any hole drilled into the bedrock should be terminated at this elevation if no fresh water zones have been penetrated in the overlying rocks.


Water from wells in Shelby County contain varying quantities of carbon dioxide and, in some cases, methane gas. These gases are colorless, odorless, and tasteless. Methane gas is lighter than air whereas carbon dioxide is

 Excellent groundwater possibilities. Municipal water supply developments at Shelbyville and Findlay indicate that large quantities of groundwater for industrial, municipal, and farm use can be developed from wells tapping the thicker and more permeable parts of this aquifer. Wells generally range in depth from 50 to 150 feet, and yields in excess of 100 gpm have been obtained. Nonpumping water levels vary with topography from about 25 to 50 feet below ground level. The bedrock is usually not tapped as a source of supply.

 Good groundwater possibilities. Small-diameter wells predominate. From Henton north to the county line, wells are usually from 70 to 125 feet deep. In the Middlesworth-Windsor area, they range from 80 to 110 feet deep. Well yields of 10 gpm or more are common. Nonpumping water levels vary with topography from about 25 to 50 feet below ground level. The bedrock is usually not tapped as a source of supply, and water may be salty below depths of 75 feet in the lowland areas near the Kaskaskia River and below depths of 250 feet in the highlands of north and east Shelby County.

 Fair groundwater possibilities. Small-diameter drilled wells usually are 35 to 75 feet deep; the large-diameter dug wells present are 20 to 40 feet in depth. Most wells yield less than 10 gpm, but a few located on the floodplain of the Kaskaskia River have yielded 100 gpm or more. Nonpumping water levels vary with topography from about 25 to 50 feet below ground level. The bedrock is not usually tapped as a source of supply, and water may be salty below depths of about 50 feet in the Kaskaskia River bottoms and below depths of 200 feet in the highland areas between Findlay and Westervelt.

 Poor groundwater possibilities. Large-diameter dug wells ranging in depth from 15 to 35 feet predominate. Yields in excess of 5 gpm are rare and some wells go dry during long drought periods. Nonpumping water levels vary with topography and range from 10 to 25 feet below land surface.

 Areas known to be underlain by shallow sandstone aquifers containing fresh water. Near Tower Hill bedrock wells are around 50 feet deep and produce up to 9 gpm. Nonpumping water levels are about 10 feet below ground level. In the Clarksburg-Stewardson area wells usually are 75 to 175 feet deep, and yield from less than 3 to about 8 gpm. Nonpumping water levels vary with topography from about 15 to 30 feet below land surface. Water from elevations less than about 475 feet above sea level is too mineralized for most domestic uses.

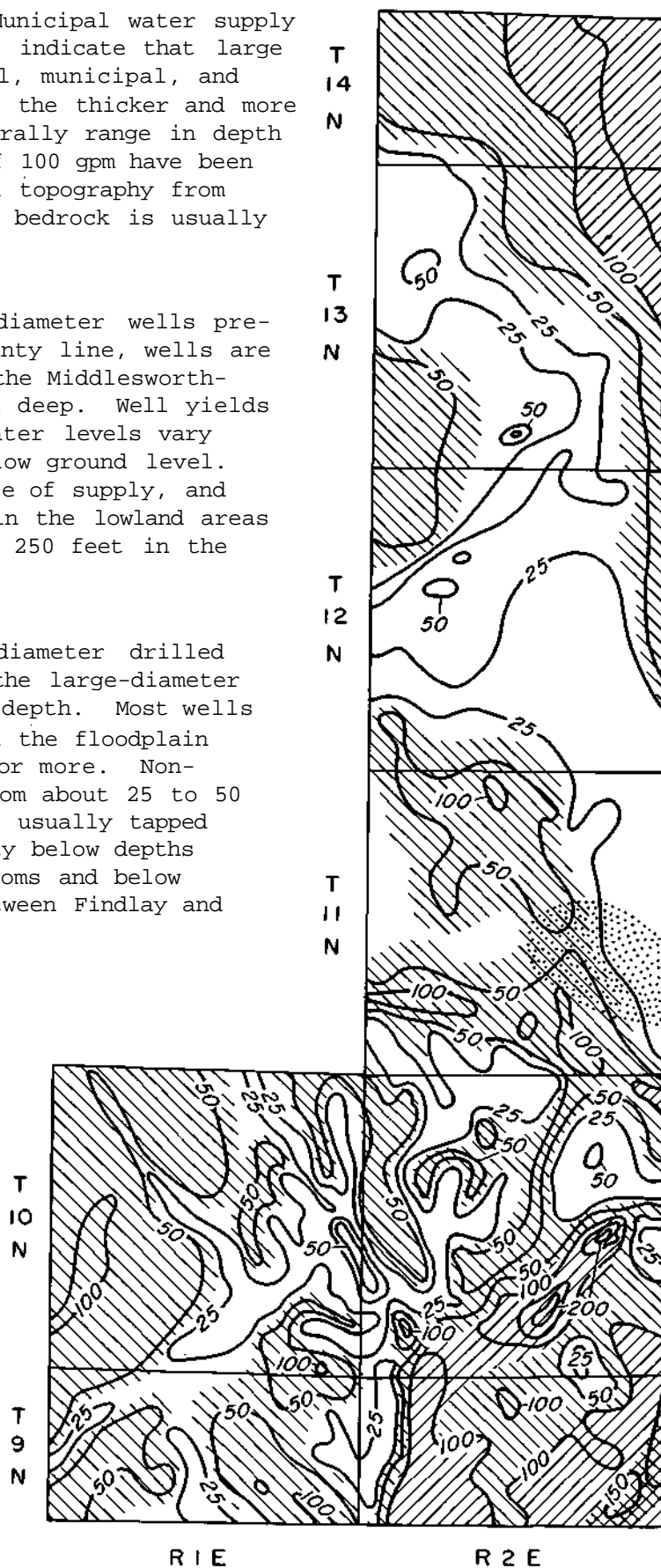
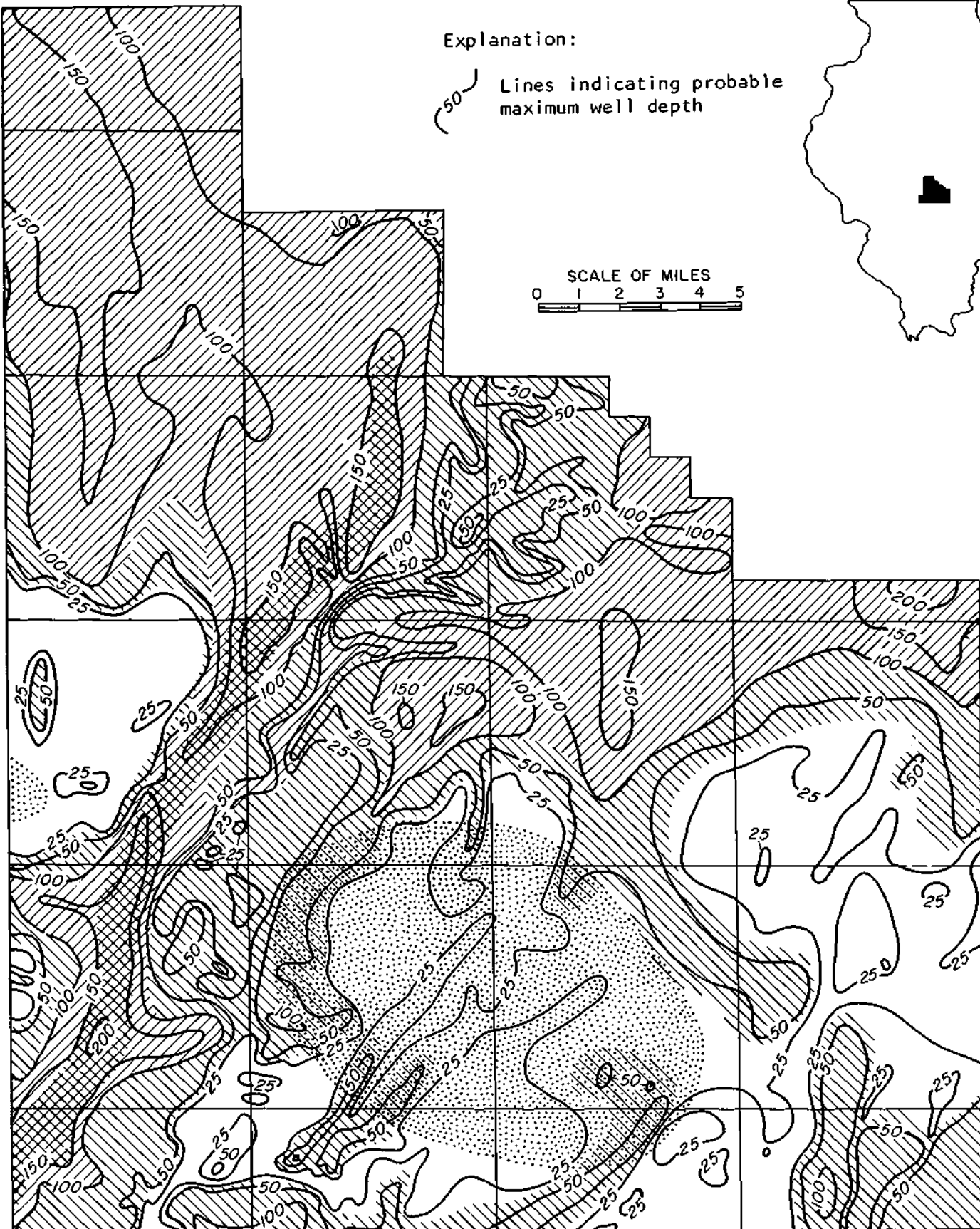
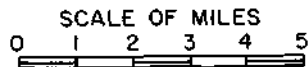


Plate 1. Groundwater availability and

Explanation:

Lines indicating probable maximum well depth



probable maximum depth of wells finished in glacial deposits

Table 1. Elements and substances commonly found in groundwater in Shelby County

<u>Constituent and recommended upper limit</u>	<u>Source</u>	<u>Remarks</u>
Iron (Fe) 0.3 ppm	Dissolved from common iron-bearing minerals present in practically all rocks, clays, and soils; may also be derived from iron pipes, pumps, and other equipment.	On exposure to air, iron oxidizes to a reddish-brown sediment. More than about 0.3 ppm stains laundry and porcelainware reddish brown; objectionable for food and beverage processing. Of 76 county samples analyzed, 56 had an iron content greater than 0.3 ppm.
Manganese (Mn) 0.05 ppm	From soils and sediments. Less abundant in rocks than is iron.	Resembles iron in chemical behavior and has same objectionable features, except stains are brown to black. Of 43 samples analyzed, 17 had more than 0.05 ppm manganese.
Nitrate (NO ₃) 45 ppm	Results from decayed organic matter such as that from barnyards, feedlots, manure piles, septic tank fields, as well as from silage juices and animal tissue. Usually occurs in waters from shallow wells of less than 50-foot depth, often as the result of poor well construction permitting drainage into the well at or near the surface.	Values higher than a few ppm may suggest pollution. More than about 45 ppm nitrate may cause methemoglobinemia (blue babies) when such water is used in preparation of infant feeding formulas. Of 53 samples analyzed, 4 showed more than 45 ppm nitrate content.
Chloride (Cl) 250 ppm	Dissolved from rocks and found in large amounts in ancient brines, sea water, and industrial brines.	In concentrations over about 250 ppm chloride gives a salty taste to water and increases its corrosiveness. Of 77 samples analyzed, 3 showed a chloride content greater than 250 ppm.
Sulfate (SO ₄) 250 ppm	Dissolved from rocks and soils containing gypsum, iron sulfides, and other sulfur compounds. Present in waters from coal mine drainage and some industrial wastes.	Sulfate in water containing calcium forms a hard scale. In large amounts, sulfate in combination with sodium or magnesium has a laxative effect, most noted by infrequent users of the water. Of 42 samples analyzed, 2 had a sulfate content greater than 250 ppm.
Alkalinity [bicarbonate (HCO ₃) and carbonate (CO ₃)]	Results from action of carbon dioxide or acid in water on carbonate rocks such as limestone and dolomite.	In the presence of calcium, carbonates may produce a carbonate scale; they decompose on heating with release of carbon dioxide gas and attendant formation of calcium carbonate scale. Of 76 samples, 2 had an alkalinity content less than 200 ppm; 43 were between 200 and 400 ppm; 27 were between 400 and 600 ppm; and 4 were between 600 and 800 ppm.
Hardness (as CaCO ₃)	Caused by calcium and magnesium which occurs in some amount in almost all rocks but especially in limestone, dolomite, and gypsum.	Before a lather will form, hard water precipitates soap, forming a sludge which causes deposits on bathtubs and is responsible for gray laundry and dingy glassware. Hard water also forms scale in boilers, hot water heaters, and pipes. Of 77 samples analyzed, 11 had a hardness less than 200 ppm; 44 were between 200 and 400 ppm; 17 were between 400 and 600 ppm, and 5 were between 600 and 1000 ppm.
Total dissolved minerals 500 ppm	Includes all mineral ingredients dissolved from rocks and soil.	Mineralization of more than 500 ppm is normally detectable to taste; over 1000 ppm is undesirable for most domestic purposes; livestock may tolerate concentrations up to 7000 ppm. ² Of 77 samples analyzed, 29 had a total mineral content less than 500 ppm; 43 were between 500 and 1000 ppm; and 5 were between 1000 and 1400 ppm.

¹U. S. Public Health Service. 1962. Drinking water standards. Publication No. 956.

²South Dakota State College Agricultural Experiment Station. 1959. Salinity and livestock water quality. Bulletin 481.

heavier. When methane gas is mixed with air in concentrations of 5 to 15 percent, it is highly explosive if ignited. Most dangerous points of concentration are in the well house, within the air cushion of pressure tanks, and in hot water heaters. All such points should be vented to the outside air if methane gas is detected in a water supply. All new wells constructed should be checked for methane gas by the driller before the installation is placed in service. Further, no one should ever enter a large-diameter dug well without previously checking for the presence of methane gas or carbon dioxide, both of which can cause asphyxiation. These gases can be readily removed from water by standard aeration procedures.

Mineral analyses of groundwater from throughout the county are included in appendix A of this report. Wells of comparable depths near these sampling points generally should produce a similar quality of water.

Temperature

Temperature of groundwater varies with the location and depth of the aquifer, the origin and time of occurrence of recharge, and the proximity of the aquifer to bodies of surface water. Where the primary source of recharge is precipitation, the groundwater is fairly cool, because water enters the groundwater reservoir mostly during early spring and late fall, and little temperature variation occurs because the earth materials provide insulation. In Shelby County such aquifers generally contain water ranging from about 54 to 57 degrees Fahrenheit (see appendix A).

Shallow aquifers near the Kaskaskia, or any river, may have a wider temperature fluctuation, and range from about 50 to as high as 68 degrees throughout any given year. Water temperature fluctuations in such an aquifer are primarily controlled by the quantity and range in temperature of surface water entering the groundwater reservoir by floods, artificial recharge, or as a result of heavy pumping. However, changes in water temperature in a well a few hundred feet from a river's edge may lag behind temperature changes in the surface water by as much as several months because of slow water movement through the aquifer.

DEVELOPMENT

Types of Wells

Wells may be classified into types according to the method used in sinking the hole. The most common types of wells in Shelby County are drilled, dug, and augered; however, some driven and bored wells also exist. The type of well chosen for a given location depends on the aquifer and the needs and economic limitations of the user.

Drilled wells with a diameter of 4 to 12 inches are generally used in aquifers occurring from about 40 to as much as 265 feet below the surface. Data on 286 drilled wells are included in appendix B.

Dug or augered wells 2 to 5 feet in diameter are common where water-bearing materials are not highly permeable (cannot transmit much water) and where they are less than about 40 feet below the surface. Most of the 209 large-diameter dug and augered wells inventoried (appendix B) are between 3 and 5 feet in diameter and from 15 to 40 feet deep.

Driven wells, constructed by driving a pointed screen and attached pipe directly into the aquifer, are feasible only where the aquifer is shallow (less than about 30 feet below the surface) and overlain by easily penetrated material such as that in the floodplain of the Kaskaskia River. They usually are 1 to 2 inches in diameter. Only two driven wells are recorded.

Bored wells 6 to 12 inches in diameter were commonly sunk prior to 1930 when home and farm water demands were small. Most of the 75 recorded bored wells range in depth from about 40 to 90 feet. Because the bored well has a small capacity for receiving and storing water from the aquifer, it usually is inadequate for present-day water requirements.

Drilling Methods

Drilled wells, which are most common in Shelby County, may be constructed by the cable-tool or hydraulic-rotary methods. An explanation of these drilling procedures follows.

In the cable-tool method, the earth materials are broken into small fragments by the alternate raising and dropping of a heavy chisel-edged bit, and these fragments are removed from the hole at intervals by a bailer. In an unconsolidated formation, an open hole is maintained by driving the casing as drilling progresses. After the aquifer has been penetrated, the well screen usually is placed opposite the water-bearing formation, the casing pulled upward to expose the screen, and the screen sealed to the casing.

In the conventional hydraulic-rotary method, the drill pipe with a bit attached to the lower end is rotated to break up the material into small particles. A thin mud is pumped through the drill pipe, then out through the openings in the bit, and up to the surface through the space between the drill pipe and the walls of the hole. The circulating mud thus removes the drill cuttings and prevents caving by plastering and supporting the formations penetrated until the final well casing and screen are placed in the hole.

In reverse hydraulic-rotary drilling, the flow of the drilling fluid is reversed from that in the conventional rotary method. The drilling fluid, usually a relatively clear water rather than mud, moves slowly down through the opening between the drill pipe and the bore hole, picks up formation cuttings loosened by the drill bit, enters the drill pipe through the holes in the bit, and by suction pumping moves to the surface where the cuttings settle in a surface pit. The fluid level in the hole must be kept at ground level at all times, because the difference in pressure between the water column in the hole and that in the aquifer prevents caving of the hole until the final well casing with attached screen is installed.

Construction Features

Construction features vary with the type of well and the characteristics of the aquifer to be utilized. Some of the features commonly employed in various types of wells in the county are casing, screening, gravel packing, and grouting (figure 3). These features, along with those for a dug well shown in detail in figure 4, are recommended by the Sanitary Engineering Division of the Illinois Department of Public Health. They are designed to minimize contamination from the surface. Detailed requirements are given in the rules and regulations of the Illinois Water Well Construction Code.

Casing. Wells are cased to maintain an open hole and to assist in protecting the quality of the water supply. Wells penetrating bedrock aquifers are cased opposite the overlying unconsolidated materials (figure 3a) and opposite any bedrock formations subject to caving. Drilled wells tapping water-bearing sand and gravel deposits are cased to the top of the well screen (figure 3b,c). Steel casing is used in drilled wells and some large-diameter dug wells; bored and dug wells may be cased with 6-inch thick reinforced concrete to a minimum depth of 10 feet with the lower portion usually lined with clay or concrete tile or uncemented brick (figures 3d and A). Sanitary protection for bored and dug wells also may be provided by placing a concrete slab at a point at least 10 feet below ground level and by filling in above the slab with compacted earth (figure 4b).

Screening. Most successful drilled wells tapping sand and gravel are equipped with a length of commercially made well screen placed opposite the water-bearing formation (figure 3b,c). A properly selected and installed screen is designed to retain the aquifer material yet permit water to freely enter the well. Torch-cut and hand-sawed slotted casing sometimes is substituted for commercially made well screens; this practice is not recommended because the openings in such a casing are usually too large to retain the aquifer material and too few to allow free water flow into the well. Most wells so equipped have a history of silt or sand pumping, low yield, and short production life.

Wells finished in bedrock aquifers not subject to caving do not require well screens.

Gravel packing. Drilled wells finished in sand and gravel (figure 3c) usually are equipped with a screen that will retain the coarser 30 to 60 percent of the aquifer material immediately adjacent to the well screen; the fine grains are removed from this area by surging, pumping, and bailing. If the aquifer is uniformly fine-grained (figure 3b), an artificial gravel pack envelope at least 6 inches thick may be required around the outside of the screen to prevent migration of fine material into the well. The grain size of this gravel pack should be about five times as large as the average grain size of the water-bearing material.

Some drillers partially fill the well casing with gravel to hold back the aquifer and term this gravel packing. This procedure, however, greatly reduces the yield-capability of the well and is a very poor substitute for a true gravel pack.

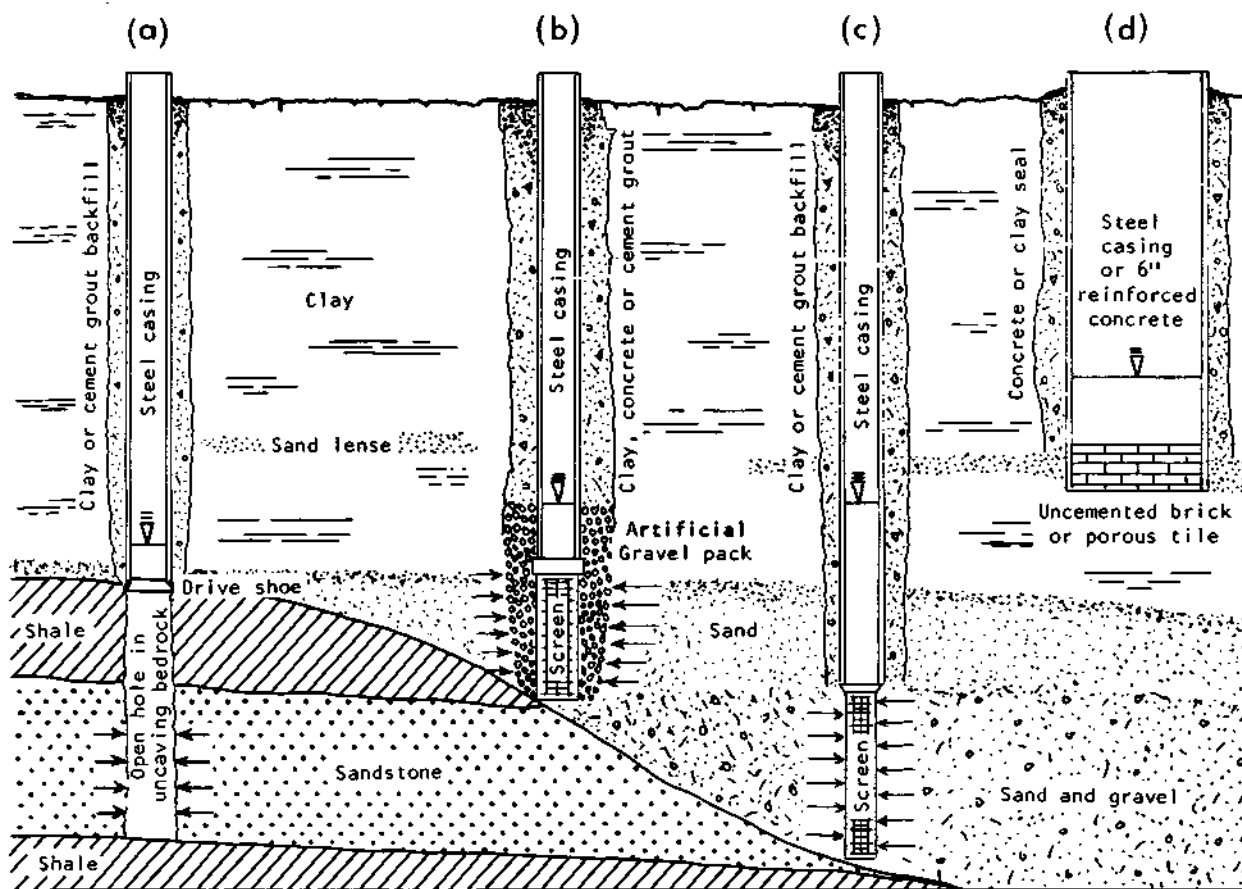
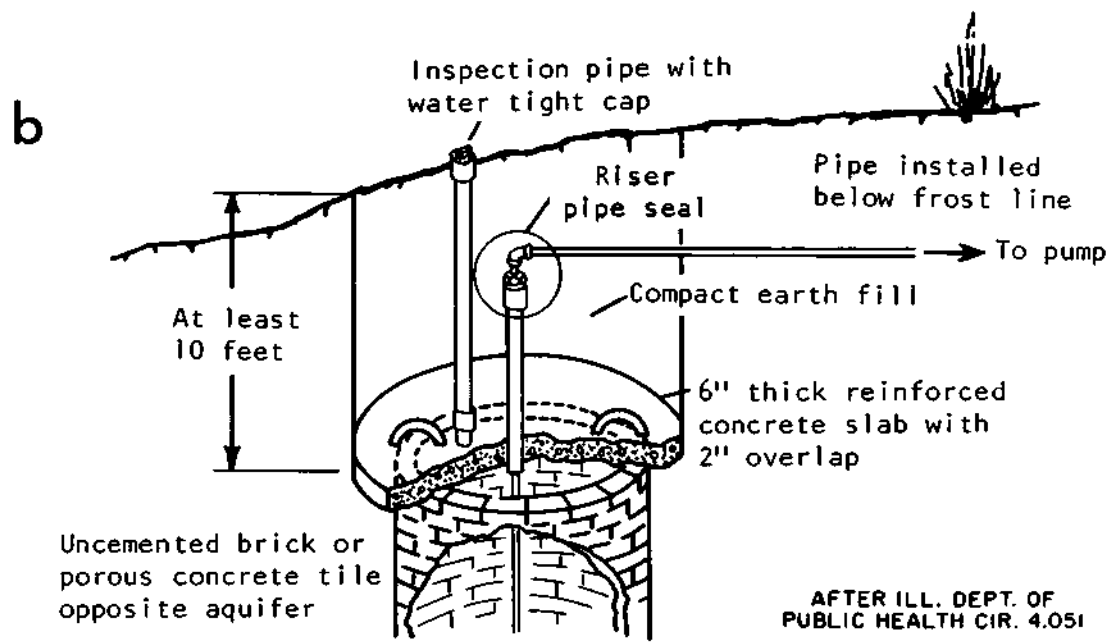
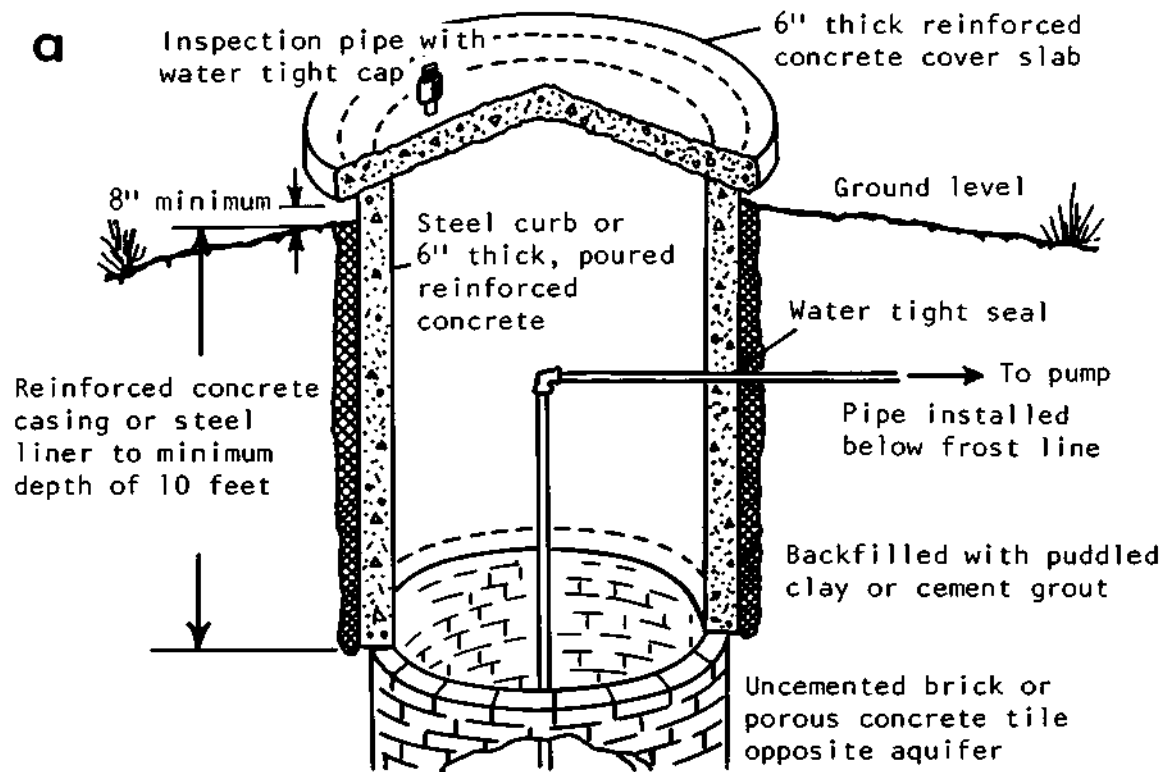


Figure 3. Construction features used in Shelby County wells--(a) drilled well finished in bedrock sandstone; (b) artificial gravel-packed drilled well in fine-grained unconsolidated material; (c) drilled well in coarse-grained unconsolidated material; and (d) shallow large-diameter dug or augered well in relatively impermeable silt, sand, or gravel

Grouting. The annular space between the casing and the bore hole must be sealed to minimize the chance of contamination from the surface. In drilled, dug, or bored wells (figure 3a,c,d), a clay slurry or cement grout must be used to seal the opening between the casing and the bore hole above the aquifer. A cement or concrete grout is required to insure an adequate seal for artificial gravel-pack wells (figure 3b).

Methods of Pumping Water

Most wells in Shelby County are equipped with electrically driven pumps of the suction, jet, cylinder, or turbine types. Suction pumps can be used only where the pumping level is less than about 18 feet. For greater lifts, deep-well jet, cylinder, or vertical or submersible turbine pumps are required. Most of the commercially available submersible and vertical turbine pumps are used in wells with a minimum inside diameter of 4 inches. Vertical turbine pumps are usually installed on large-capacity municipal supply wells in the county.



AFTER ILL. DEPT. OF PUBLIC HEALTH CIR. 4.051

Figure 4. Recommended construction features for large-diameter dug well

Disinfection of Wells

New wells, or old installations after rehabilitation, usually are contaminated and should be disinfected prior to being placed in service. The Illinois Department of Public Health recommends disinfection procedures using a strong chlorine laundry bleach. The correct amount to use can be determined from figure 5, as explained in their instructions which follow.

1) Calculate the amount of water in the well by multiplying the storage capacity per foot (from figure 5a) by the number of feet of water in the well. For example, a 24-inch diameter well with 10 feet of water stores 23 gallons per foot times 10 feet, or 230 gallons. (Considering the well full of water will be satisfactory in most cases since a slight overdose does no harm.)

2) Determine the amount of laundry bleach (from figure 5b) and mix this total amount in about 10 gallons of water. For example, 230 gallons would require 6.9 cups, say 7 cups, of laundry bleach (5.25 percent chlorine).

3) Pour this solution into the well between the casing and the drop pipe. (This may involve raising the pump about 4 inches to allow sufficient space for the addition of the solution and for the placement of a sanitary well seal.)

4) Connect one or more hoses from faucets on the discharge side of the pressure tank to the top of the well and let water from these flow back into the well for at least 15 minutes. Then open each faucet in the system until a chlorine odor or taste appears. Close all faucets. Seal the top of the casing.

5) Let stand for several hours, preferably overnight.

6) Operate the pump, discharging water from all outlets until all chlorine odor and taste disappears.

Chlorine always should be used outside or in well-ventilated places because breathing the fumes is dangerous. In heavy concentrations, chlorine also is harmful to the skin and clothing.

Additional instructions on safe water supplies from wells can be obtained from the Sanitary Engineering Division of the Illinois Department of Public Health, Springfield.

Summary of Major Water Supplies

Municipalities and industries in Shelby County pump about 0.9 million gallons of water per day (mgd) from wells. Groundwater pumped for farm and domestic purposes is estimated to be 1.2 mgd.

The major portion of the municipal pumpage occurs at Shelbyville where more than 500,000 gallons a day is withdrawn from three wells tapping sand and gravel

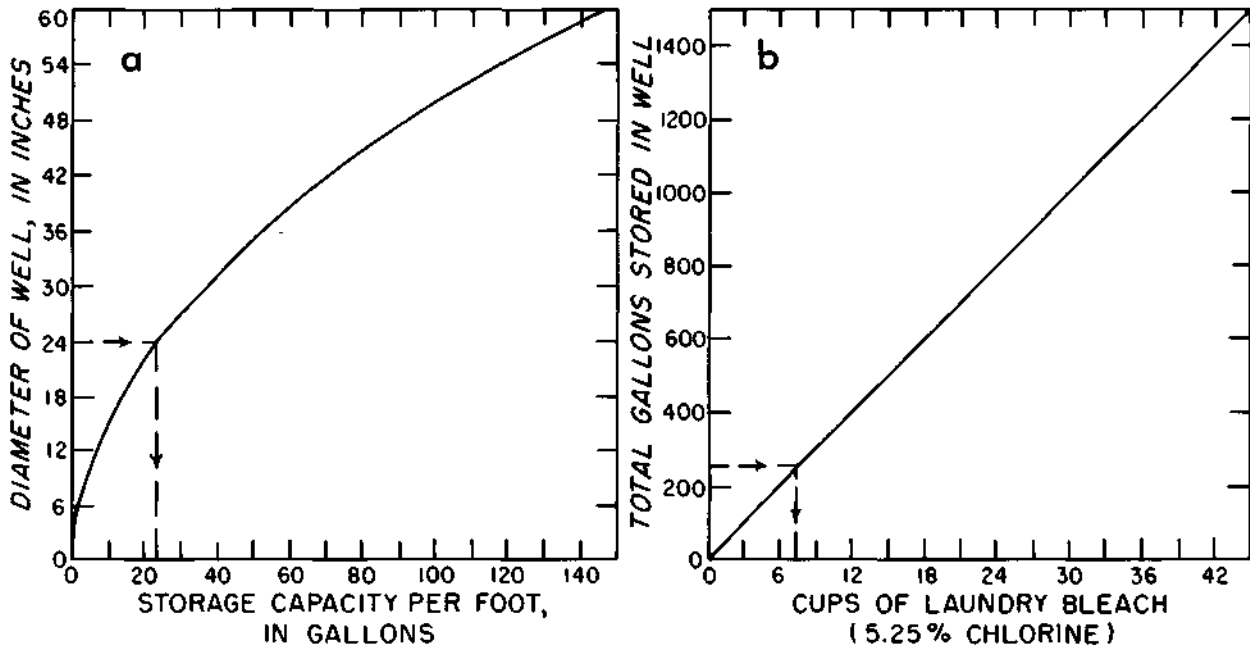


Figure 5. Method for figuring recommended chlorine dosages for well disinfection

deposits in the buried Kaskaskia valley. This buried valley roughly follows and lies near the present Kaskaskia River in Shelby County. Recent studies of the long-term safe yield (the quantity of water that can be withdrawn without exceeding the long-term recharge rate to the formation) of the Shelbyville aquifer indicate that the existing municipal wells can produce about 1.3 mgd on a continuous basis. More than 2 mgd probably can be obtained by constructing additional wells adjacent to the river in areas where that stream and the aquifer are interconnected.

Similar developments also should be possible in other parts of the buried Kaskaskia aquifer where comparable geohydrologic conditions exist (Plate 1). Outside the area underlain by this aquifer, groundwater development for farm and domestic use is largely from scattered and fairly thin layers of sand and gravel, usually within the lower part of the drift deposits.

Wells tapping the bedrock aquifers generally produce less than 10 gpm and usually range from 1 to 4 gpm. The only municipality in the county utilizing the bedrock as a source of supply is Tower Hill. It is estimated that the long-term safe yield of the two-well system tapping the shallow sandstone formations at Tower Hill is 25,000 gallons a day. Similar quantities should be obtainable in the Clarksburg-Stewardson area as shown in Plate 1.

The following is a description of each major municipal groundwater supply in Shelby County. Population figures are taken from the 1960 census; pumping figures are the most recent available and in most cases are for 1966.

COWDEN: The village of Cowden (population 575) uses two wells, located on the bank of the Kaskaskia River, as a source of municipal water supply.

The older well (No. 2) was drilled in 1944 by E. C. Baker and Sons, Sigel, to a depth of 54 feet below ground level. It is 10 inches in diameter and equipped with 11 feet of Johnson Everdur well screen. Upon completion, the well produced 225 gpm for 9 hours with 20 feet of drawdown from a non-pumping water level 18 feet below land surface. The well is equipped with a turbine pump capable of producing 100 gpm. The other well (No. 3), located about 54 feet from Well No. 2, was drilled in 1954 by E. C. Baker and Sons to a depth of 52 feet. It is a 10-inch well equipped with 14 feet of No. 60 slot screen. Upon completion, it produced 141 gpm for 4 hours with a drawdown of 5.3 feet. This well has a 120-gpm turbine pump.

Average daily pumpage is reported to be 46,000 gallons.

Analysis of a sample (appendix A, Lab. No. 136069) showed the water to have a hardness of 340 ppm, total dissolved minerals of 404 ppm, and an iron content of 0.2 ppm.

The water is not treated.

FINDLAY: The village of Findlay (population 759) utilizes one well, located in town, as a source of water supply. It was constructed in 1935 by L. R. Burt, Decatur, to a depth of 154 feet. It is a 26-inch gravel-packed well and has a 12-inch casing to a depth of 129 feet followed by 25 feet of No. 187 slot Cook well screen. When completed, the well produced 150 gpm with a drawdown of 14 feet from a nonpumping water level 96 feet below ground surface. The well was equipped with a 150-gpm turbine pump.

Average daily pumpage is reported to be 40,000 gallons.

Analysis of a sample (appendix A, Lab. No. 115228) showed the water to have a hardness of 263 ppm, total dissolved minerals of 642 ppm, and an iron content of 4.8 ppm. Methane gas is present in the water.

The water is aerated to remove the methane gas and to aid in precipitating the iron; it is filtered to reduce iron, softened to an average of 154 ppm, and chlorinated.

HERRICK: The village of Herrick (population 440) uses one well, located about 3 miles southeast of town, for a municipal water supply. It was drilled in 1964 by E. C. Baker and Sons, Sigel, to a depth of 78 feet. It is a 6-inch well equipped with 13 feet of No. 20 slot and 7 feet of No. 50 slot Cook Red Brass well screen. Upon completion, it produced 180 gpm for 4 hours with a drawdown of 8 feet from a nonpumping water level of 5 feet below land surface. A 60-gpm submersible pump is installed in the well. The long-term safe yield of the well is estimated to be 144,000 gallons per day, or 100 gpm.

Analysis of a sample (appendix A, Lab. No. 163030) showed the water to have a hardness of 318 ppm, total dissolved minerals of 586 ppm, and an iron content of 4.3 ppm.

The water is aerated, filtered, softened, and chlorinated.

MOWEAQUA: The village of Moweaqua (population 1614) now uses 12 wells as a source of water supply. All wells are located in a 2-acre plot about 2 miles north of town in Macon County.

Four of the wells were drilled in 1947 by Cyrus Stevens, Findlay, and range in depth from 23 to 26 feet. They are 2.5 inches in diameter and equipped with 7 feet of slotted screen. Two 3-inch wells constructed in 1952 by Stevens are 28 feet deep and equipped with 7 feet of well screen. These six wells are pumped by a 120-gpm centrifugal pump through a common suction header.

Five additional wells were drilled in 1961 by Stevens and range in depth from 25 to 27 feet. They are 4 inches in diameter and are equipped with 5 feet of well screen. A 150-gpm centrifugal pump is used to pump these five wells through a common suction header.

One well, now on standby, was constructed in 1953 by Luther Burt, Decatur, to a depth of 33.5 feet. It is an 8-inch well equipped with 9 feet of No. 30 slot well screen. A 60-gpm submersible pump is installed in the well.

Average daily pumpage is estimated to be 110,000 gallons.

Analysis of a sample showed the water to have a hardness of 440 ppm, total dissolved minerals of 494 ppm, and an iron content of 1.1 ppm.

The water is aerated and filtered to reduce iron, and is chlorinated.

SHELBYVILLE: Three wells, located about 3 miles southwest of town, furnish the water supply for Shelbyville (population 4821).

The wells, spaced 300 feet apart, were drilled in 1955 by Layne-Western Company, Kirkwood, Missouri, and have finished depths of 54 to 60 feet. All are gravel packed and have a 12-inch casing and 15 feet of Layne stainless steel screen. The gravel-pack envelopes on these wells range from 11 to 13 inches thick. Production tests were conducted on the completed wells as indicated below:

	<u>No. 1</u>	<u>No. 2</u>	<u>No. 3</u>
Length of test, hours	6.5	12	24
Pumping rate, gpm	328	545	500
Nonpumping water level, ft	22.2	18.5	15.5
Drawdown, ft	6.5	11.2	10.3

Although all wells are equipped with 500-gpm turbine pumps, the practical sustained yield of the existing well field is estimated to be 900 gpm or 1.3 mgd. A total of 2.0 mgd can probably be developed if additional, widely spaced wells are constructed farther to the south nearer the Kaskaskia River.

Pumpage is reported to average 500,000 gallons a day.

Analyses of samples (see appendix A) showed the water to have a range in hardness of 275 to 475 ppm, total dissolved minerals of 400 to 500 ppm, and an iron content of 0.1 ppm.

The water is chlorinated.

SIGEL: The town of Sigel (population 387) uses one well, located near the center of town, as a source of water supply. It was completed in 1954 by Holkenbrink Drilling Company, Effingham, to a depth of 65 feet. It is 8 inches in diameter and equipped with 10 feet of No. 20 slot well screen. During a production test the well produced 31 gpm for 21 hours with a drawdown of 36.5 feet from a nonpumping water level of 15.5 feet below the top of the casing. A submersible pump is installed in the well.

Average daily pumpage is estimated to be 8500 gallons.

Analysis of a sample (appendix A, Lab. No. 144606) showed the water to have a hardness of 217 ppm, total dissolved minerals of 372 ppm, and an iron content of 2.7 ppm.

The water is aerated, settled, chlorinated, and filtered to remove iron; and is softened to an average of 89 ppm.

STEWARDSON: The village of Stewardson (population 656) now uses one well, in the southeast part of town, as a source of water supply. It was drilled in 1955 by E. C. Baker and Sons, Sigel, to a depth of 50 feet. It has a 10-inch outer casing to a depth of 40 feet and an 8-inch inner casing to a depth of 50 feet. The 10 feet of exposed 8-inch pipe was perforated with 1/8-inch slots. Upon completion, the well produced 150 gpm for 6 hours with a drawdown of 5 feet from a nonpumping water level 9.4 feet below the top of the casing. It is equipped with a 100-gpm turbine pump.

Average daily pumpage is estimated to be 27,000 gallons.

Analysis of a sample (appendix A, Lab. No. 137786) showed the water to have a hardness of 288 ppm, total dissolved minerals of 371 ppm, and an iron content of 1.0 ppm.

The water is aerated, settled, and filtered to remove iron.

STRASBURG: The village of Strasburg (population 467) uses one well, located about 4.5 miles southeast of town, for a municipal water supply. It was constructed

in 1964 by E. C. Baker and Sons, Sigel, to a depth of 37 feet. It is a 6-inch well with 6.5 feet of No. 80 slot Johnson Red Brass well screen. During a 3-hour production test, the well produced 72 gpm with a drawdown of 10 feet from a nonpumping water level 4 feet below the land surface. The permanent pump is a 50-gpm submersible.

Average daily pumpage is estimated to be 25,000 gallons.

Analysis of a sample (appendix A, Lab. No. 162782) showed the water to have a hardness of 348 ppm, total dissolved minerals of 489 ppm, and an iron content of 2.6 ppm.

The water is aerated and filtered to remove iron, softened to an average of 176 ppm, and chlorinated.

TOWER HILL: The village of Tower Hill (population 700) now uses two wells, located about 1.25 miles east of town, as a source of water supply. These are the only municipal wells in Shelby County that tap water-bearing sandstone formations in the bedrock.

The older well was drilled in 1950 by E. C. Baker and Sons, Sigel, to a depth of 50 feet below land surface. It is 24 inches in diameter to a depth of 17 feet and 10 inches in diameter from 17 to 50 feet. Upon completion, the well was pumped at a rate of 50 gpm for 14 hours with a drawdown of 15 feet from a nonpumping water level of 6 feet below ground level. The well is equipped with a 15-gpm submersible pump.

The other well, drilled in 1950 by E. C. Baker and Sons, is about 600 feet north of the first well. It is an 8-inch well 59 feet deep. During a 6-hour production test, the well was pumped at rates from 8.5 to 21.5 gpm with a final drawdown of 27.5 feet from a nonpumping water level 8 feet below ground surface. A 10-gpm submersible pump is installed in the well. The safe yield of the existing well field is estimated to be 25,000 gallons per day, or 17 gpm.

Average daily pumpage is estimated to be 20,000 gallons.

Analysis of a sample (appendix A, Lab. No. 123680) showed the water to have a hardness of 259 ppm, total dissolved minerals of 337 ppm, and an iron content of 2.6 ppm.

The water is not treated.

WINDSOR: The city of Windsor (population 1021) uses three wells as a source of water supply.

Two of the wells are located about 1.5 miles south of town and were drilled 120 feet apart in 1951 and 1952 by E. C. Baker and Sons, Sigel, to depths of 99 and 95 feet. They are 7.5-inch wells and have 1/8-inch slots

cut in the lower 6 feet of casing. Submersible pumps of 50 and 60 gpm capacity are installed in the wells.

One standby well, located in town, was drilled in 1949 by Hayes and Sims, Champaign, to a depth of 131 feet. It is a 12-inch well equipped with 11 feet of No. 25 slot well screen. Upon completion, the well produced 33.5 gpm for 7 hours with a drawdown of 87 feet from a nonpumping water level 28.5 feet below ground surface. A 20-gpm submersible pump is installed in the well.

Average daily pumpage is estimated to be 78,000 gallons.

Analysis of a sample (appendix A, Lab. No. 115143) showed the water to have a hardness of 331 ppm, total dissolved minerals of 573 ppm, and an iron content of 5.4 ppm. Methane gas is also present in the water.

The water is aerated, filtered, and chlorinated.

APPENDIX A - CHEMICAL QUALITY OF GROUNDWATER

Tabulated data of mineral content for groundwater supplies in Shelby County follow.

Symbols used in the tabulations are:

- D - glacial drift
- BR - bedrock
- * - State Department of Public Health chemical analyses

The sources and significance of the major dissolved elements and substances in groundwater, their ranges in waters of Shelby County, and U. S. Public Health Service drinking water limits (1962) are included in table 1.

Appendix A. Chemical quality of groundwater

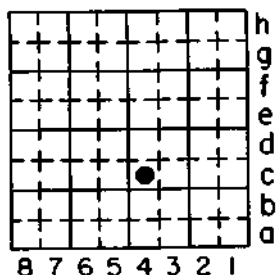
Well number	Owner	Depth	Source	Laboratory number	Iron Fe	Manganese Mn	Ammonium NH ₄	Sodium Na	Calcium Ca	Magnesium Mg	Silica SiO ₂	Fluoride F	Nitrate NO ₃	Chloride Cl	Sulfate SO ₄	Alkalinity (as CaCO ₃)	Hardness	Total dissolved minerals	Temperature °F
T9N, R2E																			
17.5d	Herrick (V)	30	D	155044	8.7	.4	--	--	--	--	--	.4	5.3	10	--	244	236	312	--
T9N, R3E																			
10.1g1	Cowden (V)	51	D	91357	.2	.0	.1	64	83.1	8.4	15	--	1.2	11	81	282	242	437	--
10.1g2	Cowden (V)	53	D	115229	.1	Tr	Tr	3	87.3	34.5	20	.1	9.5	7	62	284	361	420	--
10.1g3	Cowden (V)	52	D	136069	.2	.1	Tr	14	83.5	31.9	16	.4	10.1	7	59	292	340	404	--
15.3h	James Beaumont	45	D	136352	2.8	--	--	--	--	--	--	.1	.6	59	--	280	280	421	56
T9N, R5E																			
4.6a	Henry Vonderheid	184	BR	169753	.5	--	--	--	--	--	--	--	--	64	--	456	20	609	--
6.6h	Everett Kessler	160	BR	169750	.1	--	--	--	--	--	--	--	--	208	--	520	10	948	--
11.2d	Mound School	220	BR	169756	6.5	--	--	--	--	--	--	--	--	7	--	344	112	378	--
T9N, R6E																			
14.2f4	Sigel (T)	65	D	144606	2.7	Tr	.3	59	49.5	22.6	17	.5	.4	17	5	316	217	372	56
15.5a	Francis Stodden	267	BR	169751	1.0	--	--	--	--	--	--	--	--	3150	--	880	84	6395	--
16.1g2	Bernard Schumacher	43	D	149862	7.2	--	--	--	--	--	--	--	--	10	--	320	475	590	--
T10N, R4E																			
1.2f	Henry Biedert	24	D	81112	.0	.0	.1	47	89.2	36.0	10	--	119.6	56	53	242	371	560	--
4.8g	Government Land Bank	50	D	80885	.3	.0	.1	43	80.4	30.7	12	--	4.9	22	40	344	328	448	--
7.5a	Lewis, Heirs	48	D	80833	1.6	--	Tr	41	117.0	47.8	12	--	4.0	54	101	394	489	642	--
8.7a	Ruby Strohl	40	D	80832	.4	--	Tr	30	120.0	40.1	12	--	49.7	35	155	280	465	630	--
10.5g	Frank Wheat	34	BR	81054	Tr	Tr	Tr	54	90.0	25.9	9	--	19.5	53	105	250	332	546	--
17.8e	Ollie Reid	89	BR	169754	5.4	--	--	--	--	--	--	--	--	15	--	408	568	756	--
20.4d	T. C. Dove	31	D	81130	.0	.0	Tr	95	58.2	21.5	10	--	38.9	37	55	300	234	492	--
T10N, R5E																			
1.4a	Orville Bauer	221	BR	130618	1.0	--	--	--	--	--	--	--	--	2450	--	800	40	5028	--
3.4h	Edgar Schlecte	18	D	157055	.4	--	--	--	--	--	--	--	32.0	30	--	196	396	598	--
22.4h	John Falk	138	BR	163802	.2	--	--	--	--	--	--	--	--	20	--	544	24	615	--
23.8g	Stewardson-Strasburg Comm. Schools	170	BR	165386	--	--	--	--	--	--	--	.6	2.3	256	--	456	12	969	--
23.8h	Stewardson-Strasburg Comm. Schools	170	BR	165385	1.5	.06	--	--	--	--	--	.6	2.4	240	--	448	12	943	--
34.5c	Stewardson (V)	20	D	37049	.6	--	--	--	--	196.0	--	--	110.0	93	193	256	432	885	--
35.7a2	Stewardson (V)	50	D	137786	1.0	--	--	--	--	--	--	.3	.1	5	--	316	288	371	56
T10N, R6E																			
5.7c	R. Ensign	140	BR	90716	.5	--	--	--	--	--	--	--	--	547	17	132	192	1192	--
15.4b	Orville Wassom	35	D	150417	1.5	--	--	--	--	--	--	--	11.8	20	--	320	390	616	--
18.8d7	Strasburg (V)	37	D	162782	2.6	.27	--	--	--	--	--	.3	.9	7	--	324	348	489	56
T11N, R2E																			
15.5a	Tower Hill (V)	30	D	150927	6.5	.3	--	--	--	--	--	.5	.9	7	--	208	212	311	--
22	L. C. Cannon	30	D	85961	.1	--	--	--	--	--	--	--	--	56	--	--	568	835	--
22	Mrs. Ida Foltz	14	D	91718	.3	--	--	--	--	--	--	--	--	58	--	412	668	1060	--
22.1f	Tower Hill (V)	35	D	150926	8.7	.3	--	--	--	--	--	.3	3.3	25	--	244	420	529	--
23.2g3	Tower Hill (V)	50	BR	123680	2.6	--	--	--	--	--	--	--	--	9	--	272	259	337	56
23.2h3	Tower Hill (V)	59	BR	123246	3.6	--	--	--	--	--	--	.3	--	9	--	312	247	335	55
23.7g2	Tower Hill (V)	77	BR	144607	2.9	.1	.6	44	63.0	16.4	24	.3	.0	24	2	284	224	376	--
23.8e5	Tower Hill (V)	36	D	121992	1.5	--	--	--	--	--	--	--	--	15	3	272	251	314	55
24	Tower Hill (V)	37	D	122759	10.8	--	--	--	--	--	--	.3	.0	5	--	272	262	321	--
T11N, R3E																			
1.1d	J. I. Reed	72	D	165926	8.9	--	10.7	--	--	--	--	--	--	33	--	648	400	722	--
14.4a	T. C. Dove	72	D	140406	2.4	--	--	92	--	--	--	--	--	61	26	440	352	563	--
21.1d	R. A. Creswell	94	BR	169749	4.6	--	--	--	--	--	--	--	--	1010	--	464	252	2276	--
23.2h1	Henry Hortenstine	69	D	140407	.9	--	--	3	--	--	--	--	--	4	59	264	324	334	--
23.2h2	Shelbyville (C)	51	D	37529 ^a	.5	.15	--	33	93.0	37	--	--	6.2	31	73	332	384	470	--
25.7e	Dr. C. A. Spears	51	D	152857	.1	.0	--	--	--	--	--	--	5.7	4	--	260	286	316	--
26.3h	John & Owen Shull	55	D	135710	Tr	--	--	--	--	--	--	--	--	15	--	340	428	513	--
26.4h1	Shelbyville (C)	52	D	136376	.2	.3	--	--	--	--	--	.1	5.7	33	--	360	448	504	54
26.4h3	Shelbyville (C)	60	D	41274 ^a	.1	.2	0	5	66	29	8	.0	1.4	5	35	250	284	390	--
26.4h5	Shelbyville (C)	54	D	137848	.1	.4	.5	0	117.6	43.7	17	.1	2.2	24	87	348	474	508	55

Appendix A (Continued)

Well number	Owner	Depth	Source	Laboratory number	Iron		Manganese	Ammonium	Sodium	Calcium	Magnesium	Silica	Fluoride	Nitrate	Chloride	Sulfate	Alkalinity	Hardness	Total dissolved minerals	Temperature
					Fe	Mn	NH ₄	Na	Ca	Mg	SiO ₂	F	NO ₃	Cl	SO ₄	(as CaCO ₃)			°F	
T11N, R4E																				
--	Isaac Neut	23	D	39029	1.6	.40	--	--	28	78.4	29.5	--	--	2.6	26	86	280	341	460	--
8.6a	Shelbyville (C)	26	D	134307	1.4	1.00	--	--	--	--	--	--	--	--	48	296	364	364	856	--
8.6a	Shelbyville (C)	26	D	134514	1.0	1.70	--	--	--	--	--	--	--	--	56	338	380	332	976	--
22.1g1	John Switzer	57	D	111235	11.9	--	--	--	--	--	--	--	--	--	33	--	800	523	925	--
22.1g2	John Switzer	72	D	81111	4.4	.0	25.3	101	119.6	40.7	10	--	2.0	27	0	716	467	697	--	
T11N, R5E																				
3.2b	T. D. Hennigh	105	D	170018	1.8	--	7.9	--	--	--	--	--	--	--	8	--	464	136	529	59
12.7h1	Windsor (C)	86	D	75696	4.0	.0	8.2	109	66.2	34.2	10	--	.9	4	0	560	306	569	--	
12.7h3	Windsor (C)	101	D	115143	5.4	.0	13.4	90	71.7	36.7	29	.2	.2	3	0	560	331	573	56	
12.7h4	Windsor (C)	99	D	145470	6.0	.0	8.2	35	73	39.5	18	.7	5.5	5	0	432	344	446	56	
12.7h5	Windsor (C)	98	D	128973	5.7	.0	15.4	79	71.4	36.6	27	.4	1.2	3	2	536	329	556	57	
13.1g	J. Turner	265	BR	103064	.9	--	--	--	--	--	--	--	--	2350	--	--	678	89	5222	55
15.4a	Clark Schmidt	68	D	156933	1.1	--	11.8	--	--	--	--	--	28.6	60	--	--	548	736	1028	--
22.6d	Roy F. Schmidt	26	D	156932	10.0	--	21.9	--	--	--	--	--	9.1	11	--	--	592	502	604	--
T11N, R6E																				
20.7h	Lloyd Elson	259	BR	169752	2.9	--	--	--	--	--	--	--	--	3950	--	--	620	84	7508	--
T12N, R2E																				
1.4b	Dudley Smith	190	BR	169748	.4	--	2.2	--	--	--	--	--	.4	500	--	--	456	200	1309	--
18.6b21	Assumption (C)	24	D	115314	.3	.3	Tr	9	91	29.4	19	.3	.3	7	125	228	349	442	54	
18.6d1	Assumption (C)	28	D	141834	9.9	--	--	--	--	--	--	.2	.2	8	--	232	216	276	57	
T12N, R3E																				
16.1a	Westervelt H.S.	75	D	90574	8.7	--	--	--	--	--	--	--	--	67	--	--	448	399	648	--
T12N, R4E																				
3.8g1	Findlay (V)	167	D	75739	2.8	.0	8.5	167	51.6	26.0	12	--	1.3	86	8	512	236	659	--	
3.8g2	Findlay (V)	171	D	75860	4.0	.0	9.5	179	46.4	26.2	12	--	1.3	97	12	490	224	705	--	
3.8g3	Findlay (V)	154	D	115228	4.8	.0	11.3	148	58.9	27.9	23	.3	.4	82	0	500	263	642	58	
33.2b	Martha James	165	D	72585	1.2	0	.1	82	55.6	25.7	8	--	.4	34	44	328	245	455	--	
T12N, R5E																				
6.1e	Leroy Hugo	200	BR	137257	1.6	--	--	--	--	--	--	--	--	5700	--	--	320	730	9858	--
18.7b	Boy Scouts of America	137	BR	144060	.9	--	--	--	--	--	--	--	--	1700	--	--	524	96	3394	--
36	Windsor Comm. H.S.	100	D	78851	.6	.0	2.8	117	31.3	16.6	10	--	1.2	17	0	384	147	469	--	
36.5b3	Windsor (C)	134	D	75526	3.9	.1	3.9	148	49.6	21.0	35	--	1.6	15	0	520	211	590	--	
36.5c2	Windsor (C)	132	D	117362	4.1	.0	5.1	141	53.9	21.6	26	.3	.1	21	3	512	224	591	59	
36.6b	Windsor (C)	85	D	75175	25.0	.0	5.2	78	66.8	29.8	15	--	1.9	0	5	468	290	472	--	
T13N, R3E																				
1.1g	George Reuss	50	D	81030	1.0	.0	3.1	100	66.5	30.1	12	--	1.4	40	3	456	290	543	--	
12.1b	Henry Atkinson	95	D	81029	1.2	.0	2.0	126	36.2	10.2	12	--	1.7	22	2	376	132	426	--	
16.1e	Dr. C. D. Casey	224	BR	103710	2.0	--	--	--	--	--	--	--	--	83	--	468	125	638	--	
28.1a	J. R. Ward	241	BR	134067	3.8	--	--	--	--	--	--	--	--	69	--	484	192	746	--	
35.5h2	Roy Macklin	232	BR	134066	4.0	--	--	--	--	--	--	--	--	2650	--	296	344	5340	--	
T13N, R4E																				
16.2g	Raymond Robinson	265	BR	169755	2.3	--	--	--	--	--	--	--	--	1600	--	--	800	40	3485	--
19.1g	Birkett L. Williams	11	D	141833	1.8	.3	--	--	--	--	--	--	--	38	--	396	524	600	--	
T14N, R2E																				
20.7h	Nerjorie Porter	29	D	168986	Tr	--	--	--	--	--	--	--	95.8	58	--	--	280	660	913	--
T14N, R3E																				
21.3e	M. A. Sanner & B. S. Clark	26	D	80800	.1	.0	Tr	29	144.0	78.3	10	--	173.0	93	216	248	682	937	--	
27.1d1	Naomi Coultas	99	D	80798	3.0	.0	3.4	126	88.0	46.1	14	--	3.4	59	13	594	410	709	--	
32.6h	George Elliott	100	D	143748	39.0	--	--	--	--	--	--	--	--	100	--	428	976	1365	--	
35.1e	D. E. Baird	65	D	80799	4.0	Tr	11.3	132	126.0	50.8	14	--	2.7	77	6	726	524	882	--	

APPENDIX B - RECORDS OF WELLS

The well-numbering system used in this report is based on the location of the well, and uses the township, range, and section for identification. The well number consists of five parts: county abbreviation (SHL), township (T), range (R), section, and coordinate within the section. Sections are divided into rows of 1/8-mile squares. Each 1/8-mile square contains 10 acres and corresponds to a quarter of a quarter of a quarter section. A normal section of 1 square mile contains 8 rows of 1/8-mile squares; an odd-sized section contains more or fewer rows. Rows are numbered from east to west and lettered from south to north as shown in the diagram.



Shelby County
T11N, R4E
Section 23

The number of the well shown is SHL 11N4E-23.4c. Where there is more than one well in a 10-acre square they are identified by arabic numbers after the lower case letter in the well number.

In the listing of wells owned by municipalities, the place-name is followed by V, T, or C in parentheses to indicate whether it is a village, town, or city, except where the word City is part of the place-name.

Owners are listed according to the most current information available - the 1964 plat book and recent well records for Shelby County.

Symbols and abbreviations shown indicate the following:

- - constructed before year given, exact date unknown
- * - test hole not developed as well
- ** - abandoned well
- drv - driven well
- drl - drilled well
- bor - bored well
- drl-GP - drilled well, gravel packed

The types of wells and methods of construction used in Shelby County, their susceptibility to surface contamination, and methods of disinfection are discussed in the text of this report.

Appendix B (Continued)

Well number	Owner	Well				Screen		Land surface elev (ft above msl)	Non-pumping water level (ft)	Draw-down (ft)	Pumping rate (gpm)	Observed specific capacity (gpm/ft)	Length of test (hr)	Water-bearing formation and depth (ft)	Driller	
		Year constructed	Type	Depth (ft)	Diam (in)	Length (ft)	Diam (ft)									
T9N, R5E																
3.1g1	Stewardson (V)	1955	drl	50	--	--	--	632	--	--	--	--	--	*Mud sand, 9-14; sand & gravel, 17-47	Baker	
3.1g2	Stewardson (V)	1955	drl	42	--	--	--	632	--	--	--	--	--	*Mud sand, 10-15; mud sand & gravel, 17-35	Baker	
3.1g3	Stewardson (V)	1955	drl	40	--	--	--	632	--	--	--	--	--	*Sand & gravel, 10-40	Baker	
3.1h1	Stewardson (V)	1955	drl	32	--	--	--	630	--	--	--	--	--	*Sand, 15-25	Baker	
3.1h2	Stewardson (V)	1955	drl	37	--	--	--	630	--	--	--	--	--	*Sand & gravel, 17-37	Baker	
3.1h3	Stewardson (V)	1955	drl	47	--	--	--	630	--	--	--	--	--	*Mud sand, 9-14; sand & gravel, 17-42	Baker	
3.1h4	Stewardson (V)	1955	drl	39	--	--	--	630	--	--	--	--	--	*Sand & gravel, 10-39	Baker	
3.2h1	Stewardson (V)	1955	drl	47	--	--	--	640	--	--	--	--	--	*Sand & gravel, 19-46	Baker	
3.2h2	Stewardson (V)	1955	drl	46	--	--	--	640	--	--	--	--	--	*Sand & gravel, 10-45	Baker	
3.3g	Stewardson (V)	1955	drl	50	--	--	--	635	--	--	--	--	--	*Sand & gravel, 14-48	Baker	
3.3h1	Stewardson (V)	1955	drl	37	--	--	--	640	--	--	--	--	--	*Sand, 17-48	Baker	
3.3h2	Stewardson (V)	1955	drl	17	--	--	--	640	--	--	--	--	--	*	Baker	
4.6a	Henry Vonderheid	1964	drl	184	6	22	6	640	28	--	13	--	--	Sandstone, 162-184	Cummins	
6.1f	Oscar Kessler	1961	drl	148	4	--	--	636	30	60	4	0.1	6	Sandstone, 110-148	Fleming	
6.6h	Everett Kessler	1961	drl	160	4	--	--	628	20	70	2.5	--	7	Sandstone, 126-147, 195-160	Fleming	
11.2d	Mound School	1932	drl	220	--	--	--	650	--	--	--	--	--	Sandstone, 200-220	Baker	
12.8a	Paul Wittenberg	1958	drl	78	4	--	--	652	18	--	2	--	--	Sand, 30-37	Holkenbrink	
T9N, R6E																
1.4d	Norbert Hoene	1952	drl	67	6	--	--	645	3.8	51.2	5	0.1	2	Crack in shale, at 60'	Baker	
1.6a	Ambrose Probst	1962	drl	67	--	--	--	640	--	--	--	--	--	*	Baker	
3.4f1	Alphonse Hoene	1955	dug & drl	80	35-5	--	--	625	7	--	1.5	--	--	Lime, 60-80	Holkenbrink	
3.4f2	Alphonse Hoene	1955	drl	40	--	--	--	625	--	--	0.5	--	--	*	Holkenbrink	
5.6f	Clark Bolling	1955	drl	52	--	--	--	560	--	--	--	--	--	*	Baker	
5.8e	Bessie Bolling	1959	drl	42	6	4	6	.080 580	28	8	1	0.1	5	Sand & gravel, 38-42	Baker	
9.1h	Albert Hoene	1964	drl	100	7	70	7	.030 620	15	--	2	--	0.5	Sand & gravel, 15-35; sandstone, 55-100	Woodward	
9.8e	Charles Ludwig	1959	drl	60	6	--	--	645	20	26	3	0.1	6	Sand & gravel, 36-48; sandy lime, 48-60	Baker	
9.8h1	Eley's Tavern	1935	drl	60	6	4	6	.040 640	15	21	15	0.7	1	Sand & gravel, 48-60	Baker	
9.8h2	Sinclair Station	1935	drl	58	5	--	--	640	14	16	15	0.9	--	Sand, 54-58	Baker	
10.3g1	Harry Schutte	1955	drl	80	--	--	--	640	--	--	--	--	--	*Mud sand, 32-37, 57-62, 67-72	Baker	
10.3g2	Harry Schutte	1955	drl	42	--	--	--	640	--	--	--	--	--	*	Baker	
10.3g3	Harry Schutte	1956	drl	52	--	--	--	640	--	--	--	--	--	*	Baker	
12.2a	Lydia Baker	1952	dug & drl	59	7-6	--	--	641	12.5	1.5	3	2.0	6	Sand & gravel, 52-59	Baker	
12.4h1	Ambrose Probst	1962	drl	62	--	--	--	640	--	--	--	--	--	*Dirty sand, 46-47	Baker	
12.4h2	Ambrose Probst	1962	drl	82	--	--	--	640	--	--	--	--	--	*	Baker	
12.5h1	Ambrose Probst	1962	drl	67	--	--	--	640	--	--	--	--	--	*	Baker	
12.5h2	Ambrose Probst	1962	drl	30	--	--	--	640	--	--	--	--	--	*	Baker	
12.5h3	Ambrose Probst	1962	drl	97	--	--	--	640	--	--	--	--	--	*	Baker	
12.6h	Ambrose Probst	1962	drl	72	--	--	--	640	--	--	--	--	--	*	Baker	
14.1g	St. Michael School	1953	drl	230	--	--	--	640	--	--	--	--	--	*	Baker	
14.1h	Ben Renschen	-1915	dug	25	--	--	--	640	--	--	--	--	--	--	--	
14.2c	Victor Ezerwonka	-1921	dug	15	--	--	--	635	--	--	--	--	--	--	--	
14.2e	Sigel (T)	1954	drl	68	--	--	--	635	--	--	--	--	--	--	--	
14.2f1	Sigel (T)	1954	drl	66	--	--	--	640	--	--	--	--	--	--	*Sand, 29-33, 46-67	Holkenbrink
14.2f2	Sigel (T)	1954	drl	67	--	--	--	640	--	--	--	--	--	--	*Sand, 46-65	Holkenbrink
14.2f3	Sigel (T)	1954	drl	67	--	--	--	640	--	--	--	--	--	--	*Sand, 49-67	Holkenbrink
14.2f4	Sigel (T)	1954	drl	67	--	--	--	640	--	--	--	--	--	--	*Sand, 43-45; sand & gravel, 53-67	Holkenbrink
14.2f4	Sigel (T)	1954	drl	65	8	10	8	.020 640	14	36.5	31	0.9	21	Sand, 48-65	Holkenbrink	
14.2g	S. S. Bigler	-1912	dug	25	--	--	--	635	--	--	--	--	--	--	Sand & gravel, at 25'	--
14.3a	Carl Allhoff	1949	drl	57	6	5	6	.035 635	4.1	20.9	3	0.1	--	Sand & gravel, 50-57	Baker	
14.3f	Sigel (T)	1954	drl	76	--	--	--	635	--	--	--	--	--	--	*Sand, 47-76	Holkenbrink
14.3g1	Dr. H. H. C. Hanck	-1915	dug	22	--	--	--	635	--	--	--	--	--	--	--	
14.3g2	Charles Ludwig	1939	drl	57	--	--	--	635	--	--	--	--	--	--	*	Baker
14.4f	Rev. G. E. Faller	1929	drl	121	5	--	--	630	--	--	--	--	--	--	--	
14.5h1	Leonard Siemer	-1909	dug	35	--	--	--	630	--	--	--	--	--	--	--	
14.5h2	Leonard Siemer	1935	drl	58	6	--	--	630	14	16	20	1.3	--	Sand & gravel, at 58'	Baker	
15.3e	Henry Berchtold	-1915	dug	25	--	--	--	625	--	--	--	--	--	--	--	
15.5a	Francis Stodden	1954	drl	267	6-5	--	--	620	37	--	4	--	--	Sandstone, 165-188, 221-237; sandstone & shale, 237-267	Holkenbrink	
16.1g1	B. J. Schumacher	-1959	--	22	--	--	--	630	--	--	--	--	--	--	--	
16.1g2	B. J. Schumacher	1959	drl	43	6	5	1.25	.025 630	--	--	2.5	--	--	Sand & gravel, 39-43	Baker	
T10N, R1E																
7.8a	Everett Kuhn	--	dug	18	48	--	--	666	3.2	--	--	--	--	--	--	
9.5a	Wayne Smith	--	dug	18	24	--	--	680	5	--	--	--	--	--	--	
12.8a	Kuhn Bros. Garage	-1928	spring	--	--	--	--	480	--	--	--	--	--	--	Limestone	
14.2h	Ray M. Tuetken	1966	drl	161	--	--	--	660	--	--	--	--	--	--	*Mud sand, 141-149; lime, 157-161	Baker
29.3f	Telephone Co.	--	bor	16	18	--	--	675	--	--	--	--	--	--	--	
29.4d	Mrs. Bertha Morgan	-1918	dug	25	--	--	--	675	--	--	--	--	--	--	--	
30.7h	Margaret Gudehus	--	dug	--	36	--	--	660	3.9	--	--	--	--	--	--	
32.4h	Herman Manniken	-1867	dug	14	42	--	--	672	3.6	--	--	--	--	--	--	
T10N, R2E																
7.6c	Frances Mathewson	--	drl	96	--	--	--	645	--	--	--	--	--	--	*Sand, 34-36; sand & gravel, at 96'	Warren

Appendix B (Continued)

Well number	Owner	Well				Screen			Land surface elev (ft above msl)	Non-pumping water level (ft)	Draw-down (ft)	Pumping rate (gpm)	Observed specific capacity (gpm/ft)	Length of test (hr)	Water-bearing formation and depth (ft)	Driller
		Year constructed	Type	Depth (ft)	Diam (in)	Length (ft)	Diam (ft)	Slot size (in)								
T11N, R3E (Continued)																
26.4h3	Shelbyville (C)	1955	drl-GP	60	26-12	15	12	--	560	22.2	6.5	328	50.5	6.5	Sand & gravel	Layne-Western
26.4h4	Shelbyville (C)	1955	drl-GP	58	26-12	15	12	--	560	18.5	11.2	545	48.7	12	Sand & gravel, 19-58	Layne-Western
26.4h5	Shelbyville (C)	1955	drl-GP	54	30-12	15	12	--	560	13.2	10.3	500	48.5	26	Sand & gravel	Layne-Western
26.6e	Shelbyville (C)	1954	drl	60	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 15-60	Layne-Western
26.8h	Shelbyville (C)	1954	drl	65	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 10-65	Layne-Western
27.6h	Shelbyville (C)	1954	drl	64	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 15-64	Layne-Western
28.1a	Arch Tallman	1940	drl	67	--	--	--	--	605	--	--	--	--	--	*Sandstone, 49-57	Baker
30	Pat Mathues	1950	drl	75	7	--	--	--	600	13.5	51.5	0.5	--	--	Sandstone, 61-63	Baker
30	Pat Mathues	1950	drl	18	--	--	--	--	600	--	--	--	--	--	*	Baker
32.4a	Homer McDonald**	-1927	dug	35	--	--	--	--	640	--	--	--	--	--	Sand & gravel, at 35	--
35.1c	Don Rogers	1965	drl	100	--	--	--	--	580	--	--	--	--	--	*	Baker
36.8g1	Haufland Sand & Gravel Co.	1950	drl	46	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 5-38	Baker
36.8g2	Haufland Sand & Gravel Co.	1950	drl	42	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 5-39	Baker
36.8g3	Haufland Sand & Gravel Co.	1950	drl	7	--	--	--	--	540	--	--	--	--	--	*	Baker
36.8g4	Haufland Sand & Gravel Co.	1950	drl	28	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 4-28	Baker
36.8g5	Haufland Sand & Gravel Co.	1950	drl	37	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 3-36	Baker
T11N, R4E																
--	Don W. Anderson	-1919	dug	14	--	--	--	--	--	--	--	--	--	--	--	--
--	E. H. Harwood	-1915	drl	30	--	--	--	--	--	--	--	--	--	--	--	--
--	C. H. Hulick	-1913	bor	70	2	--	--	--	--	--	--	--	--	--	--	--
--	Lithia Springs Park	-1898	spring	6	--	--	--	--	--	1	--	--	--	--	--	--
--	Lithia Springs Park	-1898	spring	6	--	--	--	--	--	1	--	--	--	--	--	--
--	Isaac Neut	-1918	drv	23	6	--	--	--	--	--	--	--	--	--	Sand & gravel, 7-23	--
4	Arch Tallman	1937	drl	128	--	--	--	--	--	--	--	--	--	--	*Sand & gravel, 50-117	Baker
5	Shelbyville (C)	1943	drl	50	--	--	--	--	--	--	--	--	--	--	*	Baker
5	Shelbyville (C)	1943	drl	70	--	--	--	--	--	--	--	--	--	--	*	Baker
5	Shelbyville (C)	1943	drl	80	--	--	--	--	--	--	--	--	--	--	*Sand, 30-70	Baker
5	Shelbyville (C)	1943	drl	92	--	--	--	--	--	--	--	--	--	--	*Sand & gravel, 30-85	Baker
5	Shelbyville CCB.	1925	dug	15	--	--	--	--	660	--	--	--	--	--	Sand, at 15	--
5.1d	U.S. Engineers	-1912	drl	27	--	--	--	--	560	--	--	--	--	--	Sand & gravel, 12-27	--
7.8g	Shelbyville (C)	1954	drl	50	--	--	--	--	540	--	--	--	--	--	*Sand & gravel, 25-45	Stevens
8.5a	Shelbyville (C)**	1924	drl	26	8	--	--	--	550	--	--	--	--	--	Gravel	Warren
8.5b1	Shelbyville (C)**	1924	drl	26	8	--	--	--	550	--	--	--	--	--	Gravel	Warren
8.5b2	Shelbyville (C)**	1924	drl	26	8	--	--	--	550	--	--	--	--	--	Gravel	Warren
8.5b3	Shelbyville (C)**	1924	drl	26	8	--	--	--	550	--	--	--	--	--	Gravel	Warren
8.5b4	Shelbyville (C)**	1924	drl	26	8	--	--	--	550	--	--	--	--	--	Gravel	Warren
8.5b5	Shelbyville (C)**	1930	drl	23	30	--	--	--	550	--	--	--	--	--	Gravel	Thorpe
8.5b6	Shelbyville (C)**	1930	drl	24	30	--	--	--	550	--	--	--	--	--	Gravel	Thorpe
8.5b7	Shelbyville (C)**	1930	drl	24	30	--	--	--	550	--	--	--	--	--	Gravel	Thorpe
8.6a1	Shelbyville (C)**	1915	dug	26	264	--	--	--	550	--	--	--	--	--	Gravel	City
8.6a2	Shelbyville (C)**	1918	dug	26	144	--	--	--	550	--	--	--	--	--	Gravel	City
8.6a3	Shelbyville (C)**	1918	dug	26	144	--	--	--	550	--	--	--	--	--	Gravel	City
8.6a4	Shelbyville (C)**	-1918	drl	25	--	--	--	--	550	--	--	--	--	--	--	--
8.6a5	Shelbyville (C)**	1929	drl	25	26	--	--	--	550	--	--	--	--	--	Gravel	Thorpe
8.6a6	Shelbyville (C)**	1929	drl	25	26	--	--	--	550	--	--	--	--	--	Gravel	Thorpe
8.6a7	Shelbyville (C)**	1929	drl	22	26	--	--	--	550	--	--	--	--	--	Gravel	Thorpe
9.2a	Malcolm Price	1958	drl	55	6	5	6	.016	640	32	6	5	0.8	7	Sand, 49-55	Baker
9.8a	Shelbyville (C)	1954	drl	85	--	--	--	--	640	--	--	--	--	--	*	Stevens
10.5a	W. E. Kelleck	1940	drl	70	6	--	--	--	685	61	--	3	--	10	Sand, 68-70	Baker
12.2a	Floyd Hazen	-1905	dug	20	--	--	--	--	710	--	--	--	--	--	--	--
13.5h	C. H. Hulick	1939	drl	125	--	--	--	--	704	--	--	--	--	--	*Sand & gravel, 45-50; sand streaks, 85-100	Baker
13.6a	James Huffer Est.	1915	dug	23	60	--	--	--	645	8	--	7	--	1	--	Small
13.7f	Herbert Barker	--	bor	48	12	--	--	--	680	16	--	5	--	1	Sand, at 48	Flory
14.6h	Golda L. Cibak	1930	bor	82	12	--	--	--	692	42	--	7	--	5	Sand, at 82	Llay
14.8c	Rose Manning	--	dug & bor	60	48-12	--	--	--	660	40	20	6	0.3	1	Sand & gravel, at 60	--
15.3a	Daniel R. Wright	1904	bor	53	14	--	--	--	660	25	25	6	0.2	2	Sand & gravel, at 53	Abercrombie
15.8e	Lloyd E. Loflin	--	bor	46	12	--	--	--	670	25	5	5	1.0	8	Sand, at 46	--
16.6h	Lane Stewardson	--	dug	20	48-30	--	--	--	655	17	--	8	--	1	Sand & gravel, at 20	Stewardson
16.8h	Shelbyville (C)	1954	bor	65	--	--	--	--	635	--	--	--	--	1	*	Stevens
17	Donald Ingham	-1907	bor	21	12	--	--	--	--	--	--	--	--	--	Sand & gravel, at 21	--
17	Shelbyville (C)	1954	drl	88	--	--	--	--	640	--	--	--	--	--	*Sand, 5-70	Stevens
17.2h	Ethel G. Braden	--	dug	20	48	--	--	--	640	15	5	5	1.0	2	Sand & gravel, at 20	--
17.4h	Harry Riley	1940	drl	101	--	--	--	--	575	--	--	--	--	--	*Mud sand, 25-35; sandstone, 95-101	Baker
17.6f	H. Prosser	1918	bor	40	12	--	--	--	620	25	--	--	--	--	Sand, at 40	Warren
18.6g	Shelbyville (C)	1954	drl	30	--	--	--	--	580	--	--	--	--	--	*Sand & gravel, 15-30	Layne-Western

Appendix B (Continued)

Well number	Owner	Well		Screen			Land surface elev (ft above msl)	Non-pumping water level (ft)	Draw* (ft)	Pumping rate (gpm)	Observed specific capacity (gpm/ft)	Length of test (hr)	Water-bearing formation and depth (ft)	Driller		
		Year constructed	Type	Depth (ft)	Diam (in)	Length (ft)									Diam (ft)	Slot size (in)
TIAM, R2E																
--	Ralph Ayars	+1916	dug	20	--	--	--	--	--	--	--	--	--	--		
--	J. W. Dobson	+1915	dug	20	--	--	--	--	--	--	--	--	Sand, at 20	--		
--	J. W. Dobson	+1915	dug	30	--	--	--	--	--	--	--	--	--	--		
--	Moweaqua (V)**	1893	dug	--	--	--	--	--	--	--	--	--	--	--		
--	School Dist. 176	+1917	bor	28	--	--	--	--	--	--	--	--	Sand, at 28	--		
--	S. A. Thomas	+1919	bor	--	--	--	--	--	--	--	--	--	--	--		
20.7h	Marjorie Porter	--	bor	29	18	--	--	630	6.5	--	--	--	--	--		
22.1h	Ralph Clipston	1947	bor	59	12	--	--	690	--	--	--	--	--	Younger		
23.1d	Willie Bohlen	1948	drl	115	6-4	3	4	.030	705	31	36	3	0.1	4	Sand & gravel, 98-114	Baker
24.8a	William Bunning	1941	drl	156	5	1.5	5	slotted pipe	715	48	97	1	--	5	Sand, 80-82, 154-156	Baker
27.1e	Opal DeWalt	--	dug	22	42	--	--	670	4.5	--	--	--	--	--	--	
29.8c	Nathan Smith	+1912	bor	30	12	--	--	635	--	--	--	--	--	Sand, 24-30	--	
30	Whitworth Farm	+1906	drv	34	2	--	--	--	--	12	--	--	--	Sand & gravel, 30-34	--	
30	Dr. M. H. Duckwall	+1906	dug	20	--	--	--	--	--	--	--	--	--	--	--	
30	Dr. M. H. Duckwall	+1906	dug	30	--	--	--	--	--	--	--	--	--	--	--	
31	Dr. M. H. Duckwall**	1906	dug	17	--	--	--	--	--	--	--	--	--	--	--	
34.7f1	Gerald C. Drew	1963	drl	100	6	2	6	.100	635	--	--	2	--	15	Sand & gravel, 70-72	Baker
34.7f2	Gerald C. Drew	1963	drl	52	6	10	6	.040	635	1	43	10	0.2	48	Sand & gravel, 45-52	Baker
TIAM, R3E																
19.1a	Earl Otto	1911	drl	108	2	3	2	--	730	--	--	--	--	--	Sand, at 108	Payne
19.5h	Rosalie M. Weck	1918	drl	121	2	3	2	--	710	--	--	--	--	--	Sand, at 121	Payne
20.2b	T. Q. Sanner	1917	drl	95	2	3	2	--	723	--	--	--	--	--	Sand, at 95	Payne
20.8h1	T. Q. Sanner**	+1913	dug	28	--	--	--	--	723	--	--	--	--	--	--	--
20.8h2	T. Q. Sanner	1916	drl	117	2	3	2	--	723	--	--	--	--	--	Sand, at 117	Payne
21.3e	M. A. Sanner & B. S. Clark	1925	bor	26	12	--	--	--	720	12.2	4.1	7	1.7	0.5	Sand, at 26	--
21.8a	Earl Snell	1900	dug	20	36	--	--	--	725	15.3	2.3	7	3.1	0.25	Sand, at 20	--
22.1g	Mrs. Leslie Daniels	1930	bor	74	12	--	--	--	700	11.3	2.8	7.3	2.6	0.5	Sand & gravel, at 74	Younger
22.6b	Lina H. Orris	1900	drl	108	2	--	--	--	715	45.4	--	--	--	--	--	--
23.3b	Daniel Sanner	1900	dug & drl	107	36-2	--	--	--	690	--	--	--	--	--	Sand, at 107	--
23.8e	Cora Sanner	1930	bor	70	12	--	--	--	700	8.3	2.3	9.7	4.2	0.5	Sand & gravel, at 70	Younger
24.2h	J. A. McReynolds	1914	bor	65	12	--	--	--	684	21.7	3.2	7	2.2	0.5	Sand & gravel, 64-65	McReynolds
24.8f	Jean Shelton	1900	bor	60	12	--	--	--	680	9.3	2.5	7	2.8	0.5	Sand & gravel, at 60	--
25.1e	Lloyd Younger	+1900	bor	35	12	--	--	--	672	15.3	5.8	9	1.5	0.25	--	--
25.2a	Willie Bohlen	1905	bor	90	12	--	--	--	675	6	5.5	9	1.6	0.5	Sand, at 90	--
26.1a	J. E. Gorman Est.	1900	dug	25	32	--	--	--	688	24.2	--	--	--	--	--	--
26.8c	James Weakley	1911	bor	86	24-18	--	--	--	700	7	--	--	--	--	Sand, at 86	Cox
27.1d1	Naomi Coultas	1922	bor	99	12	--	--	--	702	28.8	7.3	9	1.2	0.2	Sand & gravel, at 99	Younger
27.1d2	Naomi Coultas	1952	drl	102	6	10	6	.035	702	13.5	29.5	7	0.2	24	Sand & gravel, 91-102	Baker
27.6a	I. M. Baird & L. Cook	1917	drl	106	2	3	2	--	713	--	--	--	--	--	Sand, at 106	Payne
28.1c	West Center School	1910	dug	21	24	--	--	--	715	12.3	3.3	7	2.2	0.25	--	--
28.4h	O. L. Eliss	1915	drl	137	2	3	2	--	722	--	--	--	--	--	Sand, at 137	Payne
29.6a	Sanders Bros.	1916	drl	104	2	3	2	--	730	--	--	--	--	--	Sand, at 104	Payne
29.8e	H. Cooper	1929	drl	120	2	3	2	--	730	--	--	--	--	--	Sand, at 120	Payne
30.3h	Cora Williams	1917	drl	104	2	3	2	--	732	--	--	--	--	--	Sand, at 104	Payne
30.8a	Cora Williams	1911	drl	111	2	3	2	--	724	--	--	--	--	--	Sand, at 111	Payne
30.8f1	Noble Elmers	1956	drl	116	6	5	6	.012	734	52.5	44.5	5	0.1	5	Sand, 112-116	Baker
30.8f2	Noble Elmers	1957	drl	148	6	3	6	.014	734	52.5	92.5	2	--	3	Mud sand, 95-123; sand & gravel, 147-148	Baker
31.1b	Verna King	1916	drl	120	2	3	2	--	732	--	--	--	--	--	Sand, at 120	Payne
31.1h	Ervil Pierce	1944	drl	108	4	--	--	--	727	16	25	5	0.2	1	Sand & gravel, 97-99, 106-108	Woolten
31.2e	Glen Humphrey	1917	drl	100	2	3	2	--	670	--	--	--	--	--	Sand, at 100	Payne
32.2a	Florence Alward	1905	bor	108	12	--	--	--	730	18.2	5.3	7	1.3	0.5	Sand, at 108	--
32.6h	George Elliott	--	bor	100	--	--	--	--	740	--	--	--	--	--	--	--
32.8g	Irene Goodwin	1913	drl	121	2	3	2	--	730	50	--	--	--	--	Sand, at 121	Payne
33.1g	H. L. Settle	1925	bor	88	12	--	--	--	712	31.8	2.4	9	3.7	0.5	Sand, at 88	Younger
33.4h	Guy E. Cox	1916	drl	106	2	--	--	--	723	--	--	--	--	--	Sand, at 106	Payne
33.8g	Alice Morton	1932	drl	112	2	3	2	--	730	--	--	--	--	--	Sand, at 112	Payne
34.6a	Paul Kroenlein	1925	bor	85	12	--	--	--	710	8.3	2.9	7	2.4	0.5	Sand, at 85	Younger
35.1e	D. E. Baird	1900	dug	65	24	--	--	--	682	8.5	--	--	--	--	Sand & gravel, at 65	--
35.6a	Leslie Moss	1917	drl	120	2	8	2	--	690	--	--	--	--	--	Sand, at 120	Payne
36.1d	Public Hwy.	--	dug & drl	53	20	--	--	--	670	3.9	1.7	9.7	5.8	0.5	--	Stewart
36.8e	D. E. Baird	1915	bor	66	12	--	--	--	680	5.3	3.3	7	2.2	0.5	Sand & gravel, at 66	Younger

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